

Article

# Comparative Assessment of Spatial Variability and Trends of Flows and Sediments under the Impact of Climate Change in the Upper Indus Basin

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**Abstract:** Extensive research of the variability of flows under the impact of climate change has been conducted for the Upper Indus Basin (UIB). However, limited literature is available on the spatial distribution and trends of suspended sediment concentrations (SSC) in the sub-basins of UIB. This study covers the comparative assessment of flows and SSC trends measured at 13 stations in the UIB along with the variability of precipitation and temperatures possibly due to climate change for the past three decades. In the course of this period, the country's largest reservoir, Tarbela, on the Indus River was depleted rapidly due to heavy sediment influx from the UIB. Sediment management of existing storage and future planned hydraulic structures (to tap 30,000 MW in the region) depends on the correct assessment of SSC, their variation patterns, and trends. In this study, the SSC trends are determined along with trends of discharges, precipitation, and temperatures using the non-parametric Mann–Kendall test and Sen's slope estimator. The results reveal that the annual flows and SSC are in a balanced state for the Indus River at Besham Qila, whereas the SSC are significantly reduced ranging from 18.56%–28.20% per decade in the rivers of Gilgit at Alam Bridge, Indus at Kachura, and Brandu at Daggar. The SSC significantly increase ranging from 20.08%–40.72% per decade in the winter together with a significant increase of average air temperature. During summers, the SSC are decreased significantly ranging from 18.63%–27.79% per decade along with flows in the Hindukush and Western–Karakorum regions, which is partly due to the Karakorum climate anomaly, and in rainfall-dominated basins due to rainfall reduction. In Himalayan regions, the SSC are generally increased slightly during summers. These findings will be helpful for understanding the sediment trends associated with flow, precipitation, and temperature variations, and may be used for the operational management of current reservoirs and the design of several hydroelectric power plants that are planned for construction in the UIB.

**Keywords:** suspended sediment concentrations; Upper Indus Basin; Mann–Kendall trend test; Sen's slope estimator; Gilgit basin; Karakorum climate anomaly; Karakorum-Hindukush-Himalaya

## 1. Introduction

Erosion from the relatively younger mountain ranges of the Hindukush, Karakorum, and Himalaya yield a huge amount of sediment that enters the Upper Indus Basin (UIB) and its storage reservoirs, such as the Tarbela reservoir. Average sediment loads of 250 million tons per year are produced in the Indus River by these three mountain ranges, making it one of the world's largest sediment-laden

ivers [1]. Sediment erosion is a rather complex phenomenon that primarily requires determination of parameters contained in universal soil loss equation. Recent research has covered the impacts of temperature-driven snow melt processes and patterns and climate change scenarios on the erosion rate in various catchments [2,3]. Erosion and sediment transport largely depend on hydrological and hydraulic parameters that are largely driven and/or altered by climate forces. Knowledge about the distribution of hydro-meteorological parameters and sediment yields and their magnitude variations over time is crucial to the partial validation of the impacts of climate change. The assessment and quantification of variations of fluvial sediment fluxes and their spatiotemporal patterns are important for a better understanding of river geomorphology as well as the planning, design, and operation of hydraulic structures on the river. The variation of sediment fluxes affects the depletion of reservoirs, channel erosion and deposition, abrasion of hydropower turbines, pollution of river ecosystems, and the operation and management of hydropower as well as storage reservoirs. Assessment of spatio-temporal variation of suspended sediment concentrations (SSC) or yields due to precipitation, temperature, and flow process variations are vital to adapt watershed management practices and to better understand the landslide activities in the basin.

Pakistan's economy is based on agriculture. Agricultural production contributes 18.9% of the gross domestic product (GDP) in Pakistan and makes up the major part of the country's annual exports. Pakistan is among the top ten world water stressed countries due to the vulnerability of climate changes, reduction of storages, increase of its population, and increase of water demands [4,5]. The Tarbela reservoir, which lies in the UIB, has lost 35% of its capacity since 1974 due to reservoir sedimentations [6]. This rapid depletion of storage reservoirs due to siltation not only affects the country's economy, but also endangers food security and leads to cross-border water conflicts. Post construction measures against siltation in reservoirs are indispensable, but sometimes neither technically nor financially feasible. Sedimentation also adversely affects the ecosystems of both rivers and reservoirs. To enhance the sustainability and environmental compatibility of existing reservoirs and hydropower projects planned in the UIB, the impact of climate change on both flows and suspended sediments concentrations (SSC) has to be assessed. This assessment may improve the understanding of the effects of the Karakorum climate anomaly [7,8] on the seasonal and annual trends of flows and sediment yields in the UIB. This could also be beneficial for the management of sediments. This study will also help to understand the effect of different hydrological processes, such as snow cover, snow melt, ice melt, and rainfall, on erosion and sediment yield in sub-basins of the UIB dominated by snow and ice melting.

Many techniques, statistical methods, and approaches have been used by a lot of researchers for assessing changes in climate, flows, and SSC. The moving *t*-test, Yamonato method, Carner method, order clustering approach, and filter test method were used in the past [9,10]. The parametric regression approach was applied to determine the discharge–total suspended solids relationships (Q–TSS)/SSC vs. time–discharge (T–Q) for the trend analysis of sediments [11]. In 1945, Mann and Kendall suggested the nonparametric test method for time series analysis [12]. Gerstengarbe and Werner developed the Mann–Kendall test to analyze and detect trends [13]. Furthermore, the rank sum test [14], two-sample *t*-test [15], and Pettit test [16] were used for trend analysis in many studies. In a recent study for the Loess Plateau [17], trend analysis was carried out as well, because the plateau contributes 80% of the sediment belonging to the Yellow River. For large basins like the Indus basin, the analysis and budgeting of sediments and their patterns require huge resources and a large data collection expenditure. A few years ago, sediment studies were carried out [18–20] on the UIB. These studies covered the sediment patterns, sediment distribution, and sediment budgeting on the basin scale. Sediment transport is a highly nonlinear and complex phenomenon due to multiple factors controlling and affecting the erosion process, such as channel erosion, gully erosion, snow and glacier melts erosion etc. To characterize the highly nonlinear processes of sediment erosion, flow, and climate change, most researchers have used the non-parametric test. It is the most powerful tool to assess the temporal and spatial dynamics of sediment fluxes, hydrological variables, and climate. For the trends of nonlinear processes, such as flows, climate, and sediment yields, the Mann–Kendall trend test is

used for this study. It is a robust method and has the capacity to detect the trends of time series without considering the normal distribution of the input datasets. In a regional study [21], the non-parametric Mann–Kendall test was used to detect the trends of flows and sediment discharges in the Yellow River. Similarly, the non-parametric Mann–Kendall test was applied for the trend analysis of sediment yields based on data from a few stations in the UIB in two recent studies [22,23]. According to the experience gained by many researchers, however, the Mann–Kendall test alone, and without considering the serial/auto correlations, is insufficient to detect trends. Therefore, the trend-free pre-whitening (TFPW) approach, along with the Mann–Kendall trend test, is used here to detect the trends of flows, SSC, and climatic parameters. To further quantify the magnitudes of the trends of hydro-climatic time series, Sen’s slope estimator is used after the trend-free pre-whitening (TFPW) and Mann–Kendall trend analysis. Several hydropower stations and storage reservoirs are planned to be constructed in the UIB according to the 2025–2030 vision of the Water and Power Development Authority (WAPDA). The comprehensive assessment of the trends of flows and suspended sediment concentrations (SSC) under the impact of climate change for all 13 gauging stations in the Hindukush, Karakorum, and Himalaya mountain ranges presented in this study will be particularly helpful for policymakers, hydraulic engineers, and water resources managers. Therefore, the objectives of this study are: (1) the detection and assessment of trends in different temporal intervals ranging from monthly to seasonal and annual basis for flows and SSC of the 13 hydrological stations in the UIB since 1980–2010; (2) development of seasonal and annual trend maps for representing the spatial changes in % for flows and SSC per decade in the UIB; (3) extracting the missing year’s climatic information for sparsely gauged climatic stations using the linear scaling bias correction method along with the station-recorded data and satellite grid point data; (4) assessment of the impact of changes in the trends of mean monthly rainfalls of the basins and average monthly temperatures, adjusted for lapse rate per elevation zone, on the monthly flows and SSC of the UIB; (5) assessing the significant changes in monthly, seasonal, and annual flows and SSC, which are partly due to climate change, in past 30 years for selected catchments with major SSC contributions.

## 2. Materials

### 2.1. Study Area

The Indus River is among the largest rivers in South Asia with its total length of 2880 km [24,25]. It has a total catchment area of 970,468 km<sup>2</sup> with about 56% (529,134 km<sup>2</sup>) of its area lying within the territory of Pakistan. The remaining catchment area is distributed between China, India, and Afghanistan. It also has the world’s largest irrigation networks covering an irrigated agricultural area of 181,000 km<sup>2</sup> [26,27]. Moreover, the Indus basin is an economic source of energy. Hydropower produced here has a share of 29% in the total national power generation capacity of Pakistan [27].

This study focuses on the Upper Indus Basin (UIB) upstream of the Tarbela dam with a total catchment area of 172,000 km<sup>2</sup> and a length of 1125 km [18,28]. The watershed area of the UIB upstream of the Tarbela reservoir is situated at 31°–37° E and 72°–82° N. The watershed elevation of the UIB above the Tarbela dam ranges from 360 to 8572 m above mean sea level [28]. The detailed characteristics of the gauging stations measuring flows and SSC at 13 locations in the UIB are given in the Table 1. Four gauging stations are located on the main arm of the Upper Indus, eight on the tributaries, and one station, Alam Bridge, at the confluence of two tributaries.

Figure 1 shows the details of the catchment area, the sub-basins, and the tributaries of the UIB. It starts from the frozen planes of the Tibetan Plateau in China. Then, it enters from the south eastern direction in a well-defined valley and flows parallel to the geological fault line. After crossing the fluvial plains 45 km upstream of Skardu town at an altitude of 2469 m, it joins Shyok River with 30% of glacier drainage area. Here, the valley widens and mostly consists of sediments in the form of glacial deposits down to Skardu town. Upstream of Skardu in the Deosai plains, the Kharmong River from

the central Himalayas lying in Pakistan joins the main (Indus) River with a total glacier-covered area of 4%. Kharhong River is dominated by a snow and rainfall flow regime.

**Table 1.** Geographical characteristics of the stream gauging sites measuring flows and suspended sediment concentrations (SSC) in the Upper Indus Basin (UIB).

Ser #	Discharge Gauging Stations	Region	Catchment Area km <sup>2</sup>	Elevation (m)	Elevation Range (m)	% of Glacier Cover
1	Hunza at Dainyor	Western Karakorum	13,157	1350	1426–7860	35
2	Gilgit at Gilgit	Hindukush	12,095	1430	1454–7048	10
3	Gilgit at Alam Bridge	Hunza + Gilgit	26,159	1280	1266–7843	25
4	Indus at Kharhong	Central Himalayas	67,858	2542	2478–7036	4
5	Shyok at Yogu	Eastern Karakorum	33,670	2469	2397–7553	30
6	Shigar at Shigar	Central Karakorum	6610	2438	2191–7793	39
7	Indus at Kachura	Main River UIB	112,665	2341	2107–7801	12
8	Astore at Doyian	Northwestern Himalayas	4040	1583	1580–8058	8
9	Indus at Partab Bridge/Bunji	Main River UIB	142,709	1228	1242–7889	11
10	Indus at Shatial Br.	Main River UIB	150,220	1040	906–8118	12
11	Indus at Besham Qila	Main River UIB	162,393	580	561–8118	13
12	Lower UIB at Daggar	Lower western UIB	598	700	685–2801	0
13	Lower UIB at Karora	Lower western UIB	635	880	893–4439	1

Near the town of Skardu, the Shigar River with 39% of glacier-covered area from the Himalayas drains into the main (Indus) River at an altitude of 2438 m. Downstream of Skardu town, the main River flows towards the Northwest up to Kachura Lake along the steep and narrow valley of the main River. The Hunza River runs from North to South and Gilgit River runs from the Northwest to Southeast before confluence at Alam Bridge just downstream of Gilgit town. Hunza Basin borders with Shigar Basin on the right side in the direction of flow. Hunza and Gilgit Rivers are snow- and glacier-dominated basins lying in the Western Karakorum and Hindukush mountain ranges with 39% and 10% of glacier-covered area, respectively. Astore River also joins the main River. It comes from the western Himalayas with 8% of glacier-covered area north of the Nanga Parbat massif. Downstream of Alam Bridge, the confluence of Gilgit and Hunza Rivers join the main Indus River at Bunji/Partab Bridge in Jaglot town. Jaglot is a monumental place, where all three great mountain ranges, the Himalayas, Karakorum, and Hindukush, meet. The lower part of the UIB starts at Bunji and extends up to Tarbela dam. The main river widens between Alam Bridge and Chillas and contains sediments in the form of glacier deposits. The Indus River continues its journey towards Shatial Bridge. From Shatial Bridge, the main river flows further towards the South in a steep and narrow cross section and joins Besham Qila, approx. 45 km upstream of the Tarbela reservoir. Near Besham Qila, some small and large rivers also join the Indus from both the left and the right side. The right tributaries joining the main Indus River near Besham Qila are Brandu, Gorbant, Khan Khawar, Duber Khwar, Keyal Khwar, Kandia, and Tangir. The left tributaries joining the main Indus River near Besham Qila are Siran, Chor Nullah, Spat Gah, Allai Khwar, Gunar Gah, Thor Gah, and Butto Gah.

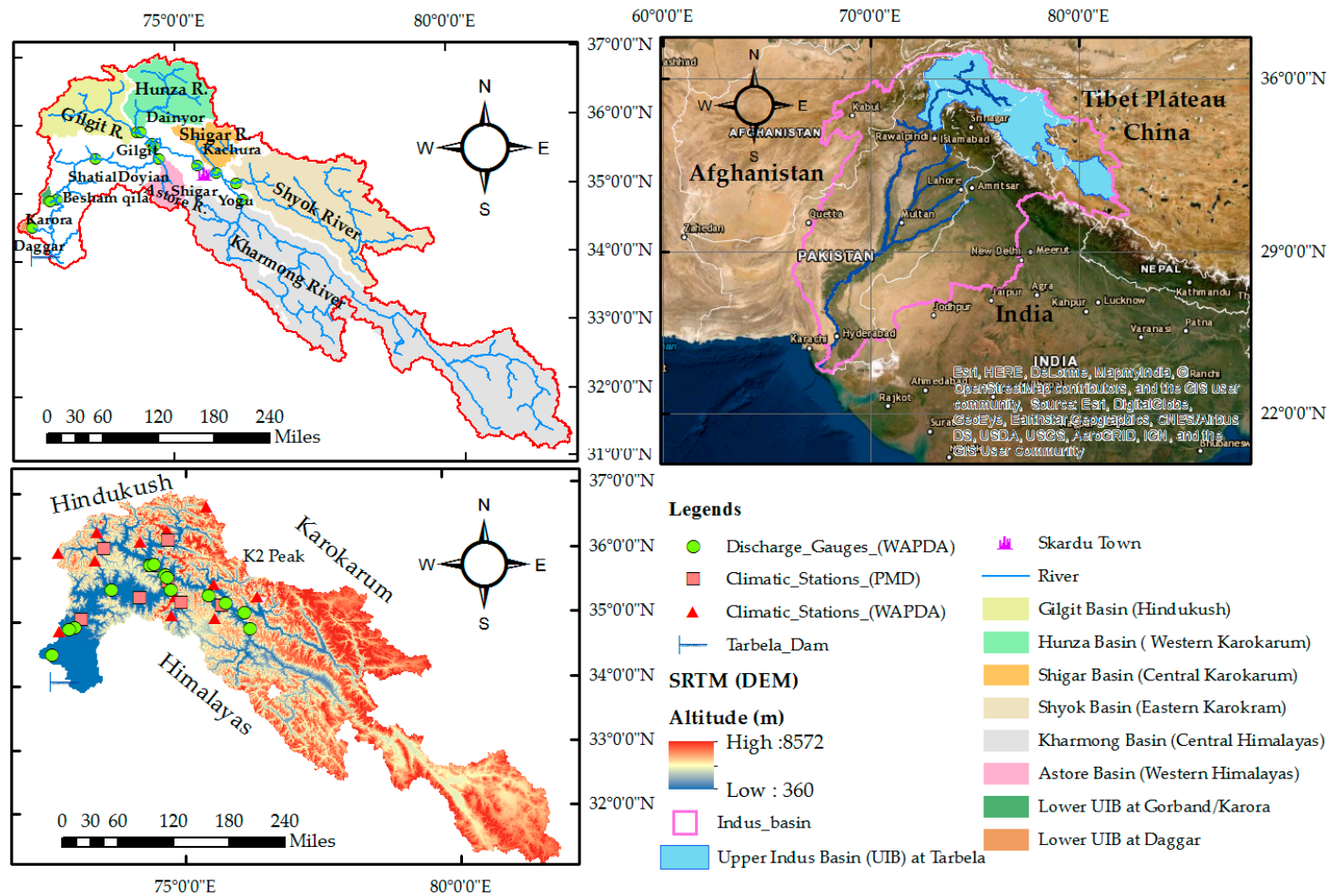


Figure 1. The map of the study area shows the locations of stream gauges, climatic stations, and sub-basins contributing to the Upper Indus Basin (UIB).

The climate of the UIB comprises two climatic fronts. One is the monsoon pattern of rainfall that originates from the Bay of Bengal and Arabian Sea. The second is the westerlies climatic front originating from the Caspian Sea and Mediterranean region during winters and springs. Most of the annual precipitation in the UIB falls in the form of snow during the winter and spring season due to westerlies [29,30]. Mean annual precipitation in Gilgit ranges from 150 mm at lower elevations to 1800 mm in the snow accumulation zone in Bagrot valley [31]. In the Western Karakorum over the Batura glaciers, annual precipitation accounts for more than 1000 mm [32]. In the Central Karakorum, annual precipitation ranges from 1000 mm to more than 2000 mm over the Hispar and Biafo glaciers [33]. In the Central Karakorum, 67% of annual precipitation fall during winter is due to the westerly front, with the remaining 33% of annual rainfall during the summer season [7,34,35]. In the lower UIB, annual rainfall ranges from 1000 mm to 1500 mm between Dasu and Besham Qila [23]. In the UIB, the maximum mean basin rainfall occurs during the spring season, with April being the month of highest rainfall [36].

In the UIB the upper glaciered has very low population density (<50 people km<sup>2</sup>) in comparison to the lower monsoon dominated region. The upper region of UIB has an estimated population of 1.5 Million. The lower region of UIB is vice versa. The upper region of UIB also has high percent of snow/ice cover, at up to 40% in Hunza, Gilgit, Shigar, and Shyok river basins. This snow/ice cover emerges as a single dominated land cover factor which explains 73.4% of the variance in sediment yields for whole UIB [37].

The trends induced by climate change on the discharges of various sub-catchments have been widely reported by many researchers [7,8,38,39], however, scientific literature on the spatio-temporal variation of SSC is far less extensive. Ali and De Boer [18] analyzed the spatial distribution of sediment in the UIB by dividing the region into three zones: upper snow- and glacial melt-dominated zone, middle snow-/glacial melt-, and rainfall-dominated mixed zone, and lower rainfall-dominated zone. Ateeq-Ur-Rehman et al. and Tarar et al. [22,23] identified the trends of SSC by using the data of two and four gauging stations of the UIB, respectively. Meanwhile, the former used the SSC of Hunza basin at Danyor, Shyok basin at Yogu, Upper Indus at Bunji/Partab Bridge, and Upper Indus at Besham Qila. They have used data from the gauge of Partab Bridge for the melt portion of the UIB and from Besham Qila to represent the load of the entire UIB. These studies found a significant inter-annual shift of SSC with the balanced annual sediment yield. However, an assessment of the spatio-temporal variation of SSC and discharges in relation to relevant hydrometeorological parameters and variables for all the sub-basins of the entire UIB has not yet been made. This study therefore covers the trends of spatial variability of SSC and discharges in the UIB using the data of 13 gauging stations located in the Karakorum, Himalayan, and Hindukush regions that represent almost the entire UIB with its eight sub-catchments (Table 1). Moreover, the previous studies used daily sediment values derived from discharge-sediment rating curves (SRC) and artificial neural networks (ANN) [22,23]. The SRC and databased models give highly overestimated and underestimated values and sometimes negative values, which are physically not possible. Especially, for the highly glacier and snow melt basins, the SRC are not reliable. In comparison to previous researches in the region that mainly relied on discharge-sediment rating curves (SRC) and artificial neural networks (ANN) for SSC, the mean of monthly sediments was taken by averaging the daily SSC gauged data. Thereby, SRC values were only used in those months where no information of sediment concentrations was available. This study is also one of the pioneer researches that used corrected estimates of rainfall and temperature for the assessment of comparative trends of precipitation, temperature, and SSC in the UIB [40,41].

## 2.2. Data Collection

For the present study, the data of 13 stream gauges in the UIB for the period 1981–2010 were collected within the Surface Water Hydrology Project (SWHP) of the Water and Power Development Authority (WAPDA). These stations measure flows on a daily basis and SSC on intermittent days.

A rating curve was used to find the missing values of SSC for monthly, seasonal, and annual trend analysis. The list and characteristics of these stations are given in Table 2.

Hunza is second largest glacier-dominated basin of the UIB in the Western Karakorum. Here, data are available for the period of 1981–2010. The Hunza River generates 12% and 12.5% of the Besham Qila flows and suspended sediment loads (SSL), respectively. The Gilgit station at Gilgit lies in the Hindukush mountain range draining the Gilgit river basin. Here, data are available for 1981–2010. Its flows and SSL are 12% and 3.9% of the Besham Qila flows and SSL, respectively. Gilgit station is important due to the contribution of snow melt, ice melt, and rainfall to flows and SSL. The Alam Bridge station is located at the confluence of Gilgit River and Hunza River, generating 26% and 18.2% of Besham Qila flows and SSL, respectively. The major influx of sediment from the Hunza River raises the SSL in the combined flow of Gilgit and Hunza at Alam Bridge.

The gauge at Kharmonig in the Central Himalayas was installed in 1982. For this reason, data of flows and SSC are available for the period 1983–2010. The main River comes from the Deosai plain and is dominated by snow melt and rainfall. Its flows and SSL are 18% and 7% of the Besham Qila flows and SSL, respectively. The Yogo station on Shyok River lies in the eastern Karakorum. Shyok River originates from the Tibetan Plateau in China and has the third largest glacier dominance in the UIB. Data are available for the period 1981–2010. The station measures 16% and 20.6% of flows and SSL, respectively, of Besham Qila. The Shigar station in the Central Karakorum generates most of the flows and SSL due to largest glacier dominance in the UIB. The station was installed in 1982, but data are available from 1985–1998 only. After 1998, Shigar station was not in operation. Its flows and SSL are 9% and 15.4% of the Besham Qila flows and SSL, respectively.

The Indus at Kachura contributes 48% and 52.1% to the Besham Qila flows and SSL, respectively. Data are available for the long term from 1981–2010. The Bunji/Partab Bridge station on the main River has data from 1981 to June 2010. Due to the 2010 flood, the gauging site was destroyed, and the data of the remaining six months of 2010 are lacking. The Bunji station generates 71% and 97.6% of the flows and SSL, respectively, of Besham Qila. The river here contains the water and sediments from glacier- and snow-dominated upstream basins, i.e., Hunza and Gilgit on upper Northeastern side and from the parallel basins of Shigar, Shyok, and Kharmonig on the lower Northeastern end. The Doyian station on Astore River also is located in a snow melt- and rainfall-dominated basin in the Western Himalayas. It generates 6% and 1.2% of the flows and SSL, respectively, of Besham Qila. Again, data are available for the entire period of 1981–2010.

The station Shatial Bridge was installed in 1983 and measures 87% of flows and surprisingly 70.0% of SSL of Besham Qila despite its location downstream of the Bunji station. Hydrological data of Shatial are available from 1983–2010. The Dagggar station on Brandu River is the right tributary of the Lower UIB near Besham Qila. It is a rainfall-dominated basin contributing 0.4% and 0.4% of discharges and SSL, respectively, of Besham Qila. The Karora station of the Gorband basin is also located on a right tributary near Besham Qila and contributes 0.7% and 0.2% of flows and SSC, respectively, of Besham Qila. Both Dagggar and Karora measured data in the period from 1981–2010.

**Table 2.** Hydrological characteristics of discharge gauging stations in the Upper Indus Basin (UIB).

Ser #	Discharge Gauging Stations	Region	Lat. (°)	Long. (°)	Period	Flows (m <sup>3</sup> /sec)	SSC (mg/L)	% of UIB Flows	% of UIB SSC	SSL (Million Tons/Year)	% of UIB SSL
1	Hunza at Dainyor	Western Karakorum	35.92	74.37	1981–2010	297	1091	12	125	33	12.5
2	Gilgit at Gilgit	Hindukush	35.92	74.3	1981–2010	293	442	12	51	10	3.9
3	Gilgit at Alam Bridge	Hunza + Gilgit	35.76	74.59	1981–2010	632	795	26	91	48	18.2
4	Indus at Khar Mong	Central Himalayas	34.93	76.21	1983–2010	453	669	18	76	19	7.0
5	Shyok at Yogu	Eastern Karakorum	35.18	76.1	1981–2010	390	1305	16	150	54	20.6
6	Shigar at Shigar	Central Karakorum	35.33	75.75	1985–1998	223	2395	9	275	40	15.4
7	Indus at Kachura	Main River UIB	35.45	75.41	1981–2010	1192	1171	48	134	137	52.1
8	Astore at Doyian	North Western Himalayas	35.54	74.7	1981–2010	140	295	6	34	3	1.2
9	Indus at Partab Bridge/Bunji	Main River UIB	35.73	74.62	1981–2010	1757	1019	71	117	257	97.6
10	Indus at Shatial Br.	Main River UIB	35.53	73.56	1983–2010	2167	919	87	106	185	0.70
11	Indus at Besham Qila	Main River UIB	34.92	72.88	1981–2010	2479	870	100	100	263	100.0
12	Lower UIB at Daggar	Lower western UIB	34.49	75.46	1981–2010	9.9	547	0.4	63	1	0.4
13	Lower UIB at Karora	Lower western UIB	34.89	72.76	1981–2010	18	355	0.7	41	0.4	0.2

Note: The flows and SSC are instantaneous values for the time the water samples are taken by sediment sampler.



The SSC data given in % of Besham Qila benchmarks do not only reflect the availability of suspended sediment, but also indicate the transport capacity. For example, Shigar at Shigar, Shyok at Yogu, and Hunza at Dainyor report 2.75, 1.50, and 1.25 times the suspended sediment concentration of Besham Qila, respectively. Similarly, the Indus at Kachura (134%), Partab Bridge (117%), and Shatial Bridge (106%) is found to contain more suspended sediment than recorded at Besham Qila, which indicates that with the increasing flows from Kachura to Besham Qila, the suspended sediment does not match with the increasing rate of discharges. However, the SSL data reveal that from Kachura to Bunji, the sediment increases from 137 to 257 million tons, but decreases to 185 million tons at Shatial, and again rises to 263 Million tons at Besham Qila. Provided that the data are correct, the decrease of sediment load from Bunji to Shatial despite the increasing discharges means that sediment is being deposited between Bunji and Shatial and there may be either channel bed erosion or an addition of suspended sediments from the lateral tributaries between Shatial and Besham Qila, which makes the SSL more than double at the tail end.

The daily climatic data of rainfall and temperatures from 18 meteorological stations in the UIB were also collected in this study. The meteorological stations of Skardu, Gilgit, Gupis, Astore, Bunji, and Chillas are operated by the Pakistan Meteorological Department (PMD).

The remaining climatic stations under the supervision of the Water and Power Development Authority (WAPDA) have been recording data since 1995. Table 3 presents the details of these climatic stations. In addition, the grid data of precipitation and temperature of  $5 \times 5$  km resolution determined in the HI-AWARE project for the Indus, Ganges, and Brahmaputra river basins were collected [40,41]. The Shuttle Radar Topography Mission's (SRTM) digital elevation model (DEM) data of 30-m resolution were applied as well. The mean basin precipitation data were extracted from the corrected rainfall data of the HI-AWARE project using DEM. The temperature data of higher elevated gauges for the missing period from 1981–1994 were extracted from grid data and corrected with station observations by using the linear biased scaling method. Estimation of glacier areas was based on the glacier polygons of the Global Land Ice Measurement (GLIMS) database. In this study, DEM is used to estimate the glacier area for each sub-basin from downloaded GLIMS polygons [42].

**Table 3.** Characteristics of climatic gauging stations in the Upper Indus Basin (UIB).

Discharge Outlet	Region	Climatic Stations	Lat. (°)	Long. (°)	Average Altitude (m)	Period
Hunza at Dainyor	Western Karakorum (WK)	Naltar	36.29	74.12	2898	1995–2010
		Ziarat	36.47	74.62	3020	1995–2010
		Khunjrab	36.83	75.4	4440	1995–2010
Gilgit at Gilgit	Hindukush (HK)	Gilgit	34.92	73.34	1460	1981–2010
		Gupis	36.17	73.4	2156	1981–2010
		Ushkore	35.99	73.25	3051	1995–2010
		Yasin	36.43	73.27	3280	1995–2010
		Shendure	36.09	72.54	3712	1995–2010
Gilgit at Alam Bridge	Hunza + Gilgit	HK + WK	-	-	-	1981–2010
Indus at Kharmong	Central Himalayas (CH)	Desosai	35.09	75.54	4149	1995–2010
Shyok at Yogu	Eastern Karakorum (EK)	Hushey	35.42	76.37	3075	1995–2010
Shigar at Shigar	Central Karakorum (CK)	Shigar	35.63	75.53	2325	1995–2010
		Skardu	35.3	75.68	2210	1981–2010
Indus at Kachura	Main River UIB	EK + CH	-	-	-	1981–2010
Indus at Partab Bridge/Bunji	Main River UIB (HK + WK + CK + EK + CH)	Bunji	35.67	74.64	1372	1981–2010
Astore at Doyian	North Western Himalayas (NWH)	Astore	35.34	74.9	2168	1981–2010
		Raitu	35.14	74.73	2718	1995–2010
		Rama	35.43	74.79	3179	1995–2010

Table 3. Cont.

Discharge Outlet	Region	Climatic Stations	Lat. (°)	Long. (°)	Average Altitude (m)	Period
Indus at Shatial Br.	Main River UIB (HK + WK + CK + EK + CH)	Chillas	35.42	74.1	1250	1981–2010
Indus at Besham Qila	Main River UIB (HK + WK + CK + EK + CH)	Chillas	35.67	74.64	1250	1981–2010
Lower UIB at Daggar	Lower western UIB	Shangla	34.87	72.6	1960	1995–2010
Lower UIB at Karora	Lower western UIB	Shangla	34.87	72.6	1960	1995–2010

### 3. Methods

The data outliers in this study do not affect the analysis, as the Mann–Kendall test used for trend analysis is non-parametric. The homogeneity of the mean annual flows and SSC was checked using the standard normal homogeneity test (SNHT) and Buishand’s range (BR) test at 5% significance level for each station. The flows and suspended sediments are considered homogeneous when the critical values of SNHT and BR are  $T_0 < 9.17$ ,  $Q/\sqrt{n} < 1.27$ , and  $R/\sqrt{n} < 1.55$ .  $T_0$  is SNHT test statistics. Similarly  $Q/\sqrt{n}$  and  $R/\sqrt{n}$  are BR test statistics. Table 4 shows the results of the SNHT and BR tests.

Table 4. Characteristics of climatic gauging stations in the Upper Indus Basin (UIB).

Sr #	Station	Stream Flows (m <sup>3</sup> /sec)			Suspended Sediments (mg/L)		
		SNHT	Buishand’s Range (BR) Test		SNHT	Buishand’s Range (BR) Test	
			$T_0$	$\frac{Q}{\sqrt{n}}$		$\frac{R}{\sqrt{n}}$	$T_0$
1	Hunza at Dainyor	4.37	0.1	1.33	4.98	0.09	1.04
2	Gilgit at Gilgit	4.32	0.08	0.85	2.19	0.05	1.13
3	Gilgit at Alam Bridge	3.29	0.06	0.98	4.36	0.08	0.98
4	Indus at Kharmong	3.55	0.10	1.5	3.51	0.10	1.42
5	Shyok at Yogu	2.42	0.06	0.97	5.68	0.10	1.91
6	Shigar at Shigar	4.50	0.13	0.99	3.00	0.12	1.18
7	Indus at Kachura	1.97	0.06	1.16	7.09	0.11	1.25
8	Indus at Partab Bridge/Bunji	4.37	0.07	0.85	3.07	0.06	1.05
9	Astore at Doyian	5.92	0.07	0.85	3.83	0.08	1.08
10	Indus at Shatial Br.	1.33	0.05	0.74	5.01	0.11	1.34
11	Indus at Besham Qila	1.23	0.05	0.83	3.31	0.06	1.14
12	Lower UIB at Daggar	1.58	0.06	0.96	5.98	0.11	1.28
13	Lower UIB at Karora	12.59	0.17	1.8	11.39	0.05	0.97

Note: For homogeneous data series  $T_0 < 9.17$ ,  $\frac{Q}{\sqrt{n}} < 1.27$ , and  $\frac{R}{\sqrt{n}} < 1.55$ .

#### 3.1. Statistical Methods

The Mann–Kendall trend test is applied to detect the trends for mean values of flow, SSC, precipitation, and average temperatures on a monthly, seasonal, and annual basis. Before applying the Mann–Kendall test, the time series data were analyzed to check serial correlation. The serial correlations were eliminated from the series using the trend-free pre-whitening (TFPW) approach.

##### 3.1.1. Serial Correlation and Trend-Free Pre-Whitening (TFPW)

Serial correlation in time series is defined as correlation of a variable with itself over successive time intervals. This correlation must be removed before trend analysis [43]. Removal of this correlation is important, because it significantly affects the results of the non-parametric test during trend analysis.

Using the two-tailed test of autocorrelation coefficient ( $r_1$ ) at a 5% significance level, the time series of flows, SSC, rainfall, and temperatures were checked for serial correlation as:

$$r_1 = \frac{\sum_{i=1}^{n-1} (X_i - X^-)(X_{i+1} - X^-)}{\sum_{i=1}^n (X_i - X^-)^2} \quad (1)$$

where  $X_i$  is the value of suspended sediments at  $i$ th time interval and  $X^-$  is the mean value of sediments.  $n$  is the number of years.

The value of the autocorrelation coefficient  $r_1$  at 95% of confidence interval in the two-tailed test is given as:

$$r_1 \text{ Suspended Sediments (95\%)} = \frac{-1 \pm 1.96 \sqrt{(n-2)}}{n-1} \quad (2)$$

The trend-free pre-whitening (TFPW) approach is proposed when the  $r_1$  value of the time series lies between the upper and lower limits of the confidence interval [44].

### 3.1.2. Mann–Kendall Test

The Mann–Kendall test is most widely used for the trend analysis of hydro-climatic data in various studies [12,45–47]. To detect the statistically significant trends in the time series, the Mann–Kendall test uses two hypotheses. The null hypothesis ( $H_0$ ) is that there is no trend overtime in flows and suspended sediment data. The ( $H_1$ ) hypothesis is that the trend in the time series is increasing or decreasing.

For the Mann–Kendall test [12,45,47],  $S$  has been computed as given below:

$$S = \sum_{i=1}^{n-1} \sum_{j=i+1}^n \text{sig}(X_j - X_i) \quad (3)$$

As  $X_i$  and  $X_j$  are the data values at times  $i$  and  $j$ , while  $n$  is the length of the dataset.

$$\text{Sgn}(X_j - X_i) = \begin{cases} +1 & \text{if } (X_j - X_i) > 0 \\ 0 & \text{if } (X_j - X_i) = 0 \\ -1 & \text{if } (X_j - X_i) < 0, \end{cases} \quad (4)$$

The positive values of the calculated  $S$  indicate the positive trend and vice versa.

The variance  $S$  of data series is computed by the equation given below:

$$\text{Var}(S) = \frac{1}{18} \left[ n(n-1)(2n+5) - \sum_{p=1}^q t_p(t_p-1)(2t_p+5) \right] \quad (5)$$

where  $t_p$  is the number of ties for  $p$ th values with  $q$  being the number of tied values. After calculating the variance  $\text{Var}(S)$  of data time series, the standard  $Z$  value is calculated as

$$Z = \begin{cases} \frac{s-1}{\sqrt{\text{Var}(S)}} & \text{if } S > 0 \\ 0 & \text{if } S = 0 \\ \frac{s+1}{\sqrt{\text{Var}(S)}} & \text{if } S < 0 \end{cases} \quad (6)$$

The calculated positive value of  $Z$  indicates an increasing trend, while a negative value of  $Z$  reflects a decreasing trend.

The  $Z$  value is compared at a significant level of  $\alpha = 1\%$ ,  $5\%$ , and  $10\%$  with the normal distribution table value. The null hypothesis ( $H_0$ ) is rejected, if the  $Z$  value is greater than  $|Z_{1-\frac{\alpha}{2}}|$ ,  $|z| > |Z_{1-\frac{\alpha}{2}}|$ , where  $|Z_{1-\frac{\alpha}{2}}|$  was obtained from the standard normal distribution table. The other hypothesis  $H_1$  is accepted.

### 3.1.3. Sen's Slope Estimator

To quantify the magnitude of change of the trend in time series data, the Sen's slope estimation method [48] was used.

$$Q_i = \frac{X_j - X_k}{j - k} \quad (7)$$

where  $Q_i$  is the slope of the  $i$ th paired of data, for  $i = 1, \dots, N$

$X_j$  and  $X_k$  are the data values at the time of  $j$  and  $k$  [ $j > k$ ], respectively. The median of  $N$  values of  $Q_i$  values is the Sen's slope estimator given as

$$Q_i = \begin{cases} T_{\frac{(N+1)}{2}} & N \text{ is odd} \\ \frac{1}{2} \left( T_{\frac{N}{2}} + T_{\frac{(N+1)}{2}} \right) & N \text{ is even} \end{cases} \quad (8)$$

## 4. Results

### 4.1. Preliminary Analysis

The preliminary descriptive statistics for discharge and SSC are given in Table 5. Besham Qila station that represents the snow- and ice melt-dominated and partly rainfall-dependent basins shows the highest values of mean (Mean), minimum (Min), and maximum (Max) flows due to the fact it is the last gauging station of the UIB that is located just upstream of the Tarbela reservoir. Shatial Bridge just upstream of Besham Qila shows the highest standard deviation (Std.) of discharges, which means that the flow pattern is more erratic than at both neighboring upstream and downstream stations. Dagggar, the basin dominated by high-intensity rainfall in spring and summer, is comparatively small and reports the highest values for the coefficient of variations ( $C_v$  (%)), coefficient of skewness ( $C_s$ ), and coefficient of kurtosis ( $C_k$ ) for discharges partly because of the low mean magnitude and high discharge variations because of rainfall storms. Dagggar basin also shows the lowest minimum and maximum flows due to rainfall contribution only and the small catchment area. Gorbant at Karora shows the lowest standard deviation. Besham Qila, Partab Bridge, and Kachura on the main Indus River are found to have the lowest coefficient of variations, coefficient of skewness, and coefficient of skewness, respectively. This is due to the fact that they represent very large catchments (Table 1).

Shigar, the glacier-dominated basin, shows the highest values of mean, maximum, and standard deviation of SSC, followed by Shyok, Kachura, and Hunza. The Indus at Shatial Bridge reaches the highest values of minimum of SSC, which may possibly be due to heterogeneous spells of channel erosion of glacier deposits. During low flow season these glacier deposits might be transported by increase of winter discharges in channel. The Gorbant at Karora is found to have the highest coefficient of variations, which could be probably due to the operation of a lot of upstream micro hydropower stations installed by the community. Gilgit, the snow melt-, ice melt-, and rainfall-dominated basin, reaches the highest coefficient of skewness and coefficient of kurtosis of SSC, which may probably be due to dominant snow melts in the basin and increased sediment transport from the upstream catchment. Astore at Doyian, a snowfall-dominated basin, is found to have the lowest SSC in terms of mean, maximum, and standard deviation. Besham Qila shows the lowest coefficient of variations, which is attributed to the fact that it is last downstream station of the UIB, where river hydraulics is the predominant factor rather than upstream catchment erosions. Moreover, the representation of a larger catchment is less prone to variations by short-term local forcings. Hunza at Dainyor shows the lowest value of coefficient of skewness. Probably, more sediments are deposited in the upstream valley and hydraulics of the river with effective discharges from the glacier melts is more dominant, which ensures steady sediment transport. The Indus at Kachura shows the lowest coefficient of skewness, probably due to high flows and steady SSC contribution along narrow river valleys with less glacier deposits in the channel. It must be noted that flow doubles and SSC is reduced by almost a quarter from Kachura to Besham Qila along the main arm of the Indus River, which means that

either the lateral sub-basin downstream of Kachura yields less suspended sediment in comparison to discharges or there is an issue of sediment transport capacity in a reach between these stations.

**Table 5.** Summary of descriptive statistics of the selected discharge and sediment gauging stations.

Sr #	Gauging Stations	Mean	Min	Max	Std.	Cv (%)	Cs	Ck
<b>Discharge m<sup>3</sup>/sec</b>								
1	Hunza at Dainyor	296.75	50.73	508.21	96.59	32.55	−0.37	1.58
2	Gilgit at Gilgit	293.21	83.14	867.86	122.76	41.87	3.42	16.67
3	Gilgit at Alam Bridge	632.39	217.61	1215.64	185.44	29.32	0.58	3.12
4	Indus at Khar Mong	452.78	118.94	782.45	140.29	30.98	0.17	1.09
5	Shyok at Yogu	390.07	180.28	545.96	89.29	22.89	−0.15	−0.36
6	Shigar at Shigar	223.25	142.06	330.64	57.44	25.73	0.29	−0.65
7	Indus at Kachura	1192.19	860.16	1547.26	197.44	16.56	0.09	−0.77
8	Indus at Partab Bridge/Bunji	1756.94	865.51	2256.72	369.38	21.02	−0.81	0.3
9	Astore at Doyian	139.9	37.52	232.24	38.58	27.58	−0.22	0.94
10	Indus at Shatial	2167.32	915.99	2981.57	499.73	23.06	−0.58	0.28
11	Indus Besham Qila	2479.37	1858.22	3304.93	378.85	15.28	0.37	−0.55
12	Lower UIB at Daggar	13.84	3.81	129.3	21.85	157.9	5.23	28.41
13	Lower UIB at Karora	18.1	5.89	37.36	7.24	40	0.97	1.26
<b>Suspended Sediments mg/L</b>								
1	Hunza at Dainyor	1091.44	53.75	2244.65	568.27	52.07	0.03	−0.04
2	Gilgit at Gilgit	441.98	76.82	2169.5	410.59	92.9	2.78	9.94
3	Gilgit at Alam Bridge	795.32	114.47	1971.58	439.5	55.26	0.86	0.53
4	Indus at Khar Mong	668.9	138.28	2150.42	537.79	80.4	1.37	1.21
5	Shyok at Yogu	1305.08	128.78	5219.85	1208.13	92.57	1.79	2.96
6	Shigar at Shigar	2395.08	257.48	7571.67	1883.03	78.62	1.6	3.64
7	Indus at Kachura	1171.27	131.29	2390.97	554.23	47.32	0.36	−0.16
8	Indus at Partab Bridge/Bunji	1019.35	306.92	3294.93	658.31	64.58	1.79	4.21
9	Astore at Doyian	295.27	96.8	675.68	164.53	55.72	0.94	0.04
10	Indus at Shatial	919	352.8	1993.22	355.1	38.64	1.32	2.34
11	Indus at Besham Qila	870.71	322.08	1584.57	286.96	32.97	0.43	0.63
12	Lower UIB at Daggar	546.85	76.8	1649.01	401.73	73.46	1.43	1.85
13	Lower UIB at Karora	355.03	48	1739.6	417.81	117.68	2.17	4.54

Note: **Red** color shows highest values and **blue** color indicates the lowest values for both flows and SSC.

#### 4.2. Monthly, Seasonal, and Annual Trends of Discharges

Using longer time steps, such as decadal accumulative discharges, means to avoid inter-annual fluctuations when determining long-term trends for the Indus basin [22,49]. In this study, monthly trends of discharges (m<sup>3</sup>/sec) with 90% confidence interval are determined per decade by using the Sens's slope estimator after the Mann–Kendall trend test, as shown in Table 6. For the flows, most of the stations show increasing trends of discharges in snow- and glacier-dominated basins in the months of from December to June but decreasing discharges in July–August on higher altitudes like Hunza and Khar Mong. The Indus at Besham Qila (main UIB), Shatial Bridge (main UIB), Partab Bridge (main UIB), and Kachura (main UIB) and the Astore at Doyian (Western Himalayas) reveal increasing trends of flows per decade in most of the months between November and April. Most of the remaining glacier- and snow-dominated basins, such as Gilgit (Hindukush), Hunza (Western Karakorum), Shigar (Central Karakorum), and Shyok (Western Karakorum), show a significant increase of flows between October and March. However, the Khar Mong (Central Himalayas) exhibits a mixed trend in these months.

Table 6. Mann–Kendall trends for monthly time series data and Sen’s slope estimates.

Discharge Gauging Stations	January	February	March	April	May	June	July	August	September	October	November	December
<b>Change in Discharge m<sup>3</sup>/sec per Decade</b>												
Hunza at Dainyor	0.87	0.89	0.66	6.35	<b>−2.27</b>	12.13	<b>−28.64</b>	<b>−174.58 *</b>	5.17	1.71	5.11 *	<b>−1.68</b>
Gilgit at Gilgit	0.29	0.09	0.9	5.36	55.82 *	33.09	53.44	2.22	13.32	19.44 *	5.6	7.39 *
Gilgit at Alam Bridge	1.1	1.3	2.74	19.82 *	19.47	109.07 *	21.57	9.62	71.52	31.03 *	18.9 *	12.44
Indus at Kharmong	<b>−1.44</b>	0.34	4.88	5.49	<b>−20.82</b>	<b>−3.24</b>	<b>−167.82</b>	<b>−7.03</b>	<b>−41.29</b>	16.33	7.74	<b>−1.58</b>
Shyok at Yogu	2.13 *	2.25 *	0.09	<b>−1.59</b>	6.41	58.34	78.54	88.74	12.55	31.21 *	10.3 *	2.01
Shigar at Shigar	1.76	2.98	8.74 *	9.58	<b>−13.7</b>	176.1	432.67	376.17 *	33.78	−2.54	−1.43	<b>−1.24</b>
Indus at Kachura	4.9	10.55 *	10.42 *	2.74	43.95	78.62	279.97	78.77	<b>−175.18 *</b>	11.71	4.69	7.43
Indus at Partab Bridge/Bunji	13.8	20.27 *	24.19 *	46.93 *	149.07 *	323.34	425.22	135.27	27.06	7.72	27.44 *	14.59
Astore at Doyian	2.62 *	2.76 *	3.93 *	13.86 *	39.44 *	22.06	<b>−6.72</b>	7.77	2.01	2.66	3.57	3.57 *
Indus at Shatial Br.	44.17 *	35.96 *	36.19 *	88.64 *	239.08	174.8	173.35	234	71.2	63.88	37.11 *	43.67 *
Indus at Besham Qila	36.31 *	41.08 *	30.48 *	6.96	170.34	16.31	214.18	125	163.75	95.17 *	57.61 *	35.84 *
Lower UIB at Daggar	0.33	0.55 *	<b>−4.31 *</b>	<b>−0.2</b>	0.1	0.4 *	<b>−1.45</b>	<b>−3.25</b>	0.01	0.15	0.07	0.2
Lower UIB at Karora	<b>−0.04</b>	1.51	<b>−2.81</b>	<b>−7.01 *</b>	<b>−7.88 *</b>	<b>−3.77 *</b>	<b>−5.73 *</b>	<b>−4.88 *</b>	<b>−2.77 *</b>	<b>−0.78</b>	<b>−0.52</b>	<b>−1.44</b>
<b>Change in Suspended Sediments mg/L per Decade</b>												
Hunza at Dainyor	24.03	8.34	42.5	−91.58	3.67	738.07	<b>−765</b>	<b>−1181.35</b>	<b>−12.73</b>	<b>−35.72</b>	15	8.67 *
Gilgit at Gilgit	15	15.89 *	17.5 *	18.86	45.53	<b>−68.77</b>	<b>−245.43</b>	<b>−472.47 *</b>	<b>−24.57</b>	<b>−0.21</b>	15.38 *	17.82 *
Gilgit at Alam Bridge	25 *	16.65 *	32.81	<b>−9.45</b>	−31.76	68.82	<b>−591.67 *</b>	<b>−858.57 *</b>	<b>−115 *</b>	<b>−9.05</b>	46.82 *	36.75 *
Indus at Kharmong	<b>−56.95</b>	<b>−22.21</b>	23.93	86.05 *	62.67	376.84	<b>−110.4 *</b>	<b>−15.11</b>	<b>−167.62</b>	<b>−14.53</b>	2.09	<b>−5.62</b>
Shyok at Yogu	8.49	<b>−16.34</b>	33.98	108.57 *	43.86	<b>−458.54 *</b>	<b>−34.56</b>	<b>−385.34</b>	<b>−63.93</b>	5.21	13.84	32.15
Shigar at Shigar	34.3	74.56	13.89	96.5	201.25	1104.17	<b>−2175.64</b>	58.3	712.78	198.34	625.42	87.68
Indus at Kachura	<b>−42.86</b>	<b>−52.18</b>	71.56	<b>−156.76 *</b>	−179.95	39.87	<b>−699.07 *</b>	<b>−667.87</b>	<b>−462.05</b>	<b>−103.08</b>	<b>−103.85</b>	<b>−75.11</b>
Indus at Partab Bridge/Bunji	25.61 *	42.75 *	28.52 *	37.15 *	−108.62	154.05	137.87	<b>−201.63</b>	<b>−230.23</b>	14.17	2.55	20.53 *
Astore at Doyian	21.09 *	32.34 *	8.69	45.49 *	60.98 *	13.55	<b>−41.11</b>	<b>−13.91</b>	54.9 *	19.53	22.15 *	17.5
Indus at Shatial Br.	50 *	38.61 *	95 *	146.1 *	117.97 *	−141.98	174.12	411.86	<b>−11.62 *</b>	<b>−0.96</b>	23.22	24.88
Indus at Besham Qila	9.01 *	4.56	6.07	13.04	182.16 *	142.51	<b>−34.44</b>	<b>−222.47</b>	26.9	<b>−18.22</b>	0.32	1.28
Lower UIB at Daggar	3.58	<b>−8.16</b>	<b>−64.75 *</b>	<b>−54.38 *</b>	−16.14 *	<b>−21.98</b>	<b>−544.84 *</b>	<b>−219.64 *</b>	<b>−67.85</b>	0.87	1	12.6
Lower UIB at Karora	5.79	7.54	<b>−34.13 *</b>	<b>−31.57</b>	−5.66	2.25	<b>−57.05</b>	11.96	6.71	<b>−4.07</b>	3.45	16.6

Note: \* significant level 90%, negative values are in bold.

The Kharmong (Central Himalayas) station shows a significant decrease of flows during the months of December, January, May, and June. In the Lower UIB, the rainfall-dominated Brandu basin at Daggar and Gorbant basin at Karora (both in the Lower UIB in the West) mostly show significantly decreasing trends of discharges in the months between December and June. During the months of July, August, and September, an insignificant weak increase or decrease in flows is observed in most of the snow- and glacier-dominated basin. The Hunza at Dainyor (Western Karakorum) and Indus at Kachura (main UIB), however, show significantly decreasing trends of flows in the months of August and September. The lower UIB basins exhibit decreasing trends of flow during the months from July to September. The Gorbant basin at Karora (Lower UIB in the West) shows a significant decrease of flows in the months from July to September. In the months of October and November, flows are generally increasing in all the basins except for the Shigar basin.

Table 7, Figure 2 and Figure 4a show the percentage of seasonal and annual changes in mean values of flows on a temporal and spatial basis in the UIB. The winter discharges of the Indus River at Besham Qila (main UIB), Shatial Bridge (main UIB), Partab Bridge (main UIB), and Kachura (main UIB) and of the Astore River at Doyian (Western Himalayas) increase significantly by 3.3–8.5% per decade. The spring discharges of the main (Indus) River at Shatial Bridge (main UIB) and Partab Bridge (main UIB) and of Astore River at Doyian (Western Himalayas) are also significantly increased by 3.3–8.5% per decade. In the Lower UIB, however, the spring discharges of Gorbant at Karora (Lower UIB in the West) are reduced significantly by up to 20% of the mean values contrary to the basins containing water from glacier and snow melting. During the summers, there are no significant changes in flows of snow- and glacier-dominated basins except for the Shigar basin. In the Brandu at Daggar and Gorbant at Karora (both Lower UIB in the West) the summer flows decrease significantly by up to 23% per decade in both basins. During the autumns, the flows at Besham Qila (main UIB) increase significantly by up to 7.27% per decade, whereas the autumn discharges are reduced significantly in Brandu River at Daggar.

**Table 7.** Annual and seasonal percentage changes per decade for mean of data periods.

Sr #	Discharge Gauging Stations	DJF (%)	MAM (%)	JJA (%)	SON (%)	Annual (%)
<b>Per Decade % of Change in Discharge for Average Data Periods</b>						
1	Hunza at Dainyor	0.41	<b>−2.74</b>	<b>−4.25</b>	<b>−1.26</b>	<b>−7.64</b>
2	Gilgit at Gilgit	3.27	14.16	4.18	5.66	5.85
3	Gilgit at Alam Bridge	<b>−0.46</b>	3.33*	4.13	7.93	4.29
4	Indus at Kharmong	<b>−1.95</b>	2.41	<b>−4.42</b>	<b>−1.55</b>	<b>−5.37</b>
5	Shyok at Yogu	4.36	<b>−1.88</b>	9.1	8.67	3.35
6	Shigar at Shigar	8.74	25.52	47.88 *	16.11	35.38 *
7	Indus at Kachura	3.32*	4.3	5.28	<b>−10.37 *</b>	0.6
8	Indus at Partab Bridge/Bunji	3.51*	9.6 *	3.18	0.44	1.14
9	Astore at Doyian	8.22 *	18.59 *	0.96	2.1	8.84
10	Indus at Shatial Br.	8.49 *	13.58 *	3.96	<b>−0.42</b>	<b>−1.49</b>
11	Indus at Besham Qila	6.95 *	5.82	0.42	7.27 *	<b>−0.71</b>
12	Lower UIB at Daggar	6.62	<b>−19.8</b>	<b>−23.61 *</b>	<b>−0.65</b>	<b>−18.30 *</b>
13	Lower UIB at Karora	<b>−4.84</b>	<b>−19.82 *</b>	<b>−22.92 *</b>	<b>−17.89 *</b>	<b>−20.34 *</b>

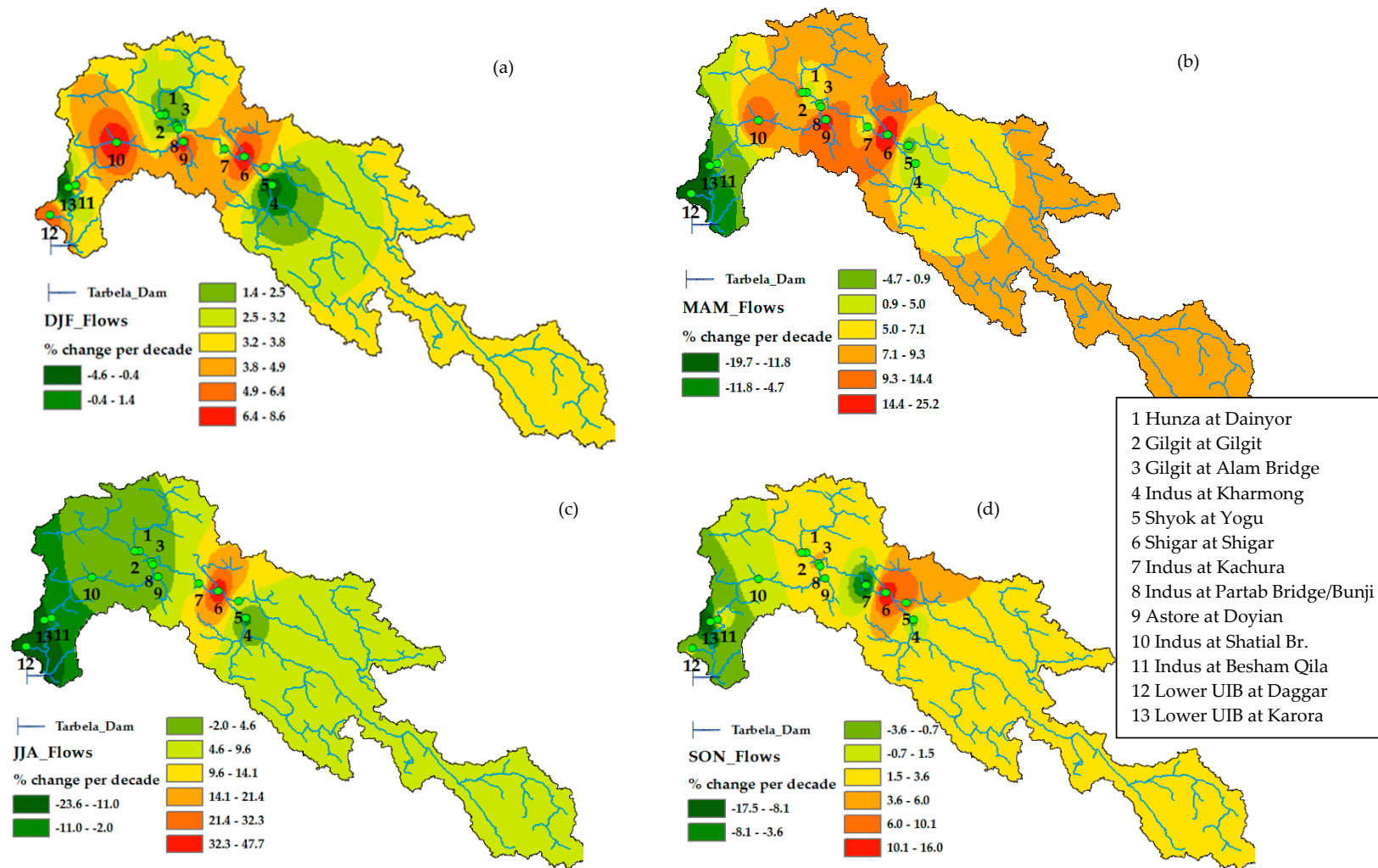
Table 7. Cont.

Sr #	Discharge Gauging Stations	DJF (%)	MAM (%)	JJA (%)	SON (%)	Annual (%)
<b>Per Decade % of Change in Suspended Sediments for Average Data Periods</b>						
1	Hunza at Dainyor	20.08 *	<b>-0.84</b>	<b>-18.53</b>	<b>-7.6</b>	<b>-17.66</b>
2	Gilgit at Gilgit	20.27 *	3.98	<b>-27.79 *</b>	<b>-2.52</b>	<b>-10.28</b>
3	Gilgit at Alam Bridge	29.03 *	2.57	<b>-18.63 *</b>	<b>-21.37 *</b>	<b>-18.56 *</b>
4	Indus at Kharmong	3.39	16.49	10.91	<b>-11.73</b>	2.46
5	Shyok at Yogu	2.52	8.85	<b>-6.59</b>	<b>-3.68</b>	<b>-14.29</b>
6	Shigar at Shigar	1.89	6.55	<b>-17.3</b>	42.87	<b>-2.2</b>
7	Indus at Kachura	<b>-12.77</b>	<b>-5.96</b>	<b>-23.84 *</b>	<b>-27.87 *</b>	<b>-24.06 *</b>
8	Indus at Partab Bridge/Bunji	28.76 *	<b>-2.01</b>	<b>-3.97</b>	<b>-6.57</b>	<b>-9.43</b>
9	Astore at Doyian	31.86 *	14.6 *	4.93	18.19	11.87
10	Indus at Shatial Br.	40.72 *	33.11 *	7.68	<b>-2.84</b>	5.75
11	Indus at Besham Qila	2.68	19.9 *	-5.75	1.8	<b>-4.51</b>
12	Lower UIB at Daggar	0.84	<b>-16.88 *</b>	<b>-24.85 *</b>	<b>-15.56 *</b>	<b>-28.2 *</b>
13	Lower UIB at Karora	5.05	<b>-5.73</b>	<b>-1.38</b>	<b>-2.44</b>	<b>-0.74</b>

Note: \* significant level 90%, negative values are in bold.

On an annual basis, flows of glacier- and snow-dominated basins do not change significantly, except for the Shigar basin. However, the annual discharges of the rainfall-dominated Brandu at Daggar and Gorband at Karora are reduced significantly by up to 20% per decade.



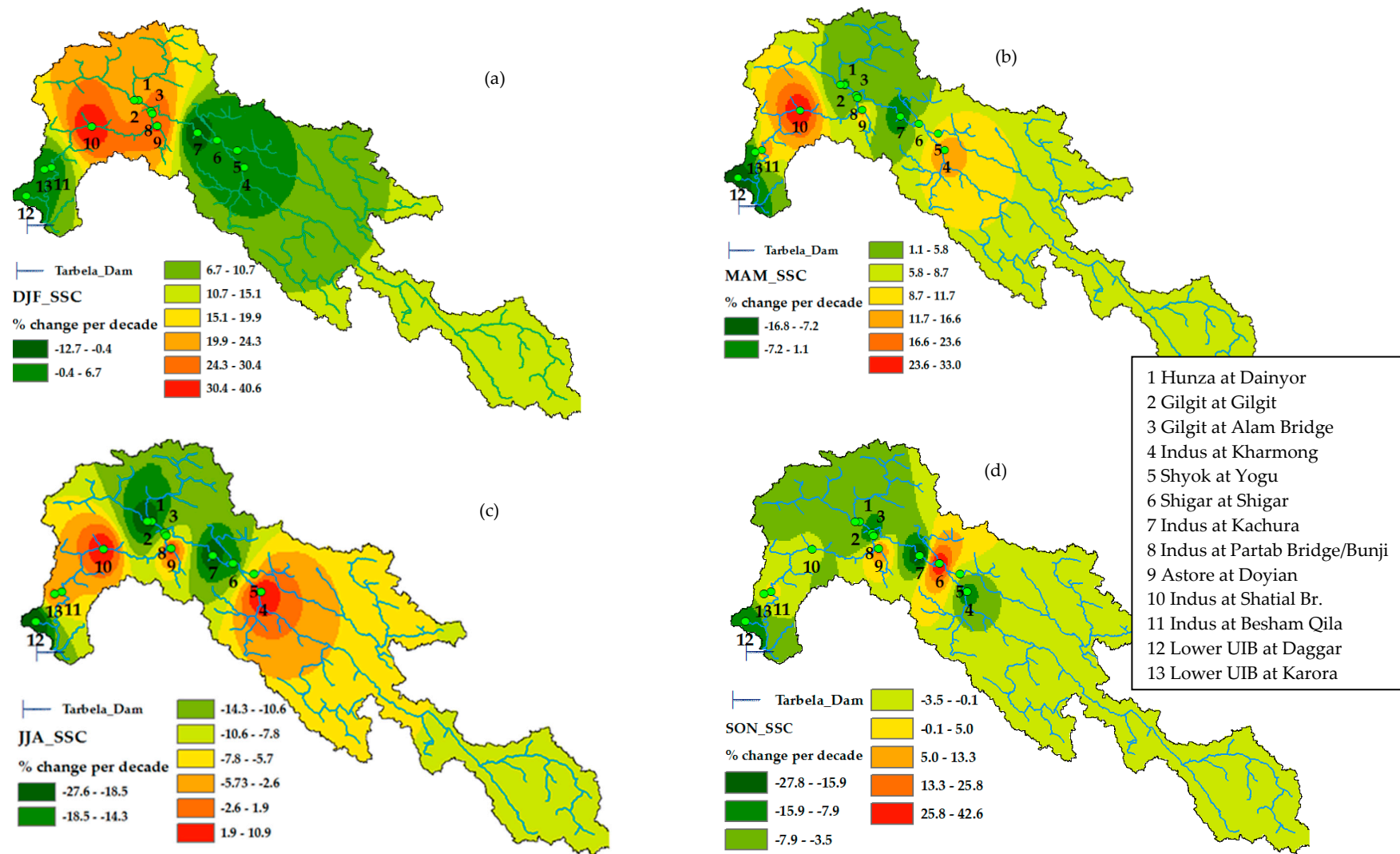


**Figure 2.** Seasonal flow trends spatially distributed over the basin with % of changes per decade in % of mean values for the whole period. (a) DJF (winter), (b) MAM (spring), (c) JJA (summer), (d) SON (autumn) seasons.

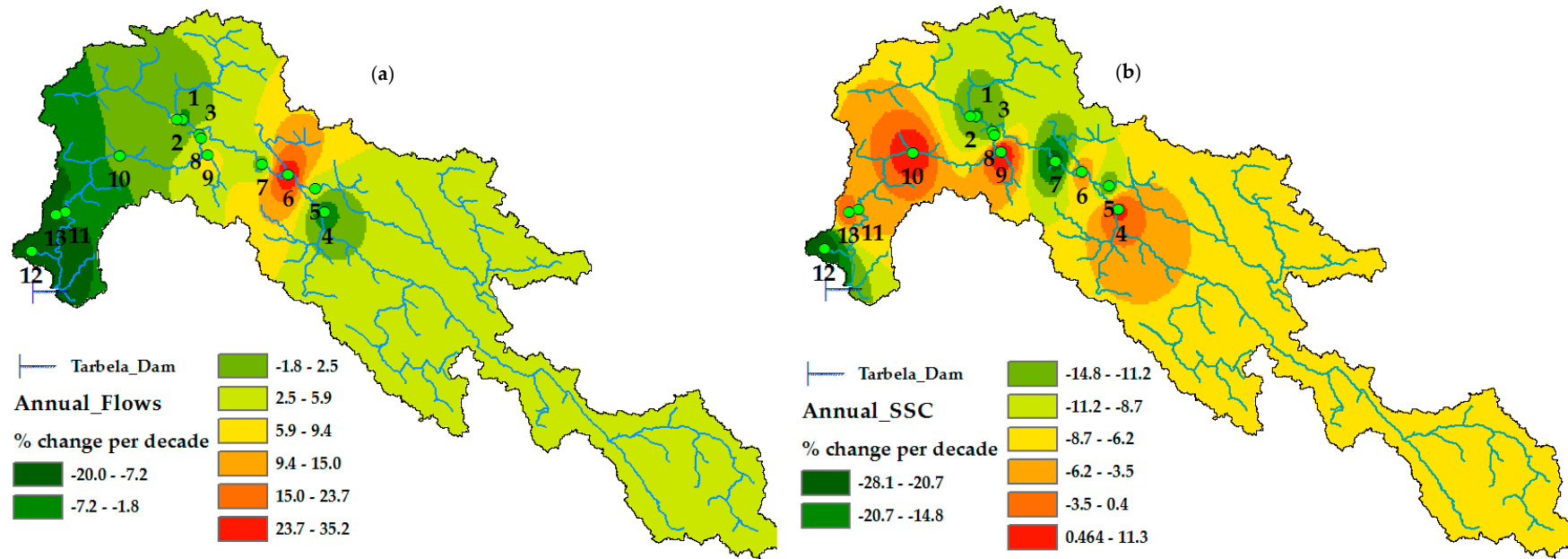
#### 4.3. Monthly, Seasonal, and Annual Trends of Suspended Sediment Concentrations (SSC)

Table 6 shows the monthly trends of suspended sediment concentrations (SSC) (mg/L) with 90% significance level, which have been obtained by using the Sens's slope estimation method. In the months from December to May the trends of SSC are mixed in upper sub-basins of the UIB. However, most of the stations located on snow- and glacier-dominated basins show an increase of the SSC from December to May along with, though disproportionate, positive discharge trends. The Indus at Partab Bridge (main UIB) and Shatial Bridge (main UIB), Astore River at Doyian (Western Himalayas), and Gilgit at Alam Bridge (Gilgit + Hunza) show significantly increasing trends of SSC (mg/L) and discharges per decade in the months between December and June. Similarly, Gilgit River at Gilgit (Hindukush) exhibits a major increase of SSC in the months of December, February, and March. The Indus at Besham Qila shows a significant increase of SSC per decade in the months of December and May. In contrast to this, the Indus at Kharhong (Central Karakorum) and Kachura (main UIB) shows an insignificant reduction of SSC in the months of December, January, and February. Contrary to the upper sub-basins of the UIB, where the SSC is increased in the months from December to May, the Lower UIB exhibits decreased SSC and discharges in the same period. The Lower UIB at Daggar and Karora is found to have a significantly reduced SSC in the months between March and May. During the months from June to October and especially from July to September, most of the glacier- and snow-fed basins show a decrease in SSC. The Gilgit at Gilgit (Hindukush) and Alam Bridge (Gilgit + Hunza), Shoyk River at Yogo (Western Karakorum), and the Indus at Shatial Bridge (main UIB), Kachura (main UIB), and Kharhong (Central Karakorum) show a significant reduction of SSC in the months between June and September with no significant trend in discharges in the same period. In the month of November, the upper glacier- and snow melt-dominated basins of Gilgit at Gilgit (Hindukush), Gilgit at Alam Bridge (Gilgit + Hunza), and Astore at Doyian (Western Himalayas) show a significant increase of SSC. The SSC in the Lower UIB, i.e., in Brandu River at Daggar (Lower UIB in the West) is found to be reduced significantly per decade with an insignificant reduction of discharges during the months of July and August.

The seasonal and annual percentage changes per decade in mean values of SSC on a temporal and spatial basis in the UIB are presented in Table 7, Figures 3 and 4b. Table 7 and Figure 3 show that during the winter season, SSC values increase significantly by 20%–40% for the basins of Gilgit at Gilgit (Hindukush), Hunza at Dainyor (Western Karakorum), Gilgit at Alam Bridge (Gilgit+Hunza), Astore at Doyian (Western Himalayas), and the Indus at Shatial Bridge (main UIB) and Partab Bridge (main UIB) in the UIB. During the spring season, the basins of Astore at Doyian (Western Himalayas) and Indus at Shatial Bridge (main UIB) and Partab Bridge (main UIB) also show a significant increase in SSC in the range of 14%–33% per decade in mean values. In the rainfall-dominated lower basin of Brandu at Daggar (Lower UIB in the West), however, the SSC decreases by up to 16.8% of its mean value per decade. During the summer season, the SSC levels in the UIB decrease significantly at most of the stations, such as Gilgit at Gilgit, Hunza at Dainyor, Gilgit at Alam Bridge, Shigar at Shigar, Shyok at Yogo, Indus at Kachura, and Brandu at Daggar. In this period, the SSC values of Gilgit at Gilgit (Hindukush), Gilgit at Alam Bridge (Gilgit + Hunza), Kachura (main UIB), and Brandu at Daggar are reduced by 18 to 28% of their mean values per decade. In the autumns, the SSC trends for Gilgit at Alam Bridge (Gilgit + Hunza), Kachura (main UIB), and Brandu at Daggar are also significantly reduced by 15%–28%. As obvious from Figure 4b, the annual SSC trends at most of the stations in the UIB decrease except for Astore at Doyian (Western Himalayas), Kharhong (Central Himalayas), and Indus at Shatial Bridge (main UIB). Figure 4b and Table 7 show that the annual SSC trends of Gilgit at Alam Bridge (Gilgit + Hunza), Indus at Kachura, and Brandu at Daggar) decrease significantly by 18%–28% per decade with no significant reduction of discharges except for Brandu at Daggar.



**Figure 3.** Seasonal suspended sediment concentration (SSC) trends spatially distributed over the basin with % of changes per decade in % of mean values for the whole period. (a) DJF (winter), (b) MAM (spring), (c) JJA (summer), (d) SON (autumn) seasons.



**Figure 4.** Annual flow and suspended sediment concentration (SSC) trends spatially distributed over the basin with % of changes per decade in % of mean values for the whole period. (a) Annual flows, (b) annual SSC.

#### 4.4. Mean Basin Precipitation Trends

Total changes in the basin's monthly precipitation (mm) for three decades (1981–2010) are shown in Table 8. In the winter months from December to January, there is an increase of precipitation in most of the basins probably due to increase of westerlies rainfall. In the month of February, Gilgit at Alam Bridge (Gilgit + Hunza) shows a significant increase in the basin-averaged precipitation, which is probably due to the effect of the Western front of the Hindukush during winters. During the spring months of March and April, precipitation is lowest, except in the Gilgit (Hindukush) basin. The month of March is the driest period in all basins.

In the springs, the basins of Gilgit at Gilgit (Hindukush) and Gilgit at Alam Bridge (Gilgit + Hunza) show a very wet period in the month of April. However, the Shyok at Yogo (Easter Karakorum), Astore at Doyian (Northwestern Himalayas), Brandu at Daggar (Lower UIB in the West), and Gorbant at Karora (Lower UIB in the West) exhibit the driest trends in March. Drying in March can be found at almost all stations, except for Shigar. During the summer seasons, most of the stations show an increase in rainfall. The stations of Gilgit at Gilgit (Hindukush) and Gilgit at Alam Bridge (Gilgit + Hunza) record a significant increase in rainfall in the months of June and July. During the summer months, however, the rainfall-dominated basins of Brandu at Daggar (Lower UIB in the West) and Gorbant at Karora (Lower UIB in the West) show reduced rainfall, except for the month of June. In the autumns, from September to November, there is an increase of precipitation in the UIB. In the month of October, however, precipitation generally is reduced in the UIB.

**Table 8.** Total changes in monthly precipitation (mm) extracted from the grid data for the Upper Indus Basin (UIB) from 1981 to 2010.

Ser #	Basin	Region	Elevation (m)	January	February	March	April	May	June	July	August	September	October	November	December
1	Hunza	Western Karakorum	3996	0.0	0.0	0.0	0.3	0.0	0.9	-0.1	0.5	0.2	0.0	0.0	0.0
2	Gilgit	Hindukush	4053	27.6	38.8	-5.4	<b>122.6</b>	12.1	<b>44.6</b>	<b>29.2</b>	17.2	11.9	-0.8	11.8	-4.4
3	Alam Br.	Hunza + Gilgit	4550	39.4	<b>49.5</b>	-21.4	<b>96.7</b>	6.3	<b>41.4</b>	17.8	14.0	8.0	-3.0	6.9	-5.5
4	Kharmong	Central Himalayas	4798	8.7	-28.8	-95.4	-23.1	0.1	14.3	-21.4	5.9	3.9	-0.8	6.8	-1.1
5	Shyok	Eastern Karakorum	5094	4.7	-16.3	<b>-103.0</b>	5.6	-3.4	1.0	-20.8	17.6	13.3	-2.3	2.3	-11.6
6	Shigar	Central Karakorum	4497	<b>241.3</b>	<b>256.7</b>	104.5	88.3	84.4	44.1	41.0	-26.5	<b>49.2</b>	-13.1	-1.3	<b>-38.9</b>
7	Kachura	Main River UIB	4922	27.8	47.8	-86.2	32.0	-1.9	9.9	-1.6	16.8	9.3	-2.0	5.0	-1.8
8	Bunji	Main River UIB	4541	24.7	61.0	-81.2	50.2	-2.1	9.4	1.8	14.6	8.1	-1.0	4.9	0.3
9	Astore	North Western Himalayas	3996	48.9	57.1	<b>-174.3</b>	45.6	-10.7	11.5	1.2	-10.3	-1.7	-13.8	0.2	-11.9
10	Shatial	Main River UIB	4488	26.3	56.0	-83.1	46.8	-2.7	8.7	0.9	15.2	8.5	-1.7	5.1	-0.9
11	Besham Qila	Main River UIB	4312	26.2	56.3	-83.1	47.0	-2.8	8.8	0.9	14.9	8.5	-1.7	5.1	-0.5
12	Daggar	Lower Western UIB	1111	5.9	25.2	<b>-178.1</b>	-48.6	-19.5	<b>35.4</b>	-7.1	-50.5	12.1	-23.4	8.5	-19.6
13	Gorband	Lower Western UIB	2257	21.9	6.9	<b>-217.1</b>	-45.1	-22.4	<b>43.9</b>	-20.1	-42.3	11.5	-25.3	12.8	-25.1
14	Tarbela	UIB	4421	25.7	29.7	<b>-107.9</b>	27.2	-19.7	14.8	-8.5	2.2	2.2	-6.1	2.7	-9.2

Note: Values in a box and written in bold represent the highest values of rainfall in mm with 90% significant level. Red color reflects the months of reduced precipitation (driest), blue color indicates the months with highest increased of precipitations (wettest).

#### 4.5. Mean Monthly Average Temperature Trends

Table 9 shows the change in mean monthly average temperature ( $^{\circ}\text{C}$ ) between 1981 and 2010. The temperature trends of 18 meteorological stations located at different altitudes in UIB are shown with ascending altitude. Generally, most of the stations at higher and lower altitudes of the UIB show warming trends during the months of February and March. The Gilgit station at 1460 m altitude shows a significant increase in average temperature in the range of 1.23–3.02  $^{\circ}\text{C}$  during the months of January, February, and March. During the months of February, March, and April, the average temperature increases by 0.58–4  $^{\circ}\text{C}$  at altitudes from 1250 m to 4440 m. In the late springs and early summers, i.e., in the months of May and June, there is an insignificant increase of average temperature at most of the higher and lower altitudes of the UIB. During summers and early autumns, i.e., in the months of July, August, and September, however, there is generally a cooling trend ranging from 0.03–1.87  $^{\circ}\text{C}$ . The stations of Gilgit in Giglit (Hindukush), Gupis in Gilgit (Hindukush), Chillas (Main UIB), and Skradu (Main UIB) show a significant decrease of average temperature by 1.34–1.87  $^{\circ}\text{C}$  during summers and early autumns. In the late autumns, average temperature increases again. In the months of October and November, temperatures at higher altitudes increase at higher rates (0.97–2.38  $^{\circ}\text{C}$ ) than valley stations (Table 9).

**Table 9.** Changes in mean monthly averaged temperatures (°C) from 1981 to 2010 for stations located at different altitudes in the Upper Indus Basin (UIB).

Ser #	Station	Sub-Basin	Elevation (m)	January	February	March	April	May	June	July	August	September	October	November	December
1	Chilas	Besham Qila	1250	0.54	0.72	<b>2.35</b>	−0.13	0.24	−28	−1.12	<b>−1.68</b>	−1.16	0.77	0.24	−0.06
2	Bunji	Partab Br.	1372	1.08	1.46	<b>3.23</b>	<b>0.52</b>	0.98	0.31	−1.05	−1.21	−0.13	0.99	<b>1.12</b>	0.58
3	Gilgit	Gilgit	1460	<b>1.23</b>	<b>1.89</b>	<b>3.02</b>	0.23	0.92	0.64	−1.11	<b>−1.34</b>	−0.50	0.61	<b>0.97</b>	0.76
4	Shangla	Gorband	1960	0.99	1.52	<b>3.24</b>	<b>1.63</b>	<b>1.51</b>	0.21	0.31	0.18	−0.08	1.08	0.82	0.68
5	Gupis	Gilgit	2156	−0.26	−0.14	1.79	−0.35	−0.28	−0.01	<b>−1.87</b>	−1.29	−0.53	<b>2.44</b>	0.98	−0.47
6	Astore	Astore	2168	0.98	<b>2.43</b>	<b>3.13</b>	1.30	0.94	<b>1.48</b>	0.00	−0.51	0.27	1.01	<b>1.67</b>	0.61
7	Skardu	Shigar	2210	−0.69	−1.03	−0.26	0.36	0.83	0.92	0.08	−0.68	−0.43	0.69	0.58	−0.05
8	Skardu	Main Kachura	2210	1.11	1.63	1.87	0.29	−0.21	−0.01	−0.80	−0.92	<b>−1.60</b>	−0.35	0.83	−0.23
9	Raitu	Astore	2718	1.17	<b>2.61</b>	<b>3.19</b>	<b>2.07</b>	1.37	0.91	0.55	0.23	−0.07	<b>1.94</b>	<b>1.03</b>	0.23
10	Naltar	Hunza	2898	1.16	<b>1.74</b>	<b>2.86</b>	1.29	0.66	1.26	0.16	0.32	0.36	<b>1.86</b>	<b>1.63</b>	0.60
11	Ziarat	Hunza	3020	1.18	<b>1.99</b>	<b>3.05</b>	<b>1.46</b>	0.69	1.21	0.37	0.46	0.49	<b>1.85</b>	<b>2.16</b>	1.25
12	Ushkore	Gilgit	3051	1.07	<b>2.20</b>	<b>2.68</b>	1.43	0.66	0.95	−0.27	−0.06	−0.16	<b>1.95</b>	1.19	0.34
13	Hushey	Shyok	3075	0.81	<b>1.77</b>	<b>3.54</b>	1.47	1.12	0.82	0.70	0.56	0.74	<b>1.67</b>	<b>1.80</b>	1.27
14	Rama	Astore	3179	1.02	<b>1.88</b>	<b>3.17</b>	<b>1.58</b>	0.93	0.60	0.46	0.30	0.18	<b>1.60</b>	0.99	0.50
15	Yasin	Gilgit	3280	1.13	<b>1.56</b>	<b>2.81</b>	1.20	0.69	<b>1.47</b>	−0.23	0.04	−0.20	1.47	−0.02	<b>−1.44</b>
16	Shendure	Gilgit	3712	1.20	<b>1.98</b>	<b>2.17</b>	1.23	0.80	0.95	−0.03	0.03	−0.21	<b>2.07</b>	0.59	0.20
17	Desosai	Kharmong	4149	1.00	<b>1.53</b>	<b>3.15</b>	<b>1.56</b>	1.23	1.35	1.06	0.24	0.44	<b>1.86</b>	1.20	0.71
18	Khunjrab	Hunza	4440	1.38	<b>1.87</b>	<b>4.00</b>	<b>1.79</b>	1.56	<b>1.37</b>	0.22	0.44	0.66	<b>2.01</b>	<b>2.38</b>	1.01

Note: Values in a box and written in bold represent the highest values of average temperature (°C) at station altitudes with 90% significant level. Red color reflects an increase in monthly temperature (warming), blue color indicates decrease in monthly temperature (cooling).



#### 4.6. Pearson's Correlations between Flows, SSC and Gridded Climatic Variables

To evaluate the dominance of different hydro-climatic variable in flows and sediment generation process the correlation between these variables was found out. The Table 10 shows the results of Pearson's correlations coefficients between the daily flows, SSC and gridded basin averaged rainfall and basin averaged mean air temperature. The Table 10 describes that correlation coefficient between flows and SSC is more significant on main Indus River for example; Indus at Shatial Bridge and Indus at Besham Qila. However, in the glacier and snow dominated basins, the correlation between flows and SSC is less significant in most of the sub basins compared to the main Indus River. Similarly, in lower UIB sub basins, the correlation between flows and SSC is least significant.

**Table 10.** Pearson's correlations between daily flows, SSC and gridded climatic datasets since 1981–2010 in the Upper Indus Basin (UIB).

Ser #	Discharge Gauging Stations	Region	Pearson Correlations		
			Q vs. SSC	Q vs. R	Q vs. Tavg
1	Hunza at Dainyor	Western Karakorum	0.51	−0.01	0.77
2	Gilgit at Gilgit	Hindukush	0.73	0.04	0.82
3	Gilgit at Alam Bridge	Hunza + Gilgit	0.64	−0.06	0.80
4	Indus at Kharhong	Central Himalayas	0.23	0.14	0.79
5	Shyok at Yogu	Eastern Karakorum	0.46	0.06	0.74
6	Shigar at Shigar	Central Karakorum	0.49	−0.02	0.72
7	Indus at Kachura	Main River UIB	0.68	−0.14	0.79
8	Astore at Doyian	North Western Himalayas	0.27	0.04	0.71
9	Indus at Partab Bridge/Bunji	Main River UIB	0.62	0.03	0.77
10	Indus at Shatial Br.	Main River UIB	0.75	0.03	0.80
11	Indus at Besham Qila	Main River UIB	0.73	0.01	0.80
12	Lower UIB at Daggar	Lower western UIB	0.47	0.37	0.05
13	Lower UIB at Karora	Lower western UIB	0.11	0.13	0.11

The correlation coefficient between flows/SSC and gridded basin averaged rainfall is non-significant on main Indus River as well as at glacierized snow and ice melted dominated sub basins. However, in lower UIB at Daggar the gridded basin averaged rainfall is better correlated than temperature. The gauging stations lying on the main Indus River as well in snow and ice melted sub basins showed significant correlations between flows/SSC and gridded basin averaged mean air temperature. Ali K.F et al. [37] found that the percent of snow/ice cover (LCs) is major land cover controlling parameter along with temperature and seasonal rainfall. Ali K.F et al. [37] also found that the combination of snow/ice cover (LCs) and climatic variables explain 98.5% of variance in sediment yield in UIB. Similarly, for the lower monsoon dominated UIB, the mean annual rainfall explains the 99.4% of variance in sediment yields.

## 5. Discussion

Except for Kharhong and Astore, the annual trends of the SSC in the glacier- and snow-dominated basins of the UIB decrease in the months of July, August, and September, which is due to less snow/glacial melt in the cooler summer, as was reported by researchers [7,8,38,39]. In contrast to the SSC trend, an insignificant increase or no change in annual flows is detected, as is shown in Tables 6 and 7. Despite the cooler summer that leads to fewer discharges in glacierized upper basins like Hunza, Kharhong, and Astore, the absence of a significantly negative trend at Besham Qila indicates that this decrease is compensated by increased discharges from other sub-basins of the UIB.

During the winter season, Gilgit (Hindukush) and Hunza (Western Karakorum) show a significant increase of SSC along with discharges. This might be associated partly with a significant increase in the average winter temperature and an insignificant increase of precipitation due to westerlies in winter.

Increasing temperature increases the snow melting rates and reduces seasonal precipitation in the form of snow during winters, which may lead to increased erosions in the catchment. A similar trend of an increase in SSC during the winter is observed for the basins of Astore (Northwestern Himalayas), Gilgit at Alam Bridge (Hunza + Gilgit), Indus at Partab Bridge (Main UIB), and Indus at Shatial Bridge (Main UIB), along with a significant increase of average temperature in the UIB.

During the spring season, the Astore at Doyian (Northwestern Himalayas), Indus at Shatial Bridge (Main UIB), Indus at Besham Qila (UIB), as well as Gilgit at Alam Bridge show a significant increase of SSC with a major increase in flows. The Indus at Khariong and Shigar at Yogo show an insignificant increase of SSC and discharges. The increased flows in the river might result in an increase in effective discharges and erosion of the deposited sediments. Moreover, the sources of sediments in the glacial and snow melt-dominated catchments might be activated by rapid melting of snow accumulated in winter because of increasing temperatures in February and March. In the lower rainfall-dominated Daggar basin (Lower UIB), however, the SSC in springs decreases significantly with decreasing flows, which could probably be due to a reduction of spring rainfall. Table 8 shows that the rainfall at Daggar, Gorbant, and Besham Qila in the Lower UIB decreases significantly in March, April, and May.

During the summer season, the SSC decreases in glacier-dominated basins, such as Gilgit (Hindukush), Hunza (Western Karakorum), Shigar (Central Karakorum), and Shyok at Yogo. This reduction of SSC in the summer at Gilgit (Hindukush) could be due to significant cooling of the summer temperature (Indus climate anomaly). Cooling of the summer temperature can be observed up to 3000 m altitude in the Gilgit basin. As in the UIB, many glaciers are located at elevations from 2500 to 7000 m. Seasonal snow accumulates in the elevation zone from 3000 m to 7000 m. Furguson [50] describes that in UIB the mass movement, supraglacial and sub glacial sediments transport supply large amount sediments in valley. These sediments are stored in alluvial fans, outwash trains and moraines. The sediment is mainly transported by snow and ice melt process. Collins [51] estimated that for Hunza basin and Indus at Bashem Qila drives 60% and 40% of sediment loads annually from glacier melting. Own et al. [52] concluded that debris transport is an important contributor in glacierized regions of the UIB. Glacier debris transport yields a large amount of coarse rock debris and sediments from the basal traction zone. The debris-covered glacier in the UIB lies below 3000 m [36,40]. This debris cover accelerates/deaccelerates melting [53,54]. This significant cooling of summer temperature up to 3000 m altitude probably reduces the supply of snow water from debris-covered glaciers and may increase the snow cover at lower and middle altitudes. This might result in a reduced SSC at the outlet. The Indus at Kachura also shows a decreasing trend of SSC over the entire year with a significant reduction in summers and autumns, which is probably also due to reduced temperatures and, hence, smaller snow water discharge rates in these months. In the Daggar (Lower UIB) basin, the SSC and flows are reduced significantly during summer, which is probably caused by a smaller rainfall intensity leading to smaller flows, sediment yields, and catchment erosion in the months of July and August.

In the autumn season, mixed trends of SSC are observed. Table 8 shows no significant change in precipitation during autumns. During the month of October, however, precipitation amounts generally are reduced in comparison to the precipitations of September and November. Moreover, as shown in Table 9, the month of September shows a lower average temperature at lower altitudes. So, the reduction of summer SSC from the UIB could probably be due to an increase in precipitation in the form of snow and a decrease of rainfall amount.

## 6. Conclusions

It may be concluded that the annual flows and SSC in the upper Indus Basin at Besham Qila are in a balanced state with a small reduction during the analysis period of three decades. However, the annual SSC values in the upper snow- and glacier-dominated Hindukush and Karakorum basins decrease, whereas the values of the mixed snow melt- and rainfall-dominated Western and Central Himalayan basins increase. The significant SSC reduction in the Hindukush and Western Karakorum basins during summers could be attributed to the Karakorum climate anomaly [7,8]. During winters, SSC values

increase along with flows due to increasing temperatures in the glacier- and snow-dominated basins. In the lower basins of Daggar and Gorband near Besham Qila, the SSC and flows decrease significantly during the spring, autumn, and summer seasons when rainfalls decrease. Contrary to the summer's SSC pattern, the reduced SSC at Kachura, coupled with increasing trends of SSC at Partab Bridge and Shatial Bridge on the main River during winter and spring, confirming the findings of [55] that sediments are deposited in the summer and erosion takes place in winter and spring between Kachura and Besham Qila. Moreover, winter erosion might cause increasing trends of discharges in this area, which may enhance the capacity of the river to erode and transport the suspended sediment. Summer sediment deposition in this area in the absence of any significant discharge trends might be due to natural morpho-dynamic processes to compensate winter erosion and maintain the balanced flow regime. These findings improve understanding of the erosion process in different sub-basins of the UIB and, sediment erosion/deposition and transport in the Indus River. Now, sediment budgets can be determined in the light of climate change. Moreover, the analysis of altering patterns of precipitation, temperature, and flow and their impacts on the SSC values of different sub-basins of the UIB will be important for the management of existing (by developing appropriate reservoir operation scenarios), and the design of future hydraulic structures on the Indus and its tributaries in the UIB. This study will also help investigate the changing patterns of water quality, ecosystems, and geomorphology of the rivers of the UIB.

**Author Contributions:** W.U.H. designed the work, highlighted the problem, formulated the work plan, and analyzed the data. M.K.S. assisted in work plan formulation and helped in the interpretation of results, drafted the general outlines of the study in a coherent paper, and reviewed it. A.C. coordinated the grid data and extracted them. F.S. and F.N. assisted by providing practical knowledge and correcting the methodology. All authors have read and agreed to the published version of the manuscript.

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