

# Estimation of the humidity flux at FINO 1 (North Sea)

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# Outline

Motivation

Humidity effects at the North Sea - FINO 1

Estimate humidity fluxes - original idea

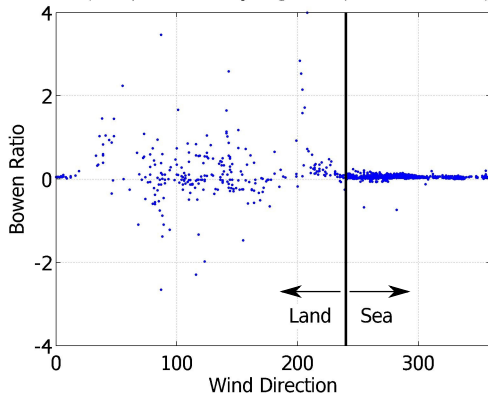
Estimate humidity fluxes - new idea

Conclusions

# Importance of humidity for marine conditions

*Sempreviva & Gryning (1996) - Humidity Fluctuations in the Marine Boundary Layer measured at a coastal site...*

Anholt (Sempreviva & Gryning, 1996), Bowen Ratio (2m)

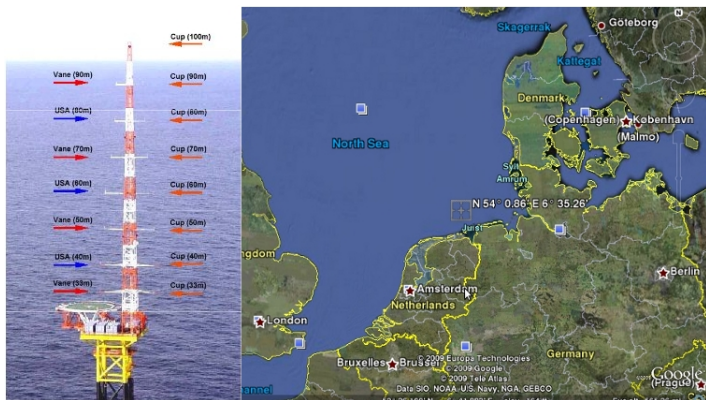


- ▶ Bowen Ratio =  $\frac{c_p \langle wT \rangle}{L_v \langle wq \rangle}$
- ▶ Marine Conditions: (Bowen Ratio  $\approx 0.1$ );

## Motivation - Importance of humidity for stability

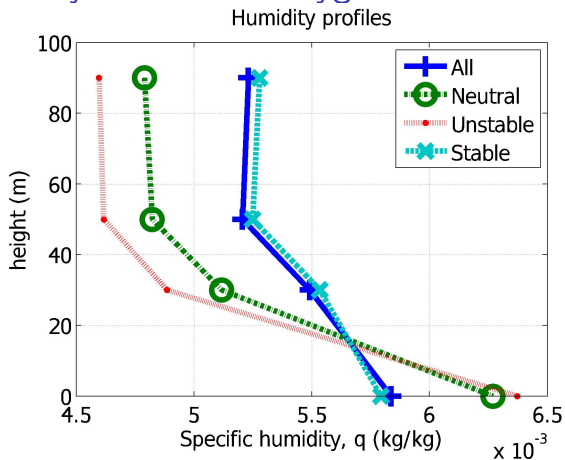
- ▶ Monin-Obukhov stability parameter:  $\frac{z}{L} = \frac{zkg\langle wT_v \rangle}{T_v u_*^3}$
- ▶ Virtual potential heat flux can be estimated from:  
 $\langle wT_v \rangle = \langle wT \rangle + 0.61T \langle wq \rangle$
- ▶ An ultrasonic anemometer measures:  
 $\langle wT_s \rangle = \langle wT \rangle + 0.51T \langle wq \rangle$
- ▶ In dry and/or cold conditions,  $\langle wq \rangle =$  humidity flux and can be neglected: hence  $\langle wT_v \rangle \approx \langle wT_s \rangle$ .
- ▶ For example, Andreas et al. (2006) - Evaluations of the von Karman constant in the atmospheric surface layer, assume  $\langle wT_v \rangle = \langle wT_s \rangle$  in arctic conditions.
- ▶ What then is the magnitude of  $\langle wq \rangle$  at FINO 1?

# FINO 1 - Measurement platform in the North Sea



- ▶ In operation since 2003, approx. 45km North of Borkum Island.
- ▶ Cup anemometers, wind vanes, temperature sensors, ultrasonic anemometers, hygrometers at multiple levels.
- ▶ Data presented here from 2005.

# Humidity at FINO 1 - hygrometers?

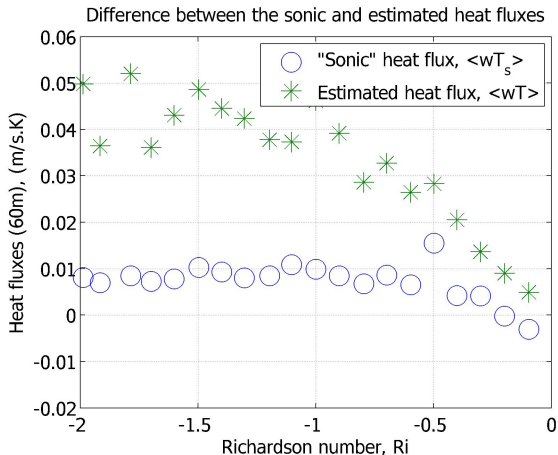


- ▶ Above 50m – Negative humidity flux for stable conditions, positive humidity flux for neutral & unstable conditions.
- ▶ It is unclear whether fluxes driven by local gradients above 50m, or by non-local gradients? Here, sea surface relative humidity assumed = 100%.

## Estimate humidity fluxes - original idea:

- ▶ Assume  $\langle wT_s \rangle = \langle wT \rangle + 0.51T \langle wq \rangle$ .
- ▶ There, use  $Ri = \frac{z}{L}$ , valid in unstable conditions, where  $Ri \approx -\frac{\kappa g z \langle wT \rangle}{Tu_*^3}$
- ▶ Rearranging gives:  $\langle wT \rangle = -\frac{RiT u_*^3}{\kappa z g}$
- ▶ Use humidity dependent sonic measurement to estimate the humidity flux:  $\langle wq \rangle = \frac{1}{0.51T} (\langle wT_s \rangle - \langle wT \rangle)$ .
- ▶ The humidity flux can then supposedly be estimated as a residual between the sonic measured heat flux and the heat flux based on a local gradient.

## Results:



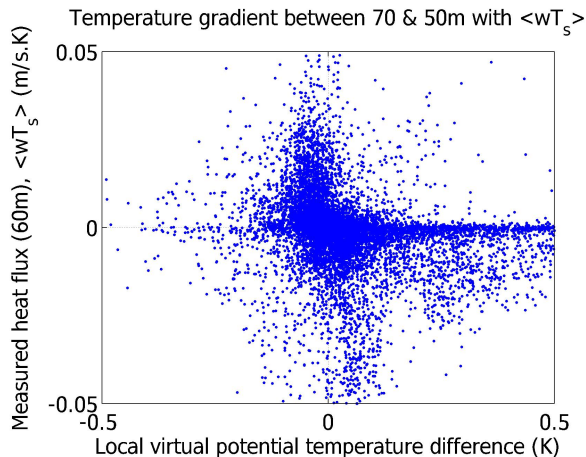
- ▶ Average humidity flux:

$$\overline{\langle wq \rangle} = \frac{1}{0.51T} (\langle wT_s \rangle - \langle wT \rangle) \approx -2 \times 10^{-4}.$$

- ▶ Typical humidity flux (Sempreviva & Gryning, 1996):  $+3 \times 10^{-5}$ .

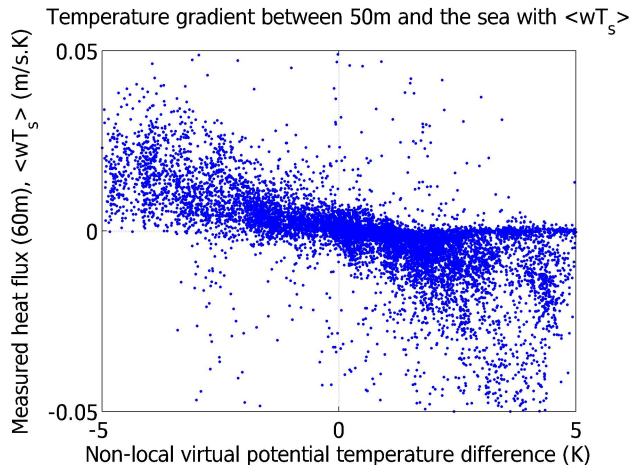


# Heat Flux - Profile Relationship



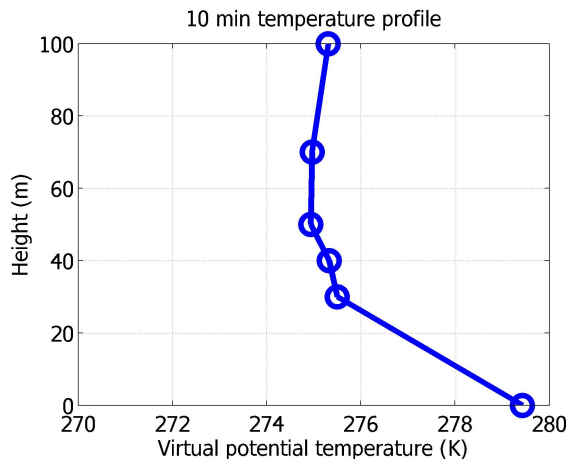
- ▶ Local heat fluxes are poorly correlated with local temperature gradients. (Correlation coefficient: -0.06)
- ▶ Using local gradients to infer fluxes not valid, hence  $Ri \neq \frac{z}{L}$ .

# Heat Flux - Profile Relationship



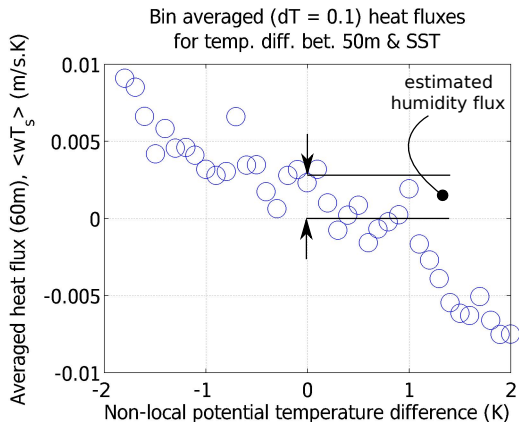
- ▶ Heat fluxes are driven by non-local temperature gradients. (Correlation coefficient: -0.47)

## Example: Temperature profile



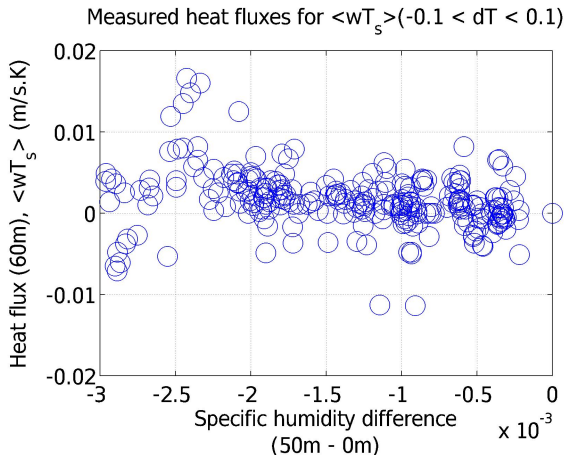
- ▶ At 60m: Locally stable, non-locally unstable.
- ▶ Heat flux,  $\langle wT_s \rangle = +0.013$  (m/s.K)

# Estimation of temperature and humidity fluxes - new idea



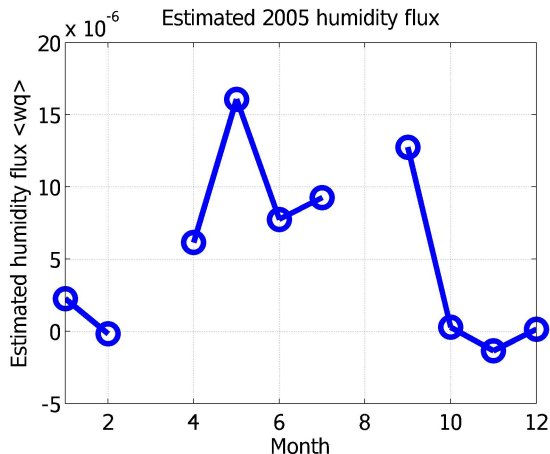
- ▶ If humidity flux were zero, averaged heat flux would pass through the origin when plotted against the potential temperature.
- ▶ Therefore, since  $\langle wT_s \rangle = \langle wT \rangle + 0.51T \langle wq \rangle$ , then  $\langle wT_s \rangle (\Delta\theta = 0) = 0.51T \langle wq \rangle$ .

# Estimated humidity fluxes



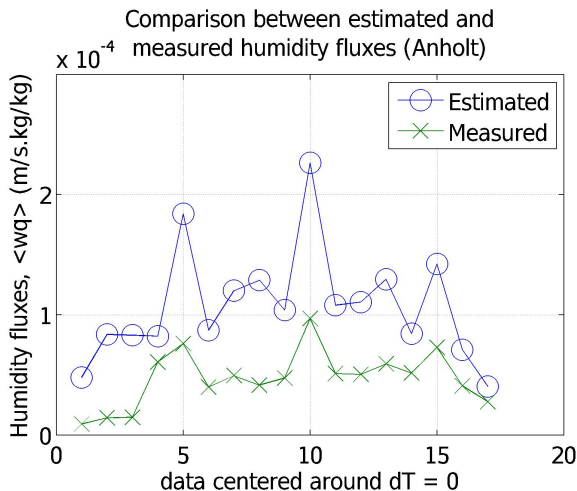
- ▶ Estimated humidity fluxes very weakly correlated with both fine humidity gradients and (shown here) bulk differences.
- ▶ Humidity flux a function of larger atmospheric scales? (ref?)

## Annual variation of humidity flux estimate:



- ▶ October-February:  $\overline{\langle wq \rangle} \approx 0$
- ▶ April-September:  $\overline{\langle wq \rangle} \approx +1 \times 10^{-5}$ . (Sempreviva & Gryning, 1996, June,  $\langle wq \rangle = +3 \times 10^{-5}$ )

# Estimate humidity fluxes - Anholt (Sempreviva & Gryning, 1996)



- ▶ Estimation possibly overestimates humidity flux by factor of 2?

# Conclusions

- ▶ Precise calculation of stability of the marine boundary layer requires direct measurement of the humidity flux since it may not be governed by local gradients.
- ▶ Currently, there is no way of directly measuring humidity fluxes at FINO 1.
- ▶ Estimates indicate that humidity flux is roughly  $1 \times 10^{-5}$  (m/s.kg/kg) during the middle of the year, but could be neglected in colder months.



# Acknowledgements

- ▶ We acknowledge the reception of valuable data and advice from Dr. Anna-Maria Sempreviva which has aided this work greatly.
- ▶ The work was done in the project VERITAS (work package 5 of OWEA) which is funded by the German Ministry of the Environment (BMU) via the PTJ (FKZ 0325060).