

Modeling pH-Dependent Adsorption of Glyphosate on Iron Hydroxides: Competition with Phosphate and Influence of Fe²⁺

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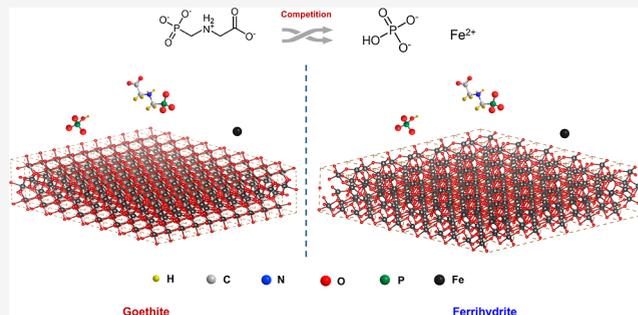
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ABSTRACT: Understanding and predicting the glyphosate (Gly) migration at the water–soil interface remains a critical challenge in environmental pollution control. Iron hydroxides, particularly goethite and ferrihydrite, serve as important sinks for pollutants and are significant for geochemical modeling. This study systematically investigates the pH-dependent competitive adsorption of Gly and phosphate (PO₄) on these iron minerals, alongside the influence of ferrous ions (Fe²⁺) on Gly adsorption—a process whose mechanistic background remains inadequately explored in previous studies. Utilizing the Charge Distribution Multisite Surface Complexation (CD-MUSIC) model, we analyzed adsorption data to evaluate surface complexation constants (log K), binding mechanisms, and competitive effects. Results revealed that Gly and PO₄ adsorption is primarily governed by pH, concentration, mineral type, and competitive interactions, with PO₄ significantly suppressing Gly retention. Nevertheless, Gly maintains measurable adsorption even under strong competition, attributed to irreversible bidentate complexation. The model's underestimation at higher loadings further points to the likely role of steric hindrance effects, a factor not accounted for in the current CD-MUSIC formulation. Fe²⁺ exhibits dual roles: it enhances Gly adsorption on ferrihydrite via coadsorption, potentially via electrostatic attraction and possible on distinct sorption sites, while its impact on goethite is negligible. These findings underscore the PO₄ critical role in reducing Gly retention in iron-rich soils and highlight the context-dependent influence of Fe²⁺. Our work advances the application of CD-MUSIC modeling to complex organic–inorganic systems and provides insights for optimizing soil remediation to mitigate Gly leaching.

KEYWORDS: *glyphosate, phosphate, ferrous ions, goethite, ferrihydrite, CD-MUSIC*



1. INTRODUCTION

Glyphosate (*N*-(phosphonomethyl)glycine, Gly), the most widely used broad-spectrum herbicide globally, has an estimated average annual application rate in European countries of 0.62 kg per hectare from 2011 to 2017.¹ Its environmental persistence has been well-documented, with residues detected in soils, surface waters, and groundwater systems,^{2–5} raising significant concerns about its environmental impact. These issues necessitate a comprehensive understanding of Gly behavior in natural systems, particularly its interactions with soil minerals and competing ions.

As a key nutrient supporting global agricultural productivity, phosphate (PO₄) shares structural and chemical similarities with Gly. This leads to competitive adsorption onto soil minerals—a process critical for environmental mobility and bioavailability of both compounds.^{6,7} Iron hydroxides such as goethite and ferrihydrite, due to their high surface reactivity and widespread distribution in soils and sediments, serve as primary sorbents for these anions. Ferrous iron (Fe²⁺), produced through anaerobic microbial respiration or mineral dissolution, further complicates the dynamics by catalyzing

mineral recrystallization, altering surface charge through reductive dissolution,^{8,9} and forming soluble complexes with Gly,¹⁰ that may enhance its mobility. However, there is no consensus on Gly surface complexation structures on mineral surfaces and its binding modes under competitive adsorption, which is ultimately relevant for understanding and modeling its pH- and concentration-dependent environmental transport.

Previous studies have employed various characterization techniques and density functional theory (DFT) calculations^{11–14} to analyze Gly surface speciation on minerals. Based on attenuated total reflection Fourier-transformed infrared (ATR-FTIR) spectroscopy interpretation, Barja and Afonso¹² proposed bridging bidentate adsorption configurations on goethite surfaces for pH 3.5–9.2 and at low-to-medium surface

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coverage. By contrast, Yan and Jing¹³ demonstrated pH-dependent transitions from binuclear/mononuclear bidentate complexes under acidic conditions to mononuclear monodentate forms at higher pH using combined spectroscopic and DFT approaches. Another molecular dynamics simulation study¹⁴ suggested the presence of intramolecular hydrogen bonding between the Gly phosphonate and amino groups and proposed adsorption primarily via monodentate complexation on goethite. The lack of a solid foundation for interpreting these spectroscopic and computational results has led to divergent conclusions across studies.

Similar research has analyzed the adsorption of PO₄ and Fe²⁺ on goethite,^{15–20} Kubicki et al.^{15,16} discussed DFT-assisted modeling of PO₄ adsorption on goethite. By comparing simulations with experimental characterization, they concluded that PO₄ tends to form bidentate bridging surface complexes at low pH. For Fe²⁺ adsorption on goethite, besides known mono- and bidentate complexes, studies have also indicated electron transfer from Fe²⁺ to ferric iron (Fe³⁺)¹⁹ and the formation of tridentate complexes involving double edge-sharing.²⁰ Given its high stability and crystallinity, goethite has been the primary focus of most research, while studies on Gly complexation with other iron minerals, such as ferrihydrite, remain scarce.

Although these studies provide new insights into adsorption in single-solute systems, predictive modeling of binary adsorption remains challenging. Established surface complexation models (SCMs), such as the Charge Distribution Multisite Complexation (CD-MUSIC) model, quantify ligand binding affinities at the mineral/water interface. The CD-MUSIC model achieves this by dissecting electrostatic contributions through CD parameters, which are often derived from molecular orbital/density functional theory (MO/DFT) analysis.²¹ These frameworks make SCMs particularly adept at identifying changes in surface complexation structures with pH and loading. Their predictive power has been successfully evidenced by the adsorption of metal ions^{22–24} and oxyanions.^{25–27} In contrast, the application of such models to organic molecules featuring multiple functional groups remains far less explored. In a recent key advance, Geysels and Hiemstra²⁸ established a set of CD-MUSIC parameters that comprehensively describes the pH- and loading-dependent sorption behavior of Gly on goethite, providing a foundation for extending this modeling approach to complex binary organic–inorganic adsorption systems.

Consequently, our work employs the CD-MUSIC framework to investigate the pH-dependent adsorption behavior of Gly on iron hydroxide, with particular emphasis on its competitive interactions with PO₄ and the influence of Fe²⁺. By combining systematic adsorption experiments with CD values derived from prior studies, we optimize the corresponding affinity constants (log K) to interrogate surface complexation mechanisms in both PO₄-Gly and Fe²⁺-Gly systems. In doing so, this study extends beyond previous work by explicitly testing CD-MUSIC predictions under higher surface loadings, stronger organic–inorganic competition, and mineralogical variability between crystalline goethite and poorly crystalline ferrihydrite. Importantly, the incorporation of Fe²⁺ introduces an additional dimension that has rarely been addressed in SCM-based descriptions of Gly adsorption, despite its environmental relevance in redox-dynamic soils. Together, these advances enable a critical evaluation of the strengths and current limitations of equilibrium SCMs for describing

organic–inorganic interactions at mineral–water interfaces. The resulting insights improve our mechanistic understanding of Gly fate in iron-rich soil and advance predictive capabilities for environmental risk management and remediation strategies.

2. MATERIALS AND METHODS

2.1. Chemicals and Materials

The chemicals used included glyphosate (96%, Sigma-Aldrich), potassium dihydrogen phosphate (KH₂PO₄, ≥99.5%, Merck), and iron(II) chloride tetrahydrate (FeCl₂·4H₂O, 99.99%, Sigma-Aldrich), all of which were obtained as commercial reagents. HEPES buffer (≥99.5%, Sigma-Aldrich) was utilized to maintain a stable pH of 7. Potassium chloride (KCl, 99.0–100.5%, Sigma-Aldrich) was employed in some experiments as the background electrolyte for experiments conducted at other pH throughout the sorption studies. All chemicals used for synthesis and quantitative analysis were of analytical grade. Laboratory solutions were prepared with ultrapure water (0.057 μS/cm) (GenPure Pro UV, Thermo Fisher Scientific, Germany) and stored at a 4 °C fridge in the dark prior to use.

Goethite was synthesized following the method described by Schwertmann and Cornell,²⁹ resulting in a N₂-Brunauer–Emmett–Teller (N₂-BET) specific surface area (SSA) of 41.9 m²/g. Synthetic two-line ferrihydrite³⁰ was prepared and used in all adsorption and titration experiments as a fresh suspension. A separate aliquot of the same batch was freeze-dried solely for characterization, yielding a BET SSA of 275.2 m²/g. It should be noted that the N₂-BET-SSA_{Feh} generally underestimates the real SSA due to irreversible aggregation of the particles during freeze-drying and outgassing.³¹ Therefore, for all CD-MUSIC modeling in this study, the SSA_{Feh} was set to 611 m²/g,³² representing the standard value for fresh ferrihydrite suspensions. Detailed synthesis and characterization are provided in Supporting Information (Figure S1 for BET analysis, Figure S2 for X-ray diffraction patterns).

2.2. Batch Sorption Experiments

All suspensions contained a background electrolyte of 0.02 M. Single and binary adsorption of Gly (0.05–2 mM) and PO₄ (0.05–2 mM) onto goethite (2.5 g/L) and ferrihydrite (0.17 g/L) was investigated through batch experiments at equilibrium pH 6, 7, and 8. These adsorption experiments were conducted under normal laboratory atmospheric conditions, as no Fe²⁺ was involved. Furthermore, the influence of CO₂ on the adsorption system was considered negligible; its potential effects on mineral surface charge and system concentration are discussed in Figure S3. Stock solutions of Gly and PO₄, as well as iron mineral suspensions, were preadjusted to the target pH using HCl (0.1–1 M) and KOH (0.1–1 M) prior to mixing. The equivalent stock target solution was then added to the suspension to achieve the desired final concentrations. For binary competitive systems, Gly and PO₄ were always introduced simultaneously and at equal concentrations across the same range (0.05–2 mM) to simulate a cocontamination scenario with a fixed 1:1 molar ratio. Fe²⁺ (0.05–2 mM) sorption experiments, conducted in the absence of PO₄ to isolate the binary Fe²⁺-Gly interaction, were performed in an N₂-glovebox (O₂ < 0.1 ppm) (UNILab plus, MBRAUN, Germany) to prevent O₂-driven redox reactions at pH 7. Notably, all solutions and suspensions were purged with helium (~1 h) prior to transferring into the glovebox to eliminate the effects of dissolved O₂. Binary sorption involved pre-equilibration with Gly (0.24 mM) before the addition of Fe²⁺. All samples were incubated in Falcon tubes and shaken for 12 h on a horizontal shaker (HS 501, IKA, Germany). Subsequently, the pH of each suspension was remeasured using a pH meter equipped with an ORP electrode (Seven2Go, Mettler Toledo, Switzerland), and phase separation was achieved by centrifugation at 12,000 rcf for 10 min to ensure complete sedimentation of colloidal particles for analysis.

The aqueous concentrations of Gly and PO₄ were quantified using ion chromatography (883 Basic IC plus, Metrohm, Switzerland) coupled with a conductivity detector (method description in Supporting Information, Figure S4) and by an ultraviolet–visible

Table 1. Estimated Log K and CD Values for Gly, PO₄, and Fe²⁺ Sorption on the Coordinated Surface Sites of Goethite^{abcdef}

complexation equation	form	log K	Δz ₀	Δz ₁	Δz ₂
≡FeOH ^{-0.5} + H ⁺ ↔ ≡FeOH ₂ ^{+0.5}		9.25	1	0	0
≡Fe ₃ O ^{-0.5} + H ⁺ ↔ ≡Fe ₃ OH ^{+0.5}		9.25	1	0	0
≡FeOH ^{-0.5} + H ⁺ + Cl ⁻ ↔ ≡FeOH ₂ ^{+0.5} ...Cl ⁻		8.95	1	-1	0
≡FeOH ^{-0.5} + K ⁺ ↔ ≡FeOH ^{-0.5} ...K ⁺		-0.3	0	1	0
≡Fe ₃ O ^{-0.5} + H ⁺ + Cl ⁻ ↔ ≡Fe ₃ OH ^{+0.5} ...Cl ⁻		8.95	1	-1	0
≡Fe ₃ O ^{-0.5} + K ⁺ ↔ ≡Fe ₃ O ^{-0.5} ...K ⁺		-0.3	0	1	0
2≡FeOH ^{-0.5} + 2H ⁺ + Gly ⁻³ ↔ ≡Fe ₂ Gly ⁻² + 2H ₂ O	B	21.15	0.69 ²⁸	-0.95 ²⁸	-0.74 ²⁸
≡FeOH ^{-0.5} + 2H ⁺ + Gly ⁻³ ↔ ≡FeHGly ^{-1.5} + H ₂ O	MH	15.40	0.28 ²⁸	-0.54 ²⁸	-0.74 ²⁸
2≡FeOH ^{-0.5} + 3H ⁺ + Gly ⁻³ ↔ ≡Fe ₂ HGly ⁻¹ + 2H ₂ O	BH	29.75	0.75 ²⁸	-0.01 ²⁸	-0.74 ²⁸
≡FeOH ^{-0.5} + 3H ⁺ + Gly ⁻³ ↔ ≡FeH ₂ Gly ^{-0.5} + H ₂ O	MH ₂	20.79	0.41 ²⁸	0.33 ²⁸	-0.74 ²⁸
2≡FeOH ^{-0.5} + 2H ⁺ + PO ₄ ⁻³ ↔ ≡Fe ₂ PO ₄ ⁻² + 2H ₂ O	B	26.71	0.46 ⁴²	-1.46 ⁴²	0 ⁴²
≡FeOH ^{-0.5} + 2H ⁺ + PO ₄ ⁻³ ↔ ≡FeHPO ₄ ^{-1.5} + H ₂ O	MH	26.52	0.28 ⁴²	-1.28 ⁴²	0 ⁴²
2≡FeOH ^{-0.5} + 3H ⁺ + PO ₄ ⁻³ ↔ ≡Fe ₂ HPO ₄ ⁻¹ + 2H ₂ O	BH	33.45	0.63 ⁴²	-0.63 ⁴²	0 ⁴²
≡FeOH ^{-0.5} + 3H ⁺ + PO ₄ ⁻³ ↔ ≡FeH ₂ PO ₄ ^{-0.5} + H ₂ O	MH ₂	33.25	0.33 ³⁸	-0.33 ³⁸	0 ³⁸
≡FeOH ^{-0.5} + Fe ²⁺ ↔ ≡FeOHFe _{II} ^{+1.5}	M	5.83	1.00	1.00	0
2≡FeOH ^{-0.5} + Fe ²⁺ ↔ ≡(FeOH) ₂ Fe _{II} ⁺¹	B	9.35	0.73 ¹⁹	1.27 ¹⁹	0 ¹⁹
2≡FeOH ^{-0.5} + Fe ²⁺ + 2H ₂ O ↔ ≡(FeOH) ₂ Fe _{III} (OH) ₂ ⁻¹ + 2H ⁺	B + e ⁻	-8.74	0.17 ¹⁹	-0.17 ¹⁹	0 ¹⁹
2≡FeOH ^{-0.5} + ≡Fe ₃ O ^{-0.5} + Fe ²⁺ ↔ ≡(FeOH) ₂ (Fe ₃ O) Fe _{II} ^{+0.5}	T	6.99	1.26 ¹⁹	0.74 ¹⁹	0 ¹⁹

^aThe CD factors are referenced from previous CD-MUSIC model studies. ^bref 19 taken from Hiemstra, T. & van Riemsdijk, W. H., 2007. ^cref 28 taken from Geysels, B. et al., 2025. ^dref 38 taken from Hiemstra, T. & Zhao, W., 2016. ^eref 42 taken from Rahnemaie, R. & Hiemstra, T., 2007. ^fM and B refer to monodentate and bidentate complexes, respectively.

Table 2. Estimated Log K and CD Values for Gly, PO₄, and Fe²⁺ Sorption on the Coordinated Surface Sites of Ferrihydrite^{abcdef}

complexation equation	form	log K	Δz ₀	Δz ₁	Δz ₂
≡FeO _c H ^{-0.5} + H ⁺ ↔ ≡FeO _c H ₂ ^{+0.5}		8.01	1	0	0
≡FeO _c H ^{-0.5} + H ⁺ ↔ ≡FeO _c H ₂ ^{+0.5}		8.01	1	0	0
≡Fe ₃ O ^{-0.5} + H ⁺ ↔ ≡Fe ₃ OH ^{+0.5}		8.01	1	0	0
≡FeO _c H ^{-0.5} + H ⁺ + Cl ⁻ ↔ ≡FeO _c H ₂ ^{+0.5} ...Cl ⁻		7.74	1	-1	0
≡FeO _c H ^{-0.5} + K ⁺ ↔ ≡FeO _c H ^{-0.5} ...K ⁺		-0.6	0	1	0
≡FeO _c H ^{-0.5} + H ⁺ + Cl ⁻ ↔ ≡FeO _c H ₂ ^{+0.5} ...Cl ⁻		7.74	1	-1	0
≡FeO _c H ^{-0.5} + K ⁺ ↔ ≡FeO _c H ^{-0.5} ...K ⁺		-0.6	0	1	0
≡Fe ₃ O ^{-0.5} + H ⁺ + Cl ⁻ ↔ ≡Fe ₃ OH ^{+0.5} ...Cl ⁻		7.74	1	-1	0
≡Fe ₃ O ^{-0.5} + K ⁺ ↔ ≡Fe ₃ O ^{-0.5} ...K ⁺		-0.6	0	1	0
2≡Fe _c OH ^{-0.5} + 2H ⁺ + Gly ⁻³ ↔ ≡Fe ₂ Gly ⁻² (c) + 2H ₂ O	B	22.15	0.69 ²⁸	-0.95 ²⁸	-0.74 ²⁸
≡FeOH ^{-0.5} + 2H ⁺ + Gly ⁻³ ↔ ≡FeHGly ^{-1.5} + H ₂ O	MH	17.40	0.28 ²⁸	-0.54 ²⁸	-0.74 ²⁸
2≡Fe _c OH ^{-0.5} + 3H ⁺ + Gly ⁻³ ↔ ≡Fe ₂ HGly ⁻¹ (c) + 2H ₂ O	BH	26.82	0.75 ²⁸	-0.01 ²⁸	-0.74 ²⁸
≡FeOH ^{-0.5} + 3H ⁺ + Gly ⁻³ ↔ ≡FeH ₂ Gly ^{-0.5} + H ₂ O	MH ₂	21.79	0.41 ²⁸	0.33 ²⁸	-0.74 ²⁸
2≡Fe _c OH ^{-0.5} + 2H ⁺ + PO ₄ ⁻³ ↔ ≡Fe ₂ PO ₄ ⁻² (c) + 2H ₂ O	B	28.31	0.46 ^{38,42}	-1.46 ^{38,42}	0 ^{38,42}
≡FeOH ^{-0.5} + 2H ⁺ + PO ₄ ⁻³ ↔ ≡FeHPO ₄ ^{-1.5} + H ₂ O	MH	23.88	0.28 ^{38,42}	-1.28 ^{38,42}	0 ^{38,42}
2≡Fe _c OH ^{-0.5} + 3H ⁺ + PO ₄ ⁻³ ↔ ≡Fe ₂ HPO ₄ ⁻¹ (c) + 2H ₂ O	BH	33.52	0.65 ³⁸	-0.65 ³⁸	0 ³⁸
≡FeOH ^{-0.5} + 3H ⁺ + PO ₄ ⁻³ ↔ ≡FeH ₂ PO ₄ ^{-0.5} + H ₂ O	MH ₂	30.22	0.33 ³⁸	-0.33 ³⁸	0 ³⁸
≡FeOH ^{-0.5} + Fe ²⁺ ↔ ≡FeOHFe _{II} ^{+1.5}	M	6.41	1.00	1.00	0
2≡Fe _c OH ^{-0.5} + Fe ²⁺ ↔ ≡(FeOH) ₂ Fe _{II} ⁺¹ (c)	B	10.13	0.73 ¹⁹	1.27 ¹⁹	0 ¹⁹
2≡Fe _c OH ^{-0.5} + Fe ²⁺ + 2H ₂ O ↔ ≡(FeOH) ₂ Fe _{III} (OH) ₂ ⁻¹ (c) + 2H ⁺	B + e ⁻	-8.40	0.17 ¹⁹	-0.17 ¹⁹	0 ¹⁹
2≡Fe _c OH ^{-0.5} + ≡Fe ₃ O ^{-0.5} + Fe ²⁺ ↔ ≡(FeOH) ₂ (Fe ₃ O) Fe _{II} ^{+0.5} (c)	T	5.41	1.26 ¹⁹	0.74 ¹⁹	0 ¹⁹
≡Fe ₃ O ^{-0.5} + Fe ²⁺ ↔ ≡Fe ₃ OFe _{II} ^{+1.5}	M	6.41	1.00	1.00	0

^aThe CD values are referenced from previous CD-MUSIC model studies. ^bref 19 Taken from Hiemstra, T. & van Riemsdijk, W. H., 2007. ^cref 28 Taken from Geysels, B. et al., 2025. ^dref 38 Taken from Hiemstra, T. & Zhao, W., 2016. ^eref 42 Taken from Rahnemaie, R. & Hiemstra, T., 2007. ^fM and B refer to monodentate and bidentate complexes, respectively.

spectrophotometer (UVSbio, Mettler Toledo, Switzerland) via the molybdenum blue method,³³ respectively. Prior to Fe²⁺ analysis, samples were acidified with 1 M HCl to stabilize Fe²⁺.³⁴ The Fe²⁺ concentration was then determined using the ferrozine assay.³⁵

2.3. Titration and CD-MUSIC Modeling

The surface properties of goethite and ferrihydrite were characterized by acid–base titration using a titrator (836 Titrando, Metrohm, Switzerland). For each mineral, separate suspensions were prepared for titration at each background electrolyte concentration (0.01 and

0.1 M KCl). HCl (0.01 M) and KOH (0.01 M) solutions were used as titrants. For a given suspension, a complete forward (base) and backward (acid) titration cycle was performed on the same sample. Each titration began by acidifying the suspension to pH 4 with HCl, followed by a titration with KOH at a rate of 0.2 mL/min. At each step, the suspension was allowed to stabilize until the pH drift was less than 1 mV/min. To prevent CO₂ interference, the titration system was continuously purged with N₂. The total solid suspension surface area at the start of the titration was at least 10 m².³⁶ Raw data were processed based on theoretical blank titrations corrected for proton

Table 3. Comparison of CD-MUSIC Parameters in This Study and Literature^{a,b}

mineral	SSA (m ² /g)	pH _{pzc}	N _{s1} (site/nm ²)	N _{s2} (site/nm ²)	C ₁ (F/m ²)	C ₂ (F/m ²)	ref
goethite	94	9.3	3.45	2.7	0.83	0.74	28
	85	/	/	/	0.92	0.92	42
	23	9	4.78	2.08	1.86	3	43
	42	9.25	3.45	2.7	1.65	0.74	this study
ferrihydrite	615	8.1	5.8	1.4	1.15	0.9	38
	611	/	5.8	1.4	/	/	32
	229	8.7	6	1.2	0.74	0.93	44
	611	8.01	5.8	1.4	1.15	0.9	this study

^aFor ferrihydrite, N_{s1} is the sum of site densities for two singly coordinated groups ($\equiv\text{FeO}_c\text{H} + \equiv\text{FeO}_c\text{H}$). ^bN_{s1} and N_{s2} indicate the site densities for singly and triply coordinated surface groups, respectively.

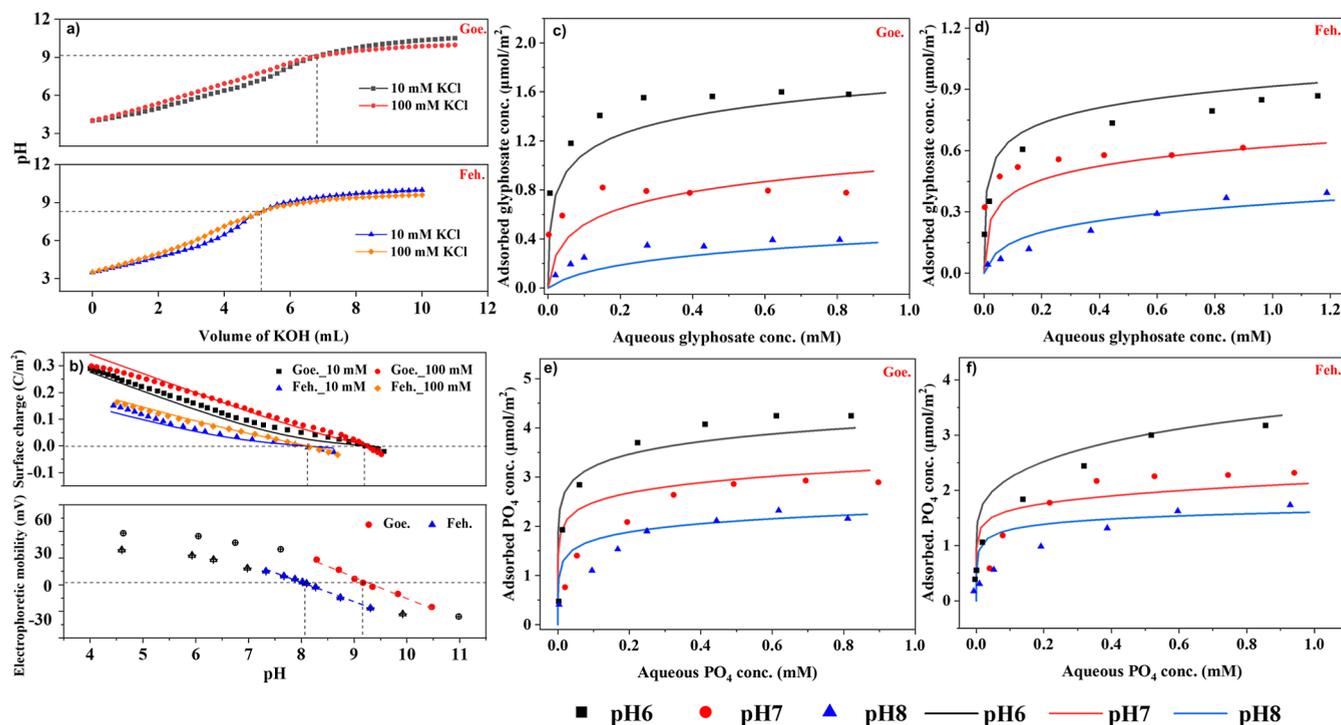


Figure 1. Surface characterization and model validation for goethite and ferrihydrite. (a) Acid–base titration curves at 10, 100 mM KCl; (b) charging behaviors (10, 100 mM KCl) and electrophoretic mobility (5 mM KCl). Adsorption isotherms of (c,d) Gly, and (e,f) PO₄ on goethite and ferrihydrite at pH 6, 7, and 8. Dots are experimental results and curves are simulated results.

balance in the presence of goethite. The isoelectric point (pH_{IEP}) of the minerals was determined using the Zetasizer Nano ZSP (Panalytical, Malvern, United Kingdom) by plotting zeta potential as a function of pH and regressing the linear portion of the data sets. The pH_{IEP} was subsequent compared with point of zero charge (pH_{pzc}) derived from titration.

The CD-MUSIC model quantitatively analyzes ligand binding affinities at mineral/water interfaces, incorporating electrostatic effects within the Stern layer, as well as the capacitances of the inner and outer-layer, into its framework. The Stern layer is partitioned into two distinct regions, represented by planes 0, 1, and 2. Consideration of both singly and triply coordinated sites on the mineral surface is typically required for iron minerals. For the synthesized goethite, the surface sites were represented by singly ($\equiv\text{FeOH}$) and triply ($\equiv\text{Fe}_3\text{O}$) coordinated surface groups, which reflect the predominant terminal and triple-bridge oxygen configurations on its predominant crystal faces, as established by crystallographic and spectroscopic studies. Its site densities (N_s) of 3.45 and 2.7 sites nm⁻², respectively, were applied in the CD-MUSIC model.³⁷ For ferrihydrite, site densities of 5.8 and 1.4 sites nm⁻² were used, following the CD-MUSIC model developed by Hiemstra and Zhao.³⁸ The $\equiv\text{FeOH}$ sites were further divided into two distinct types based on their

capacity to form different surface complexes: edge-sharing sites ($\equiv\text{FeO}_c\text{H}$, N_s = 3.0 sites nm⁻²), which can form both mono- and bidentate complexes, and corner-sharing sites ($\equiv\text{FeO}_c\text{H}$, N_s = 2.8 sites nm⁻²), which can form only monodentate complexes. The parameters log K for H⁺, K⁺, and Cl⁻, as well as the capacitance (C₁ for inner, C₂ for outer) values for the extended Stern layer model, were optimized using UCODE coupled to a modified version of FITEQL2.³⁹ The optimization followed a stepwise strategy: initial values were taken from the characterized mineral properties and published data for analogous systems; subsequently, only a constrained set of parameters were adjusted to achieve the best fit to the specific mineral charging behavior. The results are provided in Tables 1 and 2.

In modeling the Gly and PO₄ sorption, deprotonated monodentate (M), deprotonated bidentate (B), protonated monodentate (MH), protonated bidentate (BH), and doubly protonated monodentate (MH₂) complexes were considered for both iron hydroxides. Notably, bidentate complexes are assumed to form with $\equiv\text{FeO}_c\text{H}$ on ferrihydrite.³⁸ In the modeling of Fe²⁺ sorption, electron transfer from Fe²⁺ to Fe³⁺ was reported during the sorption process (B + e⁻).¹⁹ Additionally, tridentate (T) complexes can form via interaction with two $\equiv\text{FeOH}$ and one $\equiv\text{Fe}_3\text{O}$ site (double edge sharing).²⁰ All considered surface complexation reactions and their corresponding

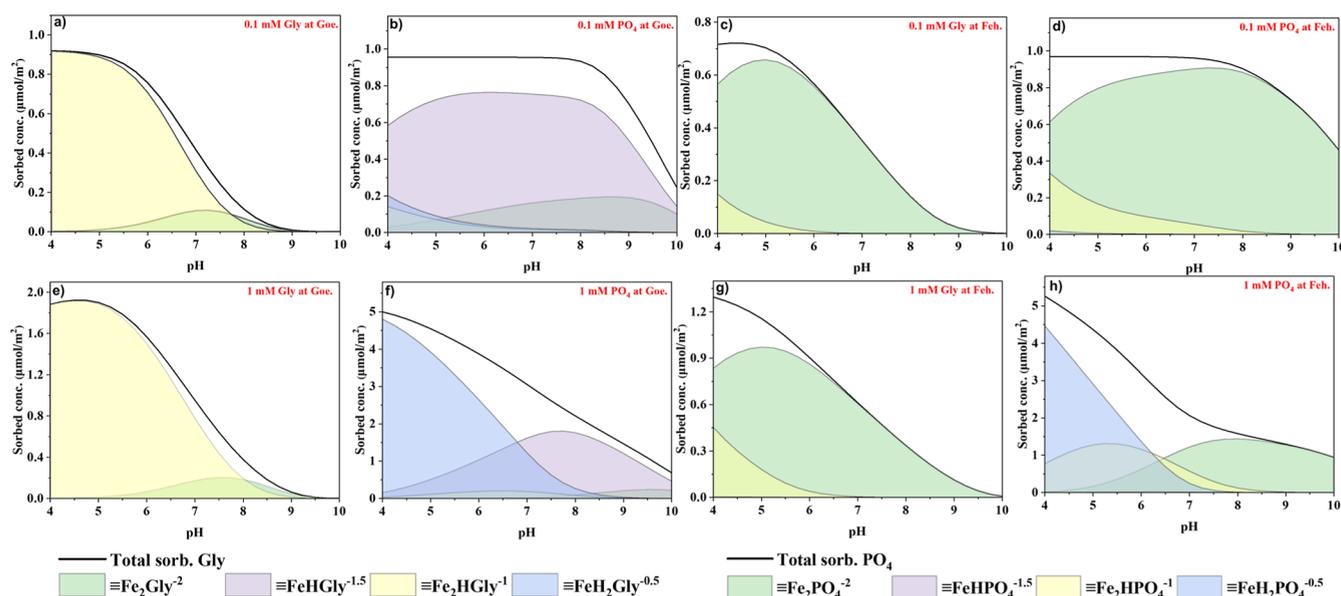


Figure 2. Simulated adsorption edge and surface complexation of Gly and PO₄ on goethite and ferrihydrite at initial conc. of (a–d) 0.1 mM and (e–h) 1 mM.

log *K* for Gly, PO₄, and Fe²⁺ were optimized using Visual Minteq 3.1 coupled with PEST.^{40,41} Atmospheric CO₂ level of 410 ppm was incorporated into all CD-MUSIC simulations to account for its influence on aqueous carbonate speciation and proton balance under experimentally relevant conditions. The charge distribution to the surface, inner Stern plane, and outer Stern plane (Δz_0 , Δz_1 , Δz_2) are sourced from the papers reported by Hiemstra and co-workers and are summarized in Tables 1 and 2. The aqueous speciation and possible surface complexation structures of Gly and PO₄ are illustrated in Figure S5, Table S1 and Figure S6.

3. RESULTS AND DISCUSSION

3.1. Comparative Analysis of CD-MUSIC Parameters

In previous studies, the CD-MUSIC model has been applied to goethite and ferrihydrite, taking into account factors such as SSA, pH_{pzc} , N_s , and *C*. Table 3 compares CD-MUSIC parameters from the literature with the final optimized set obtained in this study. For our systems, the SSA, pH_{pzc} values were fixed inputs, while the parameters N_s and *C* are the results of optimizing the model against our acid–base titration data. For goethite, our synthesized material exhibited a significantly lower SSA compared to earlier reports (42 m²/g vs 90 m²/g) by Geysels & Hiemstra²⁸ and Rahnamaie & Hiemstra,⁴² which directly led to comparatively lower overall adsorption in our experiments. Nonetheless, our study helps fill the knowledge gap regarding goethite with moderate SSA, providing a useful reference for understanding the environmental behavior of goethite across different particle sizes. This finding is consistent with observations by Wang et al.⁴³ for microsized goethite (23 m²/g), which also exhibited differences in *C* values compared to high-SSA goethite—a trend echoed in our modeled *C*₁ value being higher than those from studies using ~90 m²/g goethite. For ferrihydrite, our fitted CD-MUSIC parameters show good consistency with those reported by Hiemstra & Zhao³⁸ and by Gustafsson & Antelo.³² For our modeling, the SSA of ferrihydrite was set to 611 m²/g, representing the typical surface area of fresh, nonfreeze-dried ferrihydrite rather than the BET-measured value of the freeze-dried sample (275.2 m²/g), to accurately reflect the material

state in our adsorption experiments. It is worth noting, however, that Antelo et al.⁴⁴ investigated aged ferrihydrite and applied BET-SSA, which may account for some differences in model parameters. Although our ferrihydrite parameters are comparable to literature values, this study aims to apply CD-MUSIC model to Gly adsorption on ferrihydrite—both individually and in competition with PO₄ and Fe²⁺—thereby providing novel mechanistic insight into Gly retention under complex conditions.

3.2. Mineral Surface Charge Behavior and Single Solute Sorption

Figure 1 illustrates the surface properties and adsorption behavior of goethite and ferrihydrite through integrated experimental and modeling approaches. The potentiometric titration curves in 10 and 100 mM KCl (Figure 1a) intersect at 9.25 for goethite and 8.01 for ferrihydrite, providing an estimate of the pH_{pzc} .⁴⁵ Crucially, these pH_{pzc} values are in excellent agreement with the pH_{IEP} independently determined from the pH-dependence of electrophoretic mobility (Figure 1b). This consistency between two distinct experimental methods supports the assignment of the pH_{pzc} . While the CD-MUSIC model optimization yields discernible log *K* values for K⁺ and Cl[−] (Tables 1 and 2), the convergence of the titration intersection and the pH_{IEP} indicates that, at the macroscopic level, the net effect of electrolyte ion adsorption on the proton-conditional surface charge is minimal around the pH_{pzc} for our systems. The surface charging behaviors in Figure 1b) were described by the CD-MUSIC model, revealing distinct electrical layer structures: goethite exhibits higher inner capacitance ($C_1 = 1.65 \text{ C/m}^2$) but lower outer capacitance ($C_2 = 0.74 \text{ C/m}^2$) compared to ferrihydrite ($C_1 = 1.15 \text{ C/m}^2$, $C_2 = 0.90 \text{ C/m}^2$). The fundamental difference in their pH_{pzc} reflects the contrasting surface protonation behavior between the crystalline goethite and amorphous ferrihydrite structures. The pH_{pzc} -derived log *K* for H⁺ were subsequently used in the adsorption modeling. A broader comparison of pH_{pzc} with literature values is provided in Table S2.

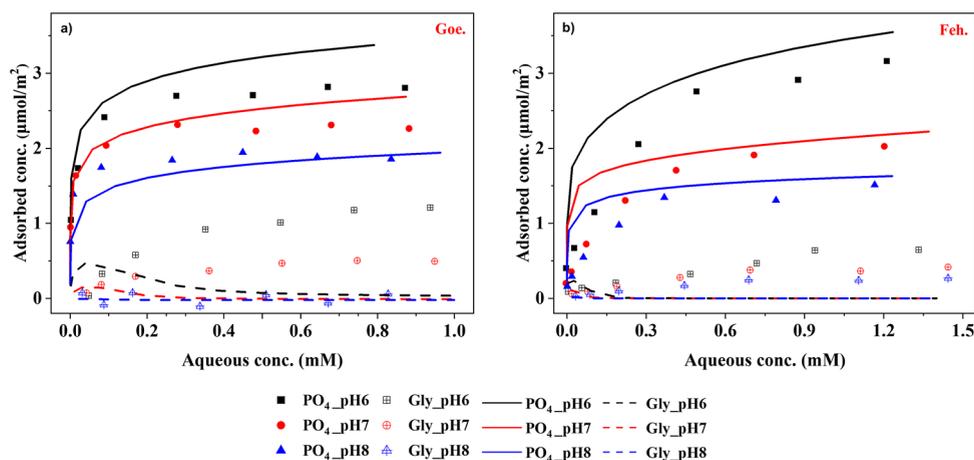


Figure 3. Competitive binary sorption from equimolar Gly-PO₄ mixtures on (a) goethite and (b) ferrihydrite at pH 6, 7, and 8. Dots are experimental results and curves are simulated results.

Figure 1c–f presents multiple adsorption isotherms for Gly and PO₄ determined at different pH values,^{6–8} demonstrating decreasing adsorption with increasing pH^{6–8} at iron hydroxides. This behavior primarily results from the decrease in positive surface charge as pH approaches the pH_{pzc} leading to weakened electrostatic interactions between the mineral surfaces and anionic adsorbates. PO₄ exhibits higher uptake compared to Gly onto minerals, consistent with previous reports on soils.^{46,47} The close match between experimental data points and CD-MUSIC model simulations (solid curves) confirms model reliability in predicting surface complexation behavior under varying chemical conditions (evaluated by R^2 , Table S4). It is noted that the model shows slight deviations from experimental data at low concentrations. These discrepancies may be related to the enhanced influence of intrinsic surface site heterogeneity at low coverage, which is challenging to fully capture with an averaged site density model. Nevertheless, across the full range of conditions studied, the CD-MUSIC model provides a quantitatively superior description of the adsorption data compared to alternative diffuse layer model (DLM) formulations (see comparative analysis in the Supporting Information). This comprehensive data set provides critical parameters for understanding contaminant adsorption mechanisms in iron hydroxide-dominated environmental systems. Additionally, surface complexation corresponding to adsorption isotherms was simulated in Figures S7 and S8. Surface sites occupied during adsorption, as calculated from experimental data, were compared with N_s implemented in the CD-MUSIC model (Table S3).

3.3. Surface Complexation of Single Gly, PO₄

Modeling parameters listed in Tables 1 and 2 can be used to calculate the speciation of Gly and PO₄ complexes on goethite and ferrihydrite surfaces as a function of pH and initial concentration. Figure 2a–h explore adsorption behaviors at representative initial concentrations of 0.1 mM (low loading) and 1 mM (high loading).

Our model reveals that, across the entire pH range and at the studied concentrations (0.1–1 mM), Gly predominantly presents on the goethite surface in BH and B forms, with the contribution of the B complex becoming increasingly prominent as pH increases (Figure 2a,e). This trend is mechanistically driven by the deprotonation of the—NH₂

group, which is the sole structural difference between the B and BH configurations (See Figure S6). These findings are consistent with the most recent results reported by Geysels and Hiemstra for Gly adsorption on goethite,²⁸ who also considered a comprehensive set of surface species including B, BH, MH, and MH₂. In our study, the contributions of MH and MH₂ complexes were optimized to be negligible, whereas they constituted a minor proportion in the cited work.²⁸ The inclusion of MH₂ in our goethite model followed the same comprehensive approach to ensure a complete description of protonation equilibria. This discrepancy in minor species distribution is not due to the CD values—which were sourced from Rahnemaie et al.⁴² but is rationally explained by distinct physicochemical properties of the goethite samples. Specifically, our material has a lower SSA (41.9 m²/g vs 94 m²/g) and a different C₁ (1.65 F/m² vs 0.83 F/m²) (Table 3). Similarly, Gly is mainly complexed on ferrihydrite in BH and B forms, with B complex being the dominant species (Figure 2c,g). Earlier research⁴⁸ using ATR-FTIR and XANES for adsorption speciation analysis yielded results consistent with ours, indicating that at low pH, binding occurs preferentially via the B complex, while at common pH, both B and MH complexes are predominant. Notably, on both iron minerals, the Gly surface species do not exhibit strong concentration dependence.

In contrast, MH and MH₂ complexes contribute significantly to PO₄ adsorption. As shown in Figure 2b, the MH complex dominates PO₄ adsorption at low loading conditions, while the MH₂ complex is the secondary contributor under acidic pH. The proportion of B rises with increasing pH, becoming comparable to the MH complex at pH 10. Numerous DFT studies^{17,42,49,50} have investigated PO₄ adsorption on goethite, all of which concluded that M bonding is the optimal geometry for PO₄. However, a critical limitation of DFT lies in its reliance on finite molecular computations, which hinders accurate prediction of high-concentration adsorption behavior. Our analysis reveals a notable shift at high loading (Figure 2f), where MH₂ replaces MH as the dominant complex in acidic pH. This transition is mechanistically attributed to reduced electrostatic potential in the Stern layer under elevated surface loading, which enhances proton adsorption and facilitates MH₂ complex formation. Intriguingly, PO₄ adsorption on ferrihydrite at low loading is predominantly mediated by B and BH complexes. (Figure 2d), which is consistent with the results

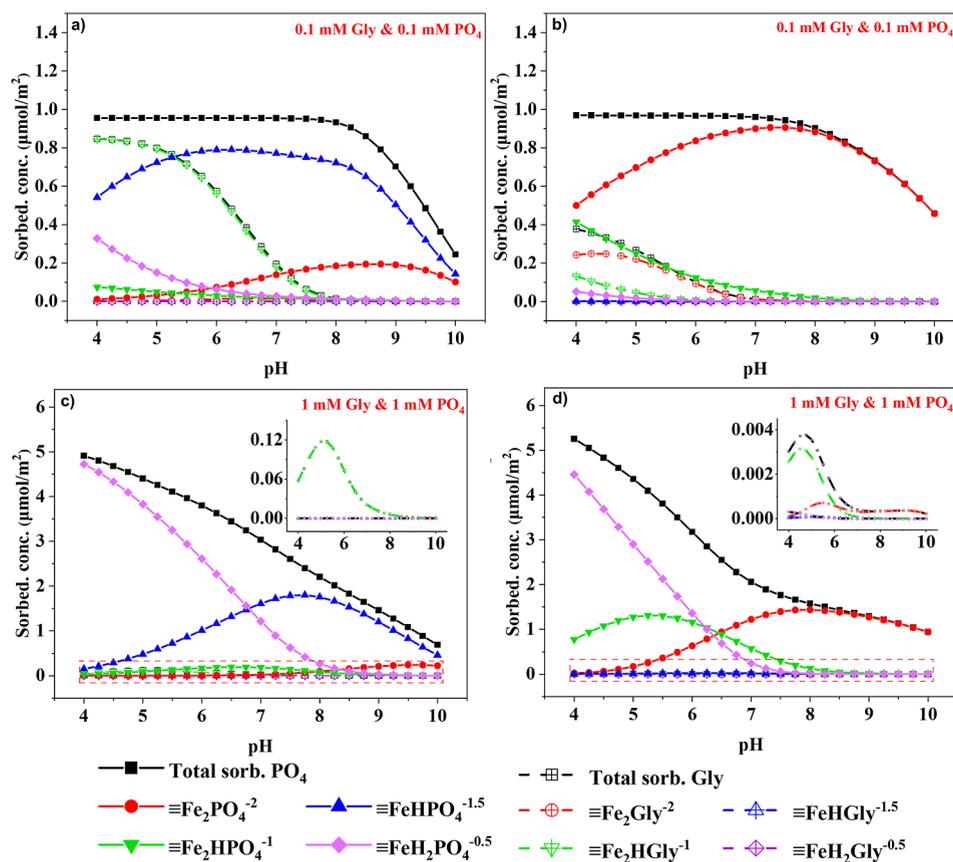


Figure 4. Simulated adsorption edge and surface complexation of Gly vs PO_4 on goethite and ferrihydrite at initial conc. of (c,e) 0.1 mM and (d,f) 1 mM. Dots are experimental results and curves are simulated results.

previously reported by Antelo et al.⁴⁴ Discrepancies emerge at high pH, where their model suggests minor M-type contributions—a divergence traceable to differences in material properties between studies: pH_{pzc} (8.0 vs 8.1), SSA (611 m^2/g vs 229 m^2/g), and C_1 (1.15 F/m^2 vs 0.74 F/m^2). At high loading, the contribution of the MH_2 complex to PO_4 adsorption on ferrihydrite under acidic conditions becomes non-negligible, as illustrated in Figure 2h). Additionally, CD-MUSIC model predictions indicate that PO_4 adsorption remains less pH-independent under low-concentration acidic conditions, and both Gly and PO_4 exhibit generally decreasing adsorption trends across the pH 4–10 range in other scenarios, consistent with typical anion adsorption behavior on iron minerals.

Based on CD-MUSIC model simulation results, it is clear that the adsorption and complexation of Gly and PO_4 on iron hydroxides are jointly influenced by environmental pH, mineral type, and initial concentration in the system.

3.4. Binary Sorption and Complexation of Gly and PO_4

To analyze competitive adsorption between Gly and PO_4 on the two iron minerals, we simultaneously added equal concentrations of both adsorbates to mineral suspensions at different pH.^{6–8} The results are shown in Figure 3a,b, where dots represent experimental data and curves denote CD-MUSIC model simulations. In the competitive adsorption system, both Gly and PO_4 exhibited decreasing adsorption with rising pH, while PO_4 consistently demonstrated stronger affinity than Gly across all pH conditions. Interestingly, the model accurately predicted PO_4 adsorption trends but showed

discrepancies in Gly adsorption behavior. Experimentally, despite the presence of strongly adsorbing PO_4 , Gly partially remained adsorbed on iron minerals, where the CD-MUSIC model predicted full Gly desorption under competitive conditions. This persistent uptake under strong competition is consistent with the irreversible adsorption behavior observed for Gly on other mineral surfaces such as $\gamma\text{-Al}_2\text{O}_3$,⁵¹ where dedicated desorption and mass balance experiments confirmed a nonlabile fraction. We also considered the possibility of Gly decomposition contributing to the apparent retention; however, this is unlikely under our experimental conditions (pH 6–8, 12 h) as our analytical method (IC-ECD) did not detect significant levels of its primary abiotic degradation product (AMPA), and such conditions are not known to promote rapid degradation. Therefore, the observed discrepancy suggests that some Gly-mineral complexes exhibit irreversible binding characteristics not captured by the equilibrium CD-MUSIC model.

The surface speciation in Figure 4a–d provides insight into the model underestimation of Gly adsorption in binary systems. In 0.1 mM competitive systems, (Figure 4a,c) Gly maintained measurable adsorption competition against PO_4 . Although Gly adsorption decreased, Gly surface species remained consistent with single-adsorbate systems across pH conditions. (Figure 2a,c) At higher concentrations (1 mM, Figure 4b,d), PO_4 maintained its single system behavior and surface species, (Figure 2f,h) while Gly adsorption dropped sharply with peak adsorption occurring near pH 5. The model predicted a shift in dominant Gly surface species to BH-type complexes under these conditions (Figure 4b,d) and

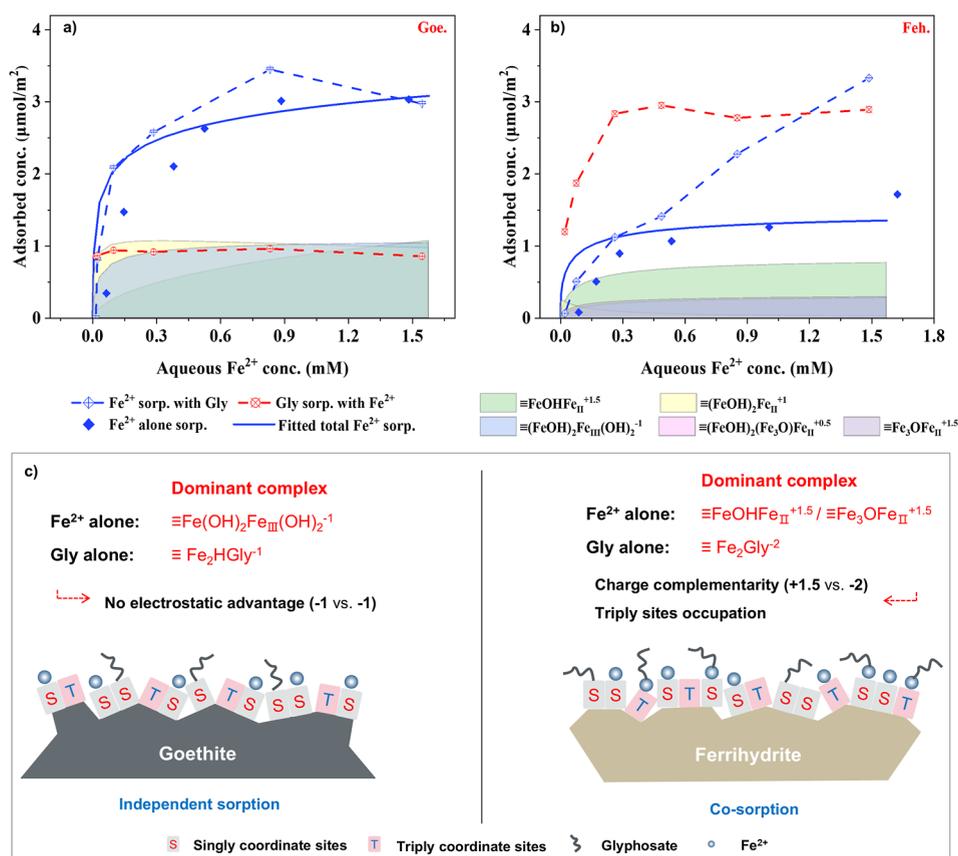


Figure 5. Simulated Fe²⁺ adsorption isotherm and surface complexation on (a) goethite and (b) ferrihydrite under isothermal conditions at pH 7. Solid lines represent Fe²⁺ single adsorption, while dashed lines indicate Fe²⁺ adsorption in competition with pre-equilibrated Gly (0.24 mM). Stacked bars depict the relative contributions of Fe²⁺ surface species at pH 7 along the single adsorption isotherm. (c) Schematic illustration of the proposed binary sorption mechanism for Fe²⁺ and Gly at iron minerals.

potentially reflected irreversible complexation at high concentrations.

Until now, no existing models have fully predicted the Gly-PO₄ competitive adsorption patterns observed here. Previous CD-MUSIC studies have primarily focused on individual Gly²⁸ or PO₄^{26,43,44} adsorption on iron minerals, as well as selected inorganic binary systems such as Mg²⁺/Ca²⁺-PO₄⁵², As(V)-PO₄³² whereas the Gly-PO₄ organic-inorganic competition remained less modeled. Recent work⁵³ is beginning to address this complexity by developing a CD-MUSIC framework for Gly adsorption on ferrihydrite in the presence of PO₄, and under a Gly concentration of 0.47 mM, a noticeable decline in model performance is found. Moreover, a complementary study⁵⁴ from the same group demonstrated that glyphosate adsorption on metal (hydr)oxide surfaces is subject to pronounced steric hindrance effects, which become increasingly important at higher surface coverages. Such steric constraints can limit access to reactive surface sites and alter surface complexation configurations, mechanisms that are not explicitly accounted for in conventional CD-MUSIC formulations. Together, these findings highlight the inherent challenges of modeling large, multifunctional organic adsorbates in competitive systems and help explain why the equilibrium CD-MUSIC model employed in this study underestimates the irreversible fraction of Gly binding. This limitation likely arises from the omission of steric effects and molecular crowding phenomena in addition to classical site competition and electrostatic interactions. By extending CD-

MUSIC to this unexplored domain, our work provides new experimental and modeling insights for organic-inorganic competitive adsorption systems.

3.5. Binary Sorption and Surface Complexation of Fe²⁺

Fe²⁺ is widely present in natural environments, especially abundant under anoxic and reducing conditions, while it readily oxidizes to Fe³⁺ in oxygenated waters.⁹ Fe²⁺ primarily exists in soil-water systems of wetlands, paddy fields, and sediments under reducing conditions, playing a crucial role in biogeochemical cycling.^{55–57} Fe²⁺-sorbed on iron hydroxides can significantly enhance the transformation of organic pollutants and stabilize organic carbon through adsorption and coprecipitation mechanisms. Previous studies have shown that Fe²⁺ barely reacts with ligands (Cl⁻, OH⁻) in natural waters at ambient pH.⁵⁸ Given that Fe²⁺ forms complexes with Gly in solution,¹⁰ (Figure S9) with no reports of solid Fe²⁺-Gly precipitation under analogous conditions, this study investigates the competitive adsorption behavior between Fe²⁺ and Gly under the premise of preadsorption of Gly.

The dissolution of minerals during adsorption was also analyzed in Figure S10, indicating that it was negligible and all measured Fe²⁺ concentrations originated from experimental additions. In Figure 5, the adsorption of Fe²⁺ on iron minerals alone and under Gly competitive conditions is presented, with the complexation of Fe²⁺ illustrated as stacked diagrams. On goethite (Figure 5a), Fe²⁺ adsorption was not affected by the presence of preadsorbed Gly, resulting in a similar overall uptake as when Fe²⁺ was adsorbed alone. At pH 7, Gly

primarily forms the BH complex carrying a -1 charge, while Fe^{2+} mainly coordinates via a B form with two $\equiv\text{FeOH}^{-0.5}$ sites (-1 charge) on the goethite surface (Figure 5c). This leads to no electrostatic advantage via Gly preadsorption to influence Fe^{2+} binding. In contrast, on ferrihydrite (Figure 5b), the adsorption amounts of both Fe^{2+} and Gly increased under coadsorption conditions compared to their single system adsorption. This indicates that Fe^{2+} and Gly mutually enhance their adsorption on ferrihydrite. In this case, Gly predominantly forms the B complex with a -2 charge, while Fe^{2+} mainly binds in the M form ($+1.5$ charge) on ferrihydrite (Figure 5c). This enhanced opposite charge complementarity between the adsorbed species facilitates coadsorption. Furthermore, when optimizing Fe^{2+} adsorption on ferrihydrite using the CD-MUSIC model, we also considered its complexation with $\equiv\text{Fe}_3\text{O}^{-0.5}$. Without including this interaction, the predicted adsorption amounts were significantly lower than the experimental data. Therefore, these sites are also potential coadsorption sites for Fe^{2+} and Gly on ferrihydrite.

Previous studies have reported that other divalent metal ions, such as Ca^{2+} ,^{59,60} Cd^{2+} ,⁶¹ Mg^{2+} ,⁶⁰ Zn^{2+} ,⁶² etc., influence PO_4 adsorption on both goethite and ferrihydrite. According to our charge results for the simulated surface species, one explanation can be that PO_4 forms the MH complex with goethite at pH 7 carrying a -1.5 charge, giving those cations an electrostatic advantage over Fe^{2+} for surface binding (-0.5 charge). The same mechanism can be applied to explain coadsorption behavior on ferrihydrite. As pointed out, the effects of Fe^{2+} on Gly adsorption on iron hydroxides remain rarely studied. Most existing research has focused on how Cu^{2+} ,⁶³ Zn^{2+} ,⁶⁴ Pb^{2+} ⁶⁵ influence Gly adsorption in soils, often showing that these cations enhance the environmental mobility of Gly under most conditions. Therefore, the dual mineral-specific roles exhibited by Fe^{2+} in its adsorption and complexation behaviors, as demonstrated by the application of the CD-MUSIC model in this study, provide new insights and theoretical foundations for understanding Fe^{2+} transport in the environment.

4. ENVIRONMENTAL IMPLICATIONS

To address the competitive effects of PO_4 and Fe^{2+} on Gly adsorption and the associated modeling challenges, we applied the CD-MUSIC framework to experimental data for goethite and ferrihydrite. Our results demonstrate that PO_4 exhibits stronger affinity than Gly for both iron hydroxides, with their surface bonding governed by pH, initial concentration, mineral type, and competitive interactions. This implies that PO_4 fertilization in agroecosystems can thus limit the retention of Gly in iron-rich soils, while natural PO_4 cycling in aquatic and terrestrial systems may critically regulate Gly mobility. The incomplete desorption of Gly under strong PO_4 competition highlights its environmental persistence; the proposed complexation mechanism provides new insights into its long-term fate.

Co-adsorption experiments with Fe^{2+} revealed dual roles: Fe^{2+} enhances Gly adsorption on ferrihydrite but has negligible effects on goethite. As iron hydroxides are key components controlling nutrient and contaminant retention and transport in soils, and Fe^{2+} is predominantly enriched in anoxic environments such as wetlands and rice paddies, our findings deepen the understanding of Gly mobility under reducing conditions. This study also extends the application of the CD-MUSIC model to organic–inorganic competitive systems.

While the model successfully describes the dominant trends and identifies key surface species, it underestimates Gly retention under strong competition. This discrepancy highlights the current limitation of equilibrium-based models in fully capturing mechanisms such as steric hindrance or specific irreversible binding, which may operate at high surface coverage. Overall, the optimized parameters and mechanistic insights provided here constitute a critical advance for modeling Gly fate in complex, iron-rich environments and enhance the predictive capability of SCMs for multicomponent systems.

■ ASSOCIATED CONTENT

Supporting Information

The Supporting Information is available free of charge at <https://pubs.acs.org/doi/10.1021/acsestwater.6c00065>.

Details about the characterization of iron hydroxides, speciation and surface complexation of adsorbates, dissolution of minerals during adsorption, comparison of CD-MUSIC parameters, and fitting comparison between DLM and CD-MUSIC model (PDF)

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Notes

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