

Capabilities and Limitations of Regional Climate Modeling for Hydrological Impact Analysis

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Stakeholders demand delineation of climate change adaptation strategies

- Flood protection measures (adaptation of infrastructure)
- Future hydropower potential (low flows)
- Water availability for agriculture
- Tourism (Skiing...)

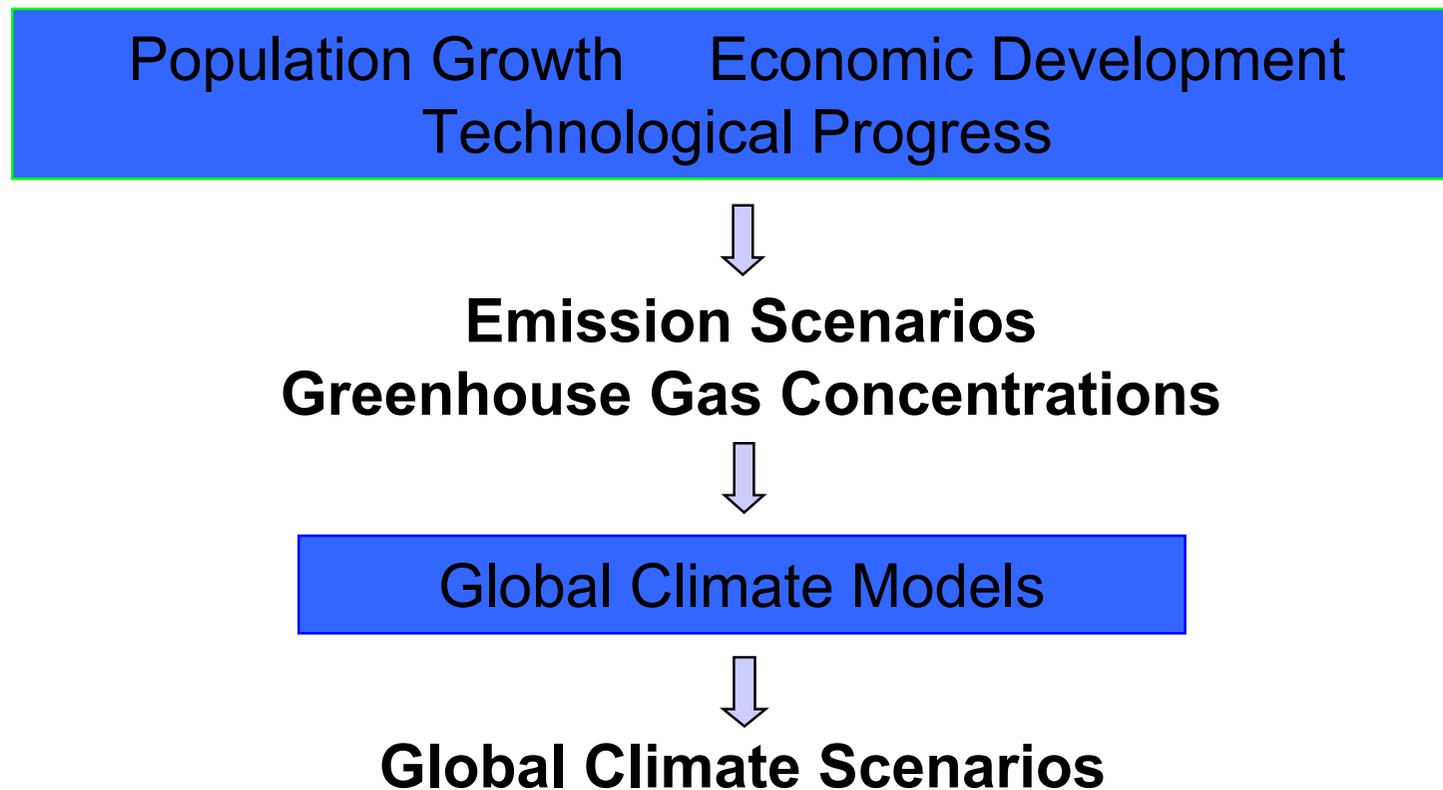
Requirements in hydrological climate change impact investigations

- High resolution spatial and temporal distribution of future temperature & precipitation

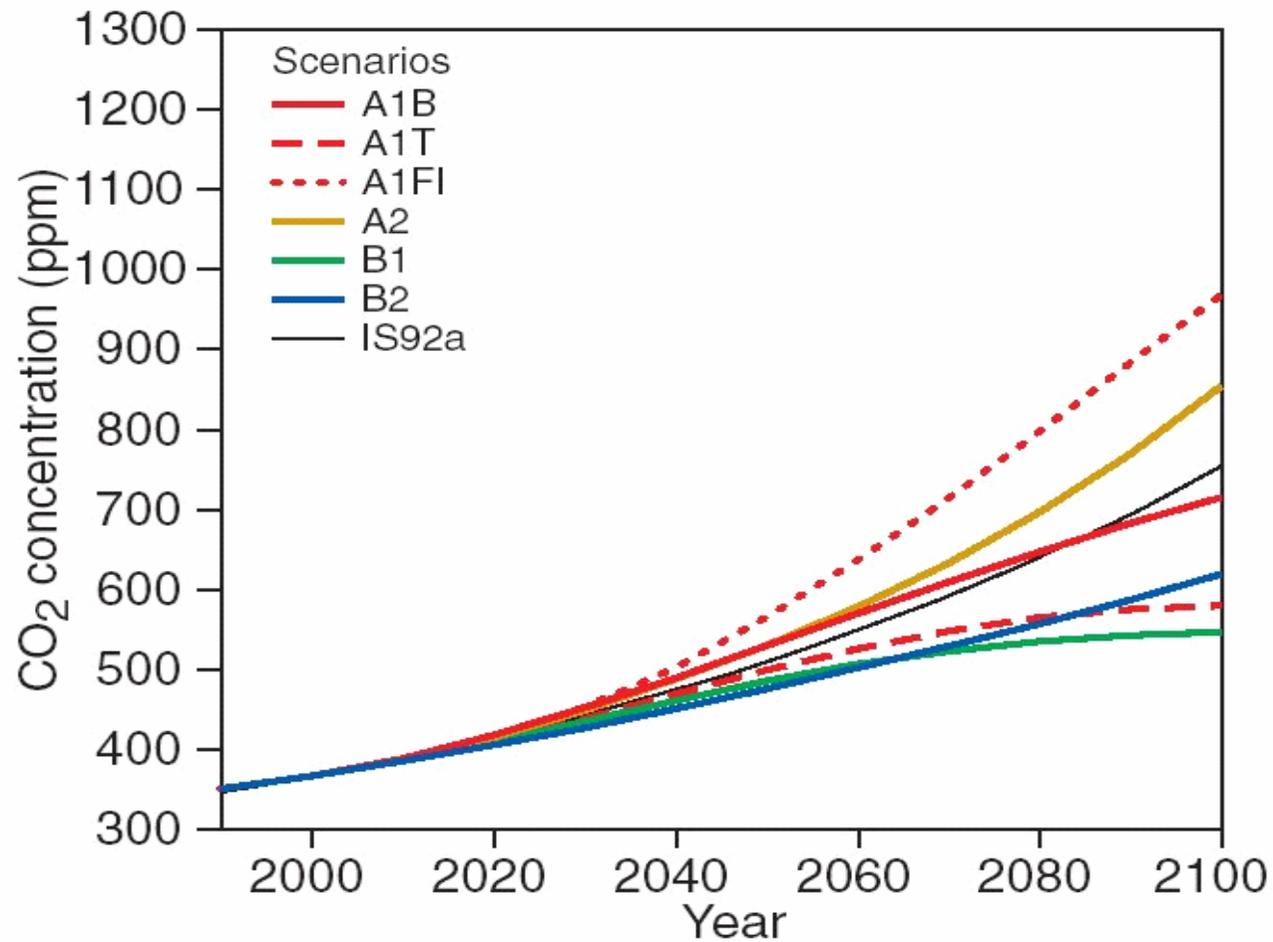
How well are current regional climate predictions suited for that purpose?

Atmosphere is primary driver of hydrological processes: brief review

- Spatial and temporal distribution of precipitation and temperature
- Global warming impacts water vapor carrying capacity & evaporation/condensation
- Projections of future terrestrial water availability must consider future atmospheric state
- Close linkage atmosphere, precipitation & land surface:
 - energy & water fluxes at land surface: latent & sensible heat fluxes
 - terrain elevation: e.g. precipitation generation by orographic blocking
 - soil: long term memory of previous precipitation & temperature
- **Interlinked atmosphere & land surface process description & analysis necessary**



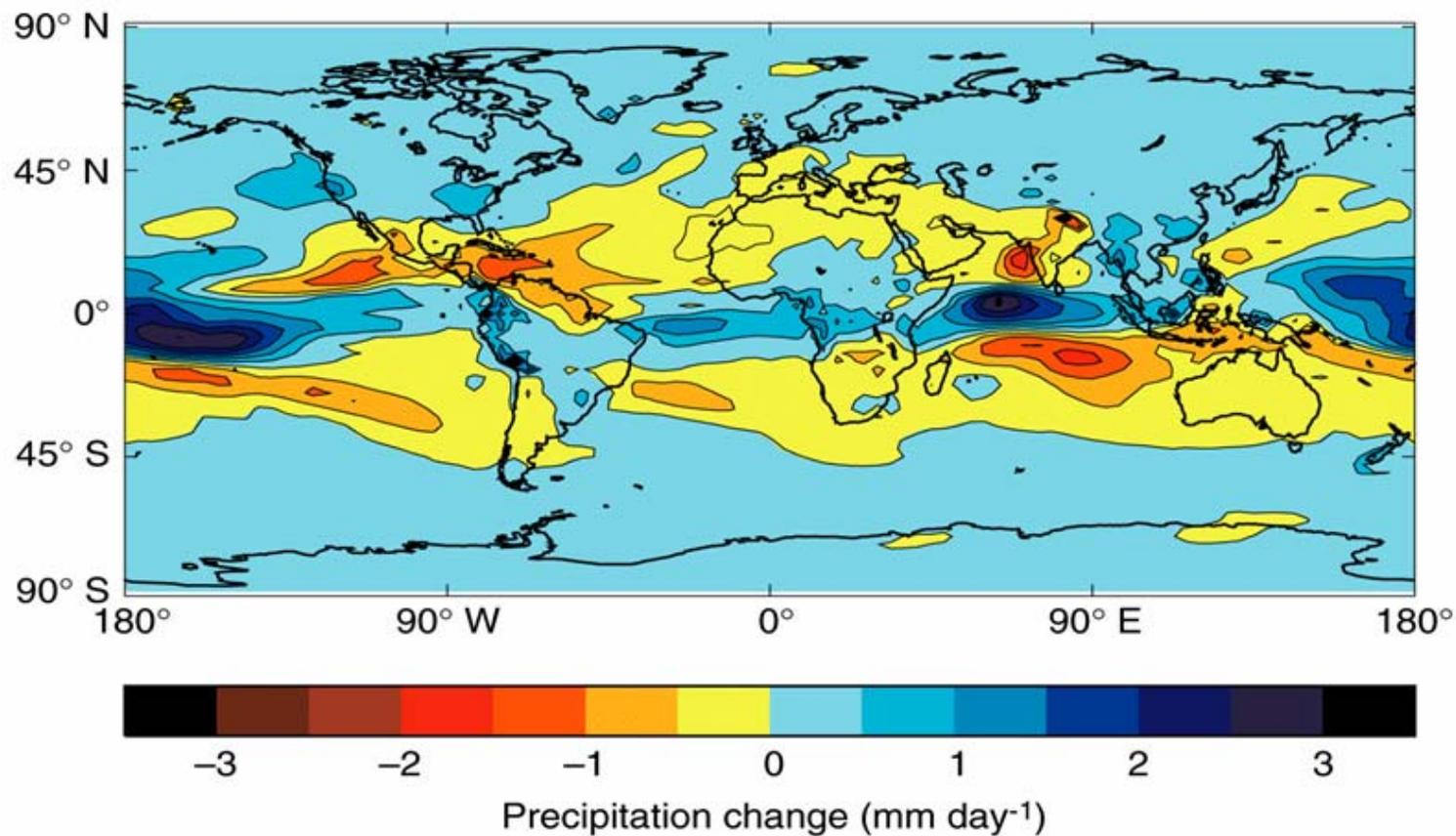
CO₂ concentrations



Global climate models

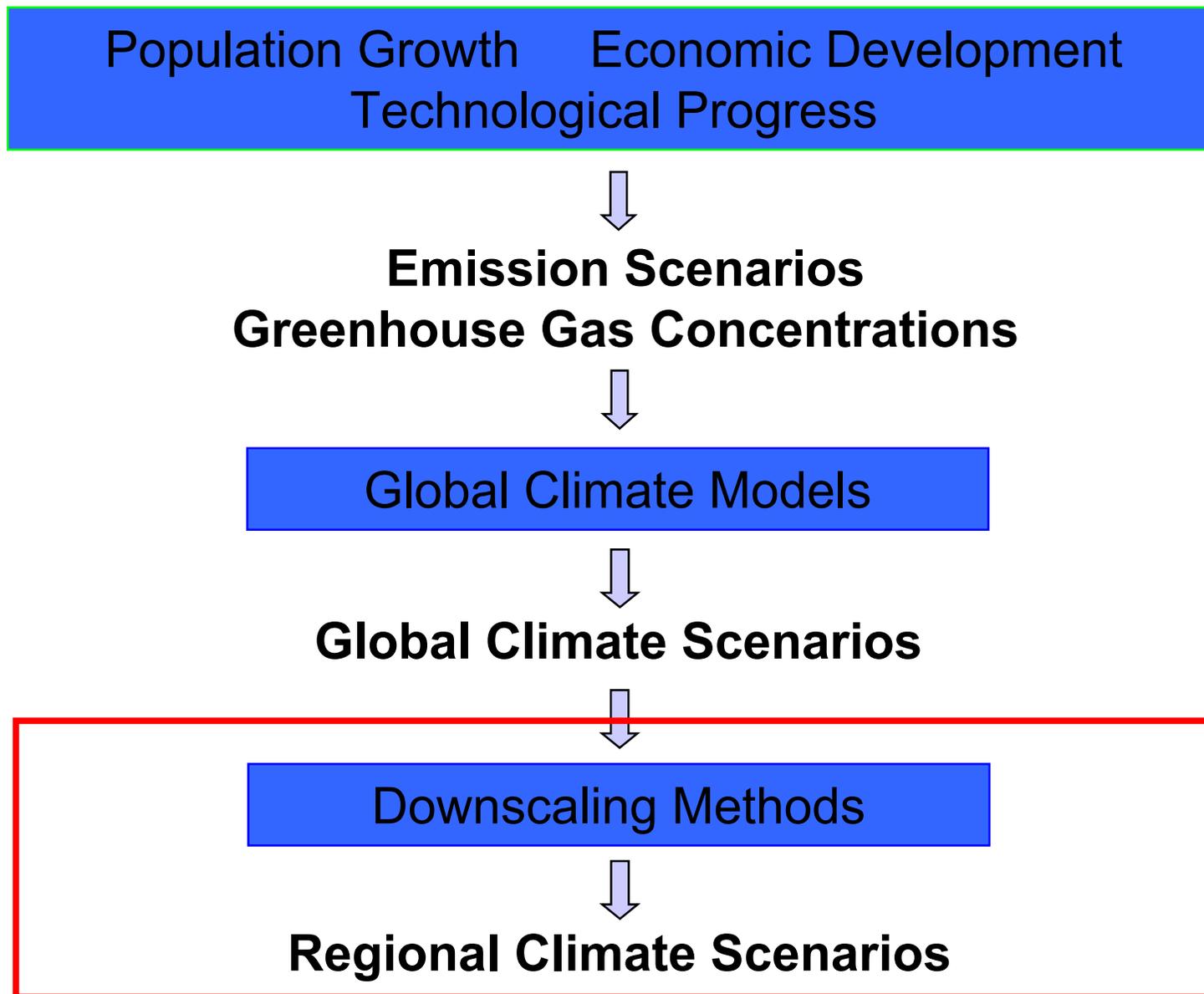
- physically based (e.g. conservation laws for energy, momentum, water/humidity variables) but: parameterizations for subgridscale processes (turbulence, convective precipitation, ...)
- resolution usually $\approx 100\text{-}300$ km
- designed to reproduce & project global trends
- limited potential for regional analysis

Global Climate Scenarios: Projected Changes in Annual Precipitation for the 2050s



Hadley Centre
for Climate
Prediction and
Research

⇒ Resolution too coarse for regional impact analysis !



Statistical downscaling

- statistical relations between large scale forcing & station observed variables
- e.g. via multiple regression, canonical correlation analysis, circulation pattern analysis
- computationally efficient
- climate change information derived only at station locations
- gridded fields obtained by spatial interpolation between station locations
- persistence of statistical relations under changing climate conditions assumed

Dynamical downscaling

- 3-dim regional atmospheric models (RCM) based on conservation laws, physical relations, parameterizations for subgridscale processes
- nested approach: global model provides lateral boundaries of regional model
- computationally expensive
- examples: *nonhydrostatic* CLM, MM5, WRF
hydrostatic REMO, RegCM, HIRHAM, PRECIS
- usually coupled atmosphere-land surface modeling systems

Momentum conservation

$$\frac{\partial \vec{v}}{\partial t} + (\vec{v} \cdot \nabla) \vec{v} = -f \vec{k} \times \vec{v} - \nabla \Phi - \frac{1}{\rho_a} \nabla p_a + \frac{\eta_a}{\rho_a} \nabla^2 \vec{v} + \frac{1}{\rho_a} (\nabla \cdot \rho_a \mathbf{K}_m \nabla) \vec{v}$$

Energy conservation

$$\frac{\partial \theta_v}{\partial t} + (\vec{v} \cdot \nabla) \theta_v = \frac{1}{\rho_a} (\nabla \cdot \rho_a \mathbf{K}_h \nabla) \theta_v + \frac{\theta_v}{c_{p,d} T_v} \sum_{n=1}^N \frac{dQ_n}{dt}$$

Gas law

$$p = \frac{nR^*T}{V}$$

Air mass conservation

$$\frac{\partial \rho_a}{\partial t} + \nabla \cdot (\vec{v} \rho_a) = 0$$

Conservation water mass

$$\begin{aligned} \frac{\partial q_v}{\partial t} + (\vec{v} \cdot \nabla) q_v &= \frac{1}{\rho_a} (\nabla \rho_a \mathbf{K}_h \nabla) q_v + R_{evap} - R_{cond} - R_{iini} - R_{idep/sub} \\ \frac{\partial q_c}{\partial t} + (\vec{v} \cdot \nabla) q_c &= \frac{1}{\rho_a} (\nabla \rho_a \mathbf{K}_h \nabla) q_c + R_{cond} + R_{iini} + R_{idep/sub} - R_{aconv} - R_{accr} \\ \frac{\partial q_r}{\partial t} + (\vec{v} \cdot \nabla) q_r &= \frac{1}{\rho_a} (\nabla \rho_a \mathbf{K}_h \nabla) q_r - R_{evap} + R_{aconv} + R_{accr} - \frac{\partial V_f \rho_a g q_r}{\partial t} \end{aligned}$$

Energy conservation at land surface

$$\begin{aligned} L_v E + H + G &= SW_{net} + LW_{net} \\ &= (1 - \alpha) SW \downarrow + LW \downarrow - \epsilon \sigma_B T_{surf}^4 \end{aligned}$$

Soil temperature diffusion

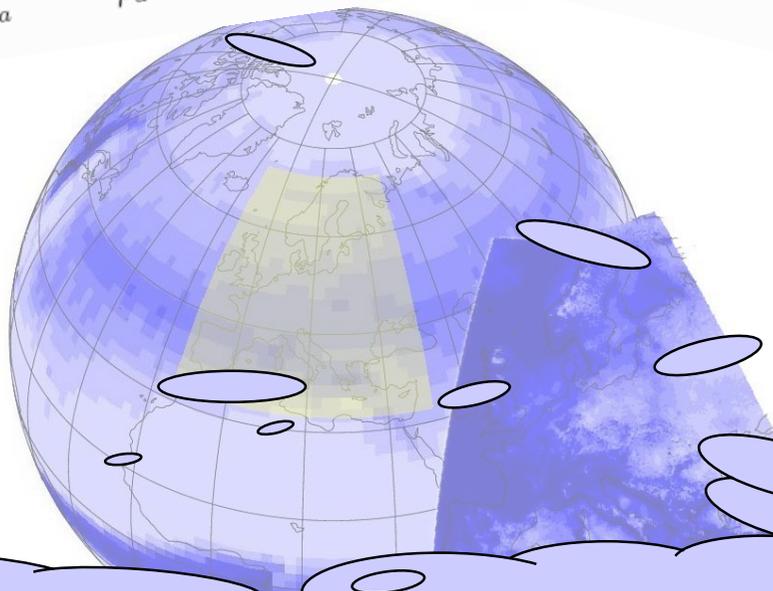
$$C(\Theta) \frac{\partial T_s}{\partial t} = \frac{\partial}{\partial z} \left[K_t(\Theta) \frac{\partial T_s}{\partial z} \right]$$

Precipitation physics

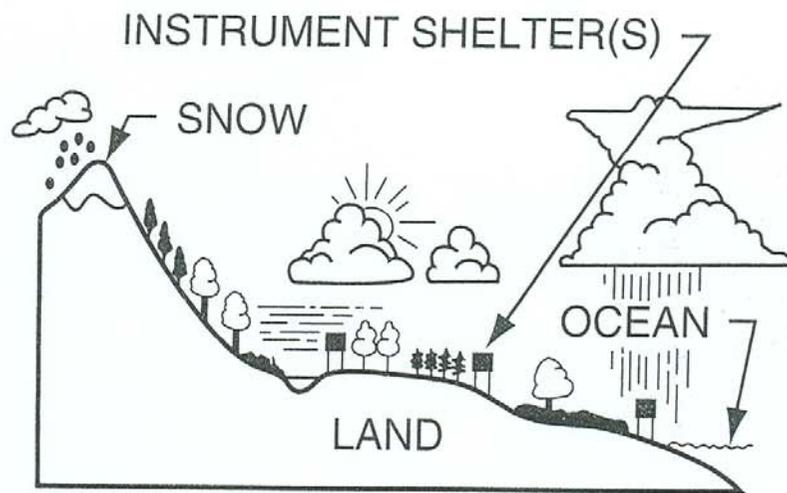
$$R_{evap(rain)} = \frac{2\pi N_{0r} (S_w - 1)}{A_r + B_r} \left[\frac{0.78}{\Lambda_r^2} + 0.32 \left(\frac{a_r \rho}{\eta_a} \right)^{1/2} S_c^{1/3} \frac{\Gamma(5/2 + b_r/2)}{\Lambda_r^{5/2 + b_r/2}} \right]$$

Soil water infiltration

$$\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial z} \left[D(\Theta) \frac{\partial \theta}{\partial z} \right] + \frac{\partial k(\Theta)}{\partial z}$$

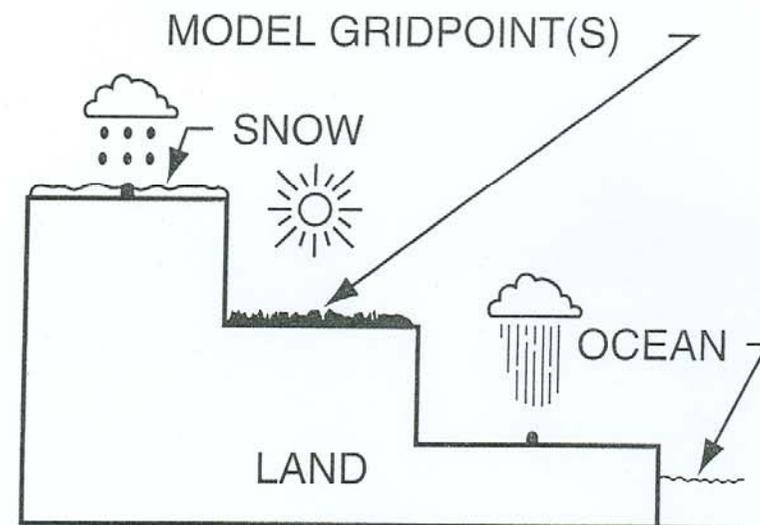


REAL WORLD



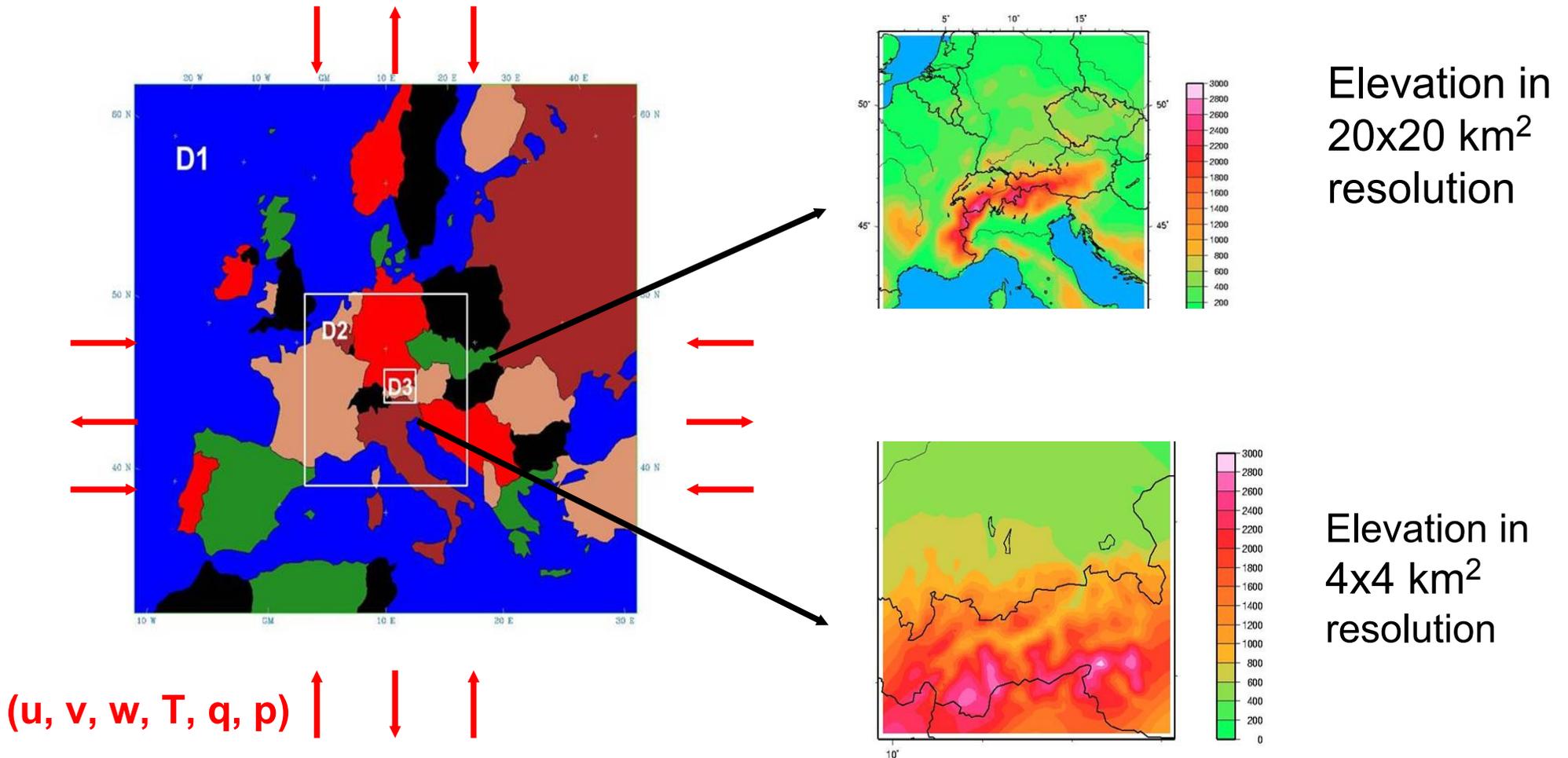
Vs.

MODEL WORLD

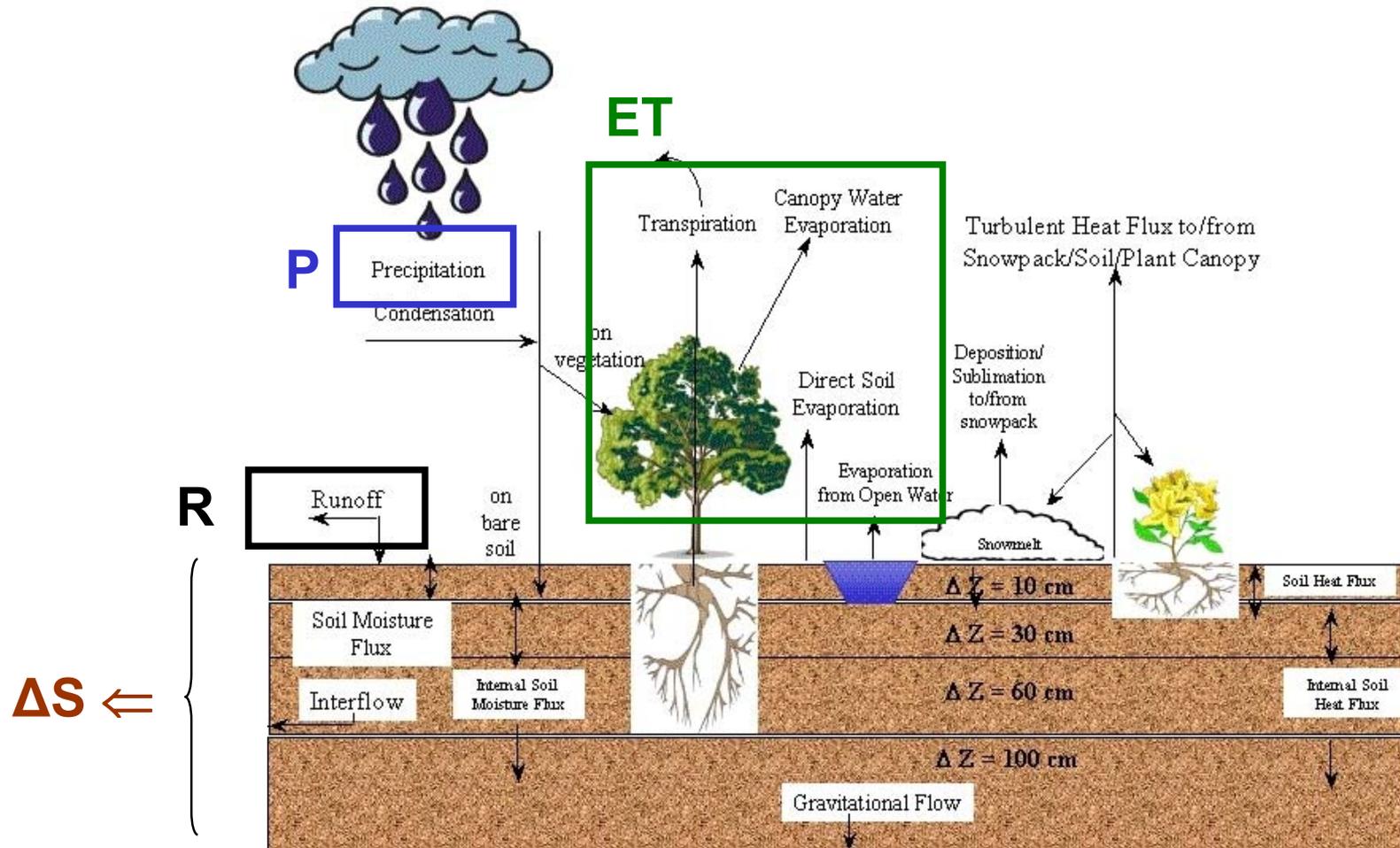


Dynamical Downscaling: Nested Simulations

RCM: forced by global model (boundary- & initial value problem)
High spatial resolution \Rightarrow detailed consideration of elevation



Dynamical downscaling: SVAT model as lower boundary at every grid point



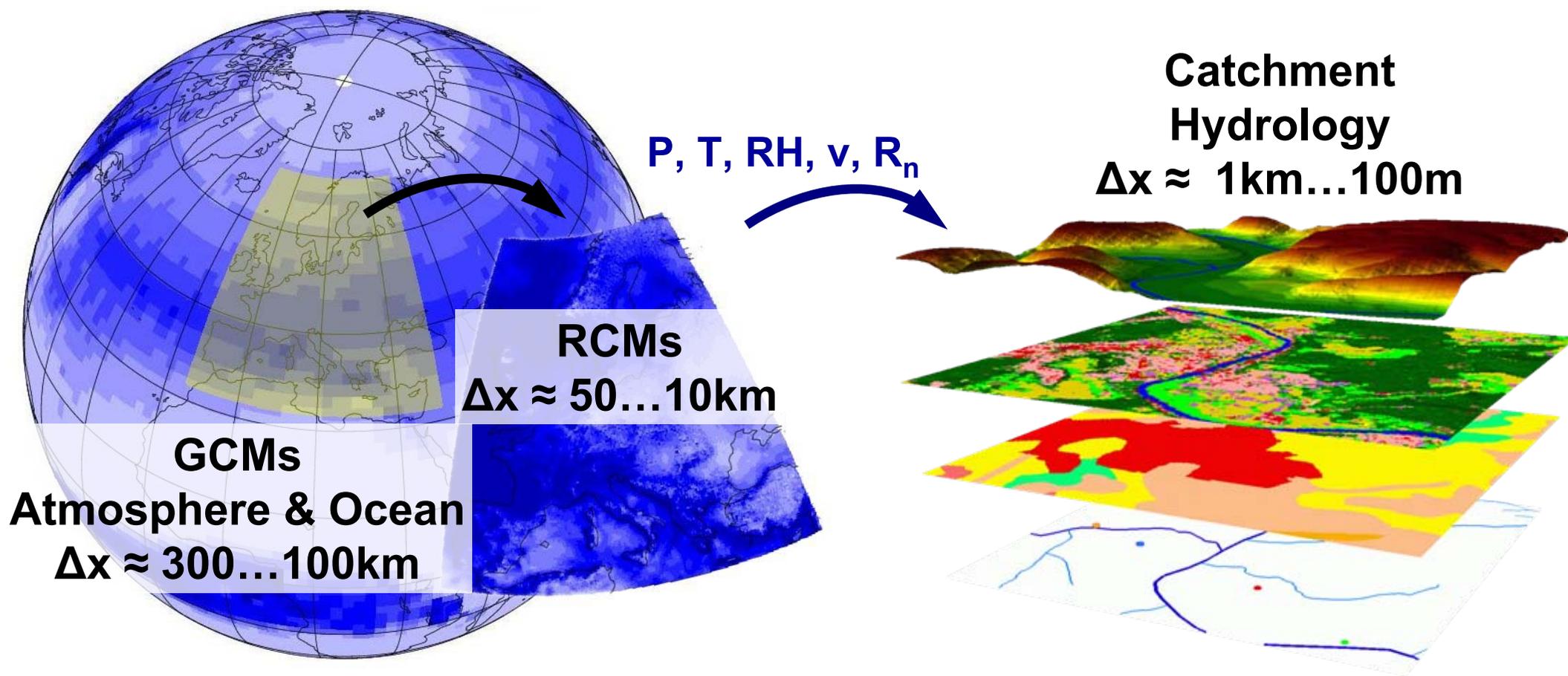
Soil-Vegetation-Atmosphere-Transfer (SVAT) Model

SVAT- hydro models (designed for atmospheric feedback purposes)

- full energy balance (soil heat & sensible heat fluxes)
- 2-way interaction with PBL
- focus on vertical water fluxes (soil moisture, ET)

“Traditional”- hydro models (designed for pure hydrological applications):

- vertical + lateral water fluxes, surface runoff routing
- deeper soils considered, finer vertical & horizontal resolutions
- often groundwater interaction, often extensions for reactive flow & transport, erosion, etc.
- **but of course: depending on specific model choice**



Evapotranspiration

$$E = \frac{s(T_a)[K + L] + \gamma K_E \rho_w \lambda_v v_a e_{sat}(T_a)[1 - RH]}{\rho_w \lambda_v [s(T_a) + \gamma]}$$

Infiltration

$$\frac{\partial}{\partial z'} \left(K_h(\theta) \frac{\partial \Psi(\theta)}{\partial z'} \right) - \frac{\partial K_h(\theta)}{\partial z'} = \frac{\partial \theta}{\partial t}$$

Conductivity

$$K_h(\theta) = K_{h,sat} \sqrt{\frac{\theta - \theta_{res}}{\phi - \theta_{res}}} \left[1 - \left(1 - \left(\frac{\theta - \theta_{res}}{\phi - \theta_{res}} \right)^{1/m} \right)^m \right]^2$$

Groundwater Flow

$$-\frac{\partial}{\partial x} \left(K_{hx} \frac{\partial h}{\partial x} \right) - \frac{\partial}{\partial y} \left(K_{hy} \frac{\partial h}{\partial y} \right) - \frac{\partial}{\partial z} \left(K_{hz} \frac{\partial h}{\partial z} \right) + R = -S_s \frac{\partial h}{\partial t}$$

Routing

$$v_h = M \left(\frac{Q_h / v_h}{B_h + \frac{2Q_h}{v_h B_h}} \right)^{2/3} \sqrt{I}$$

Wave Retention

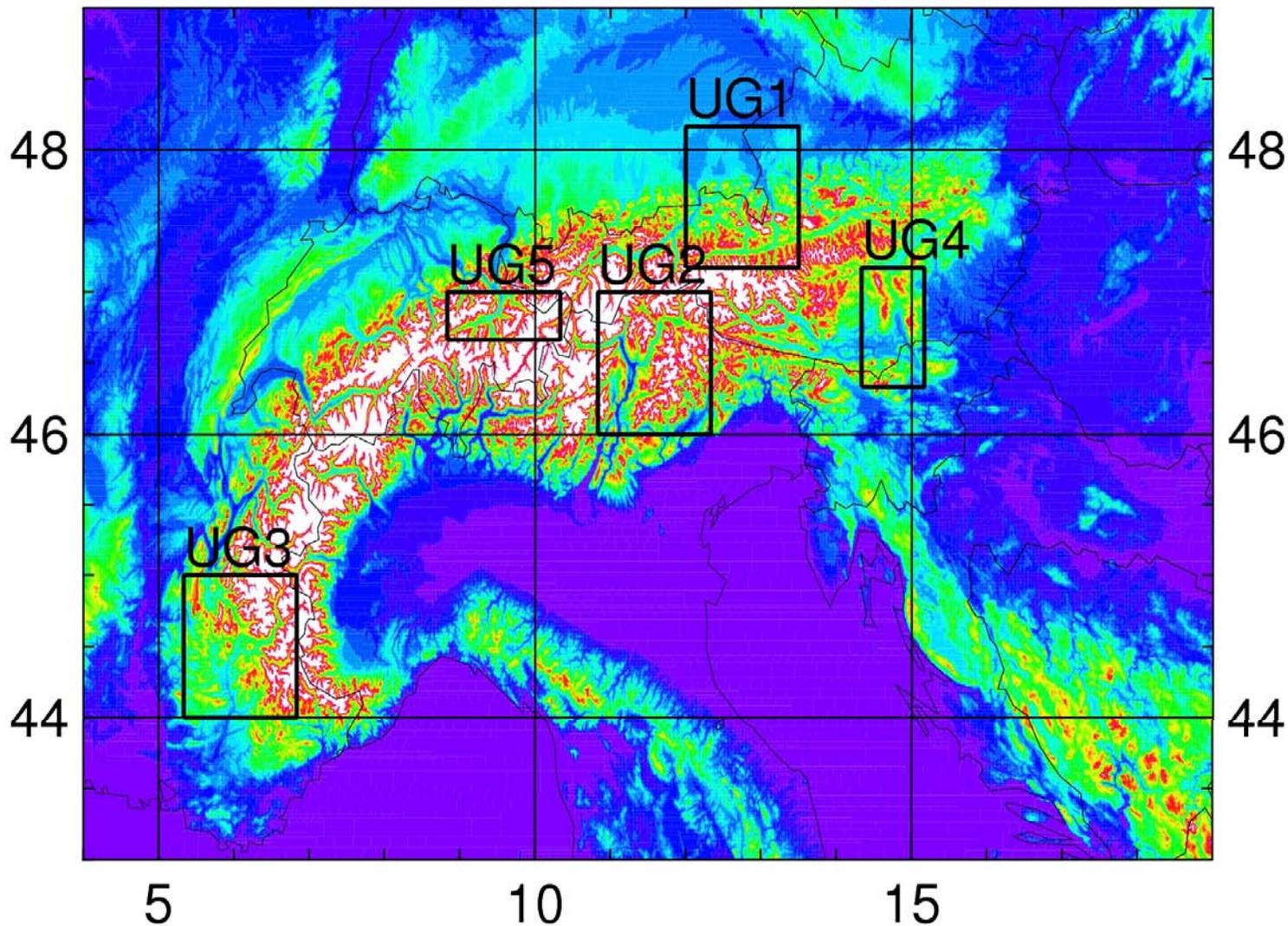
$$Q_{out} = (Q_{v,i-1} \cdot e^{-\Delta t/k_v} + Q_{v,i} \cdot (1 - e^{-\Delta t/k_v})) + (Q_{h,i-1} \cdot e^{-\Delta t/k_h} + Q_{h,i} \cdot (1 - e^{-\Delta t/k_h}))$$

$$\Delta x, \Delta y = 100\text{m}-1\text{km}, \Delta t = 1\text{h}-1\text{d}$$

How well do RCMs reproduce observed climate?

An example from the Alps

ClimChAlp Investigation areas



Greater Alpine Region (GAR)

Available High Resolution Data ($\Delta x < 20\text{km}$)

SRES	GCM	RCM	SDM	Ensembles
SRES B1	ECHAM5	CLM REMO		2?
	ECHAM5			1
	PCM		TYN	1
	Had3		TYN	1
	CSIRO2		TYN	1
	CGCM		TYN	1
SRES B2	Had3	RegCM		1
	PCM		TYN	1
	Had3		TYN	1
	CSIRO2		TYN	1
	CGCM		TYN	1
SRES A1B	ECHAM5	CLM REMO		2?
	ECHAM5			1
SRES A1FI	PCM		TYN	1
	Had3		TYN	1
	CSIRO2		TYN	1
	CGCM		TYN	1
SRES A1	PCM		TYN	1
	Had3		TYN	1
	CSIRO2		TYN	1
	CGCM		TYN	1
SRES A2	Had3	HIRHAM ReGCM REMO		1
	Had3			1
	EGHAM5			1
	PCM		TYN	1
	Had3		TYN	1
	CSIRO2		TYN	1
	CGCM		TYN	1

CLM - Germany (2001-2100)
hourly data, $\Delta x \approx 18\text{km}$

HIRHAM - DMI, Denmark(2070-2099)
daily data in the Internet (PRUDENCE)
 $\Delta x \approx 13\text{km}$

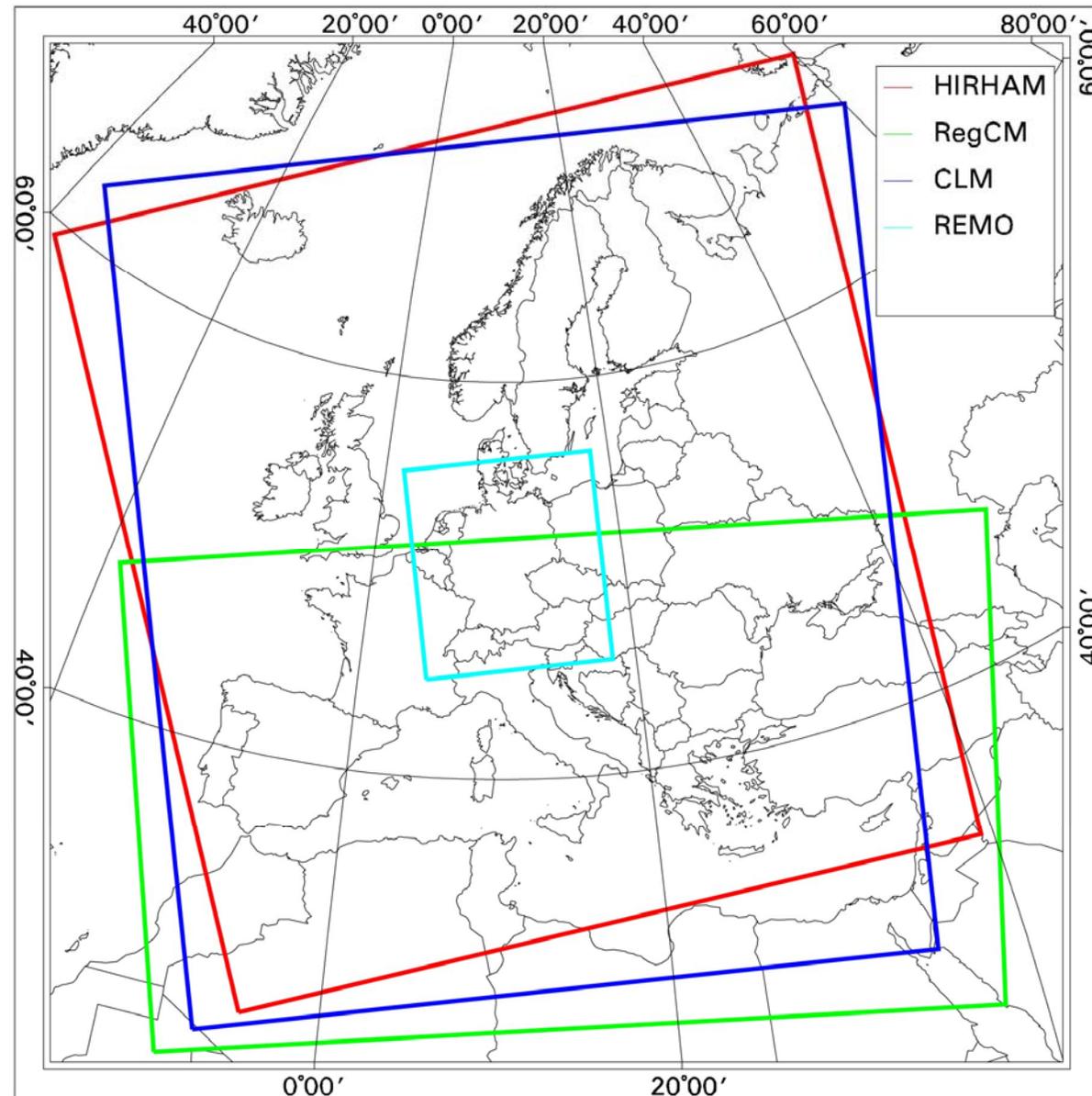
RegCM - ICTP, Italy (2070-2099)
daily data, $\Delta x \approx 20\text{km}$

REMO - MPI, Germany (2001-2100)
hourly data, $\Delta x \approx 10\text{km}$

TYN – Tyndall Centre, UK, statistical
downscaling

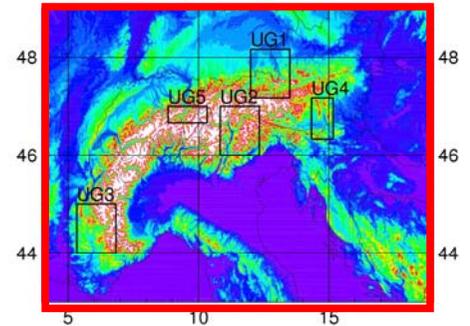
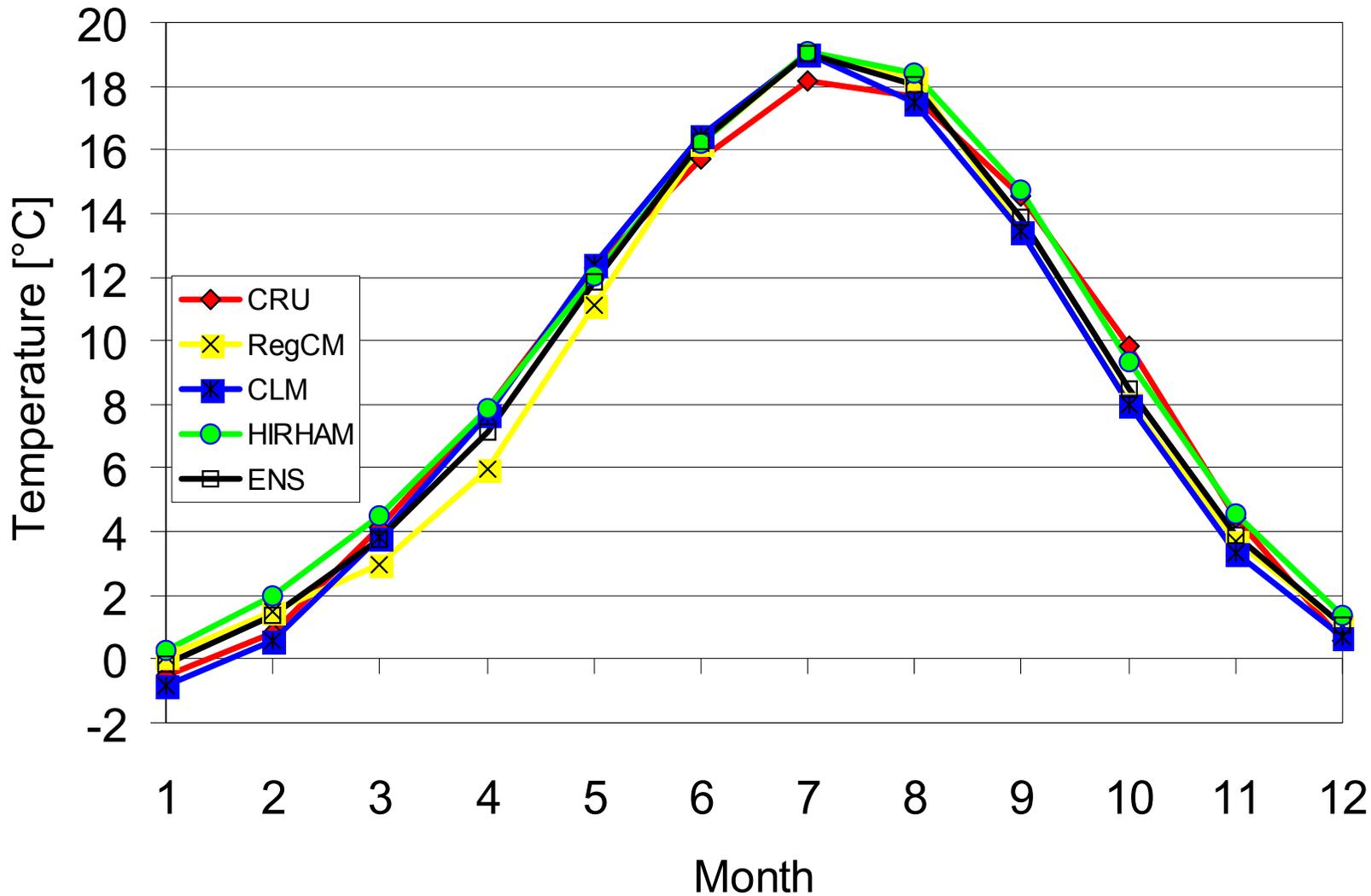
Different GCMs & different scenarios !

Extent of High Resolution Data



Performance Present Climate: **GAR (Greater Alpine Region)**

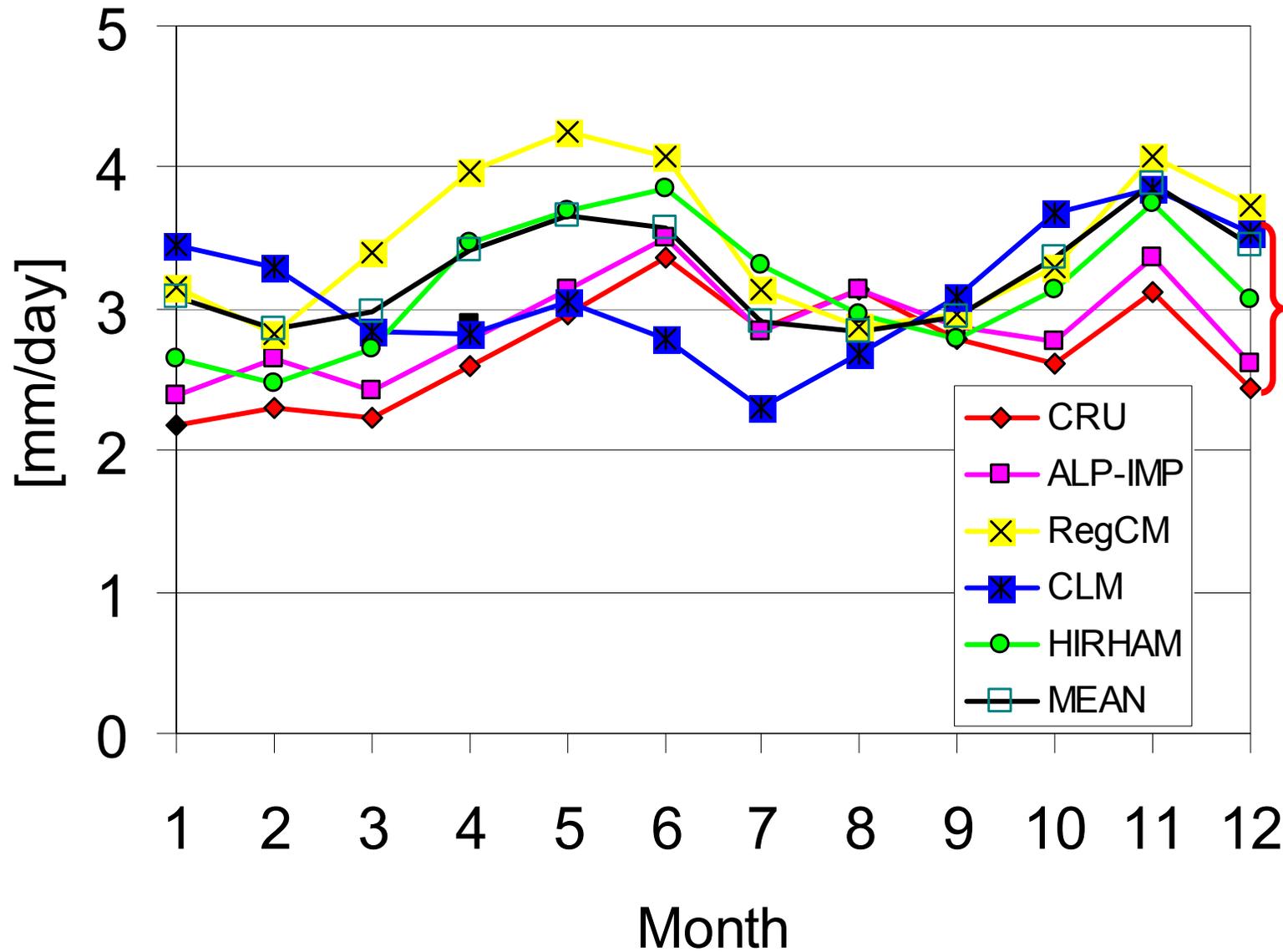
GAR



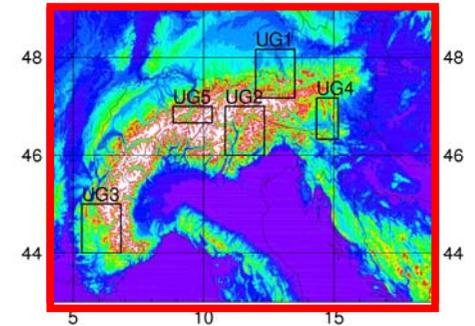
Mean monthly temperature bias +/- 1°C

1961-90

Performance Present Climate: **GAR**

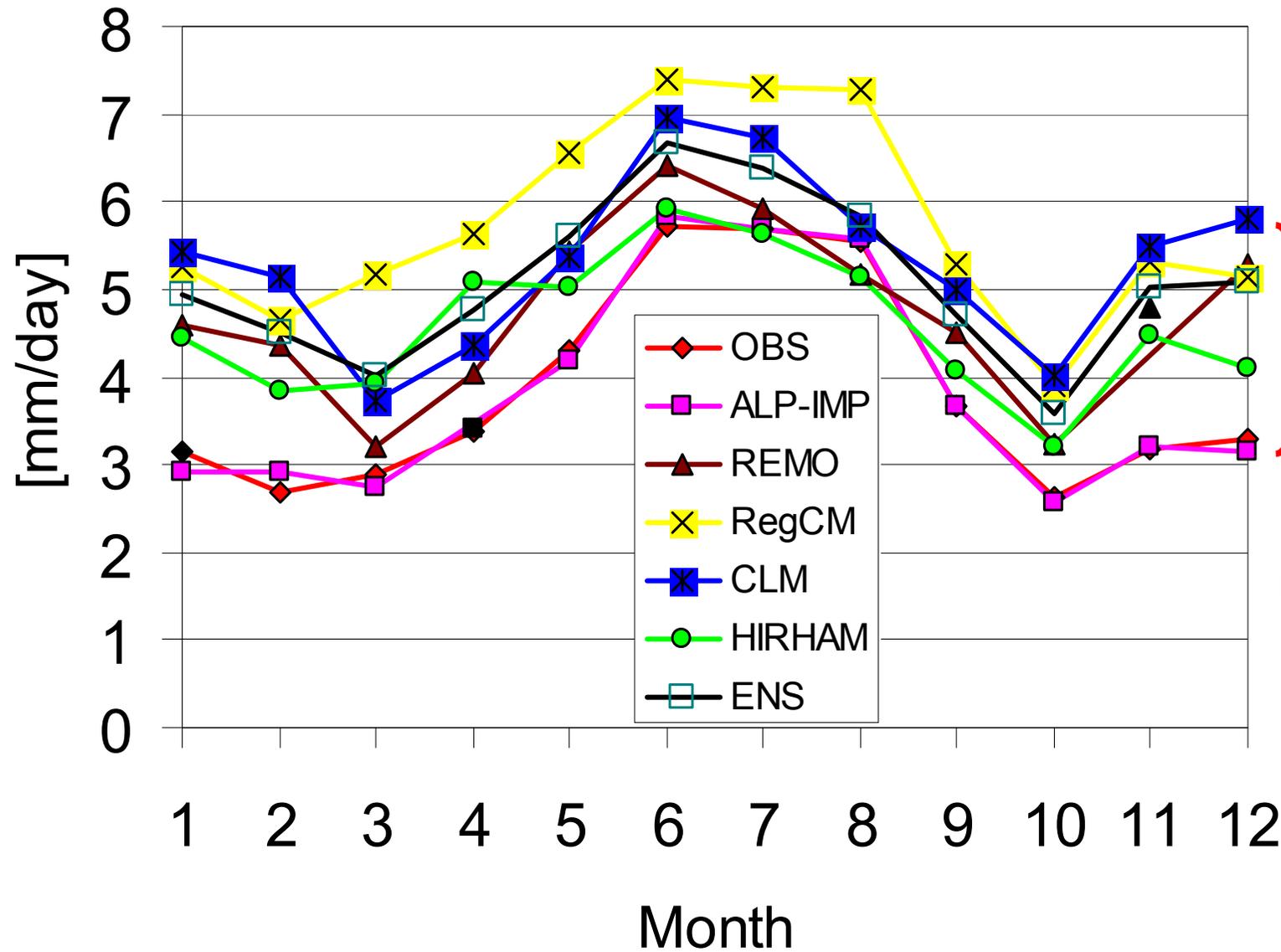


GAR

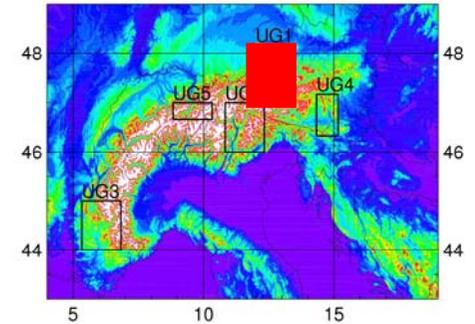


Precipitation bias up to 1.5 mm/day

1961-90



UG1

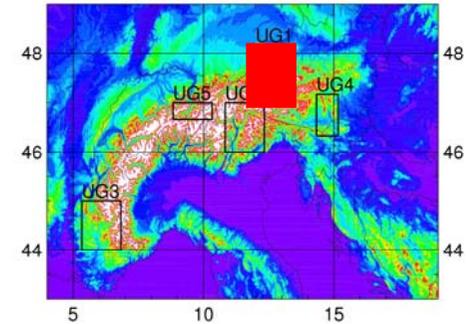


Precipitation bias up to +2.5 mm/day!

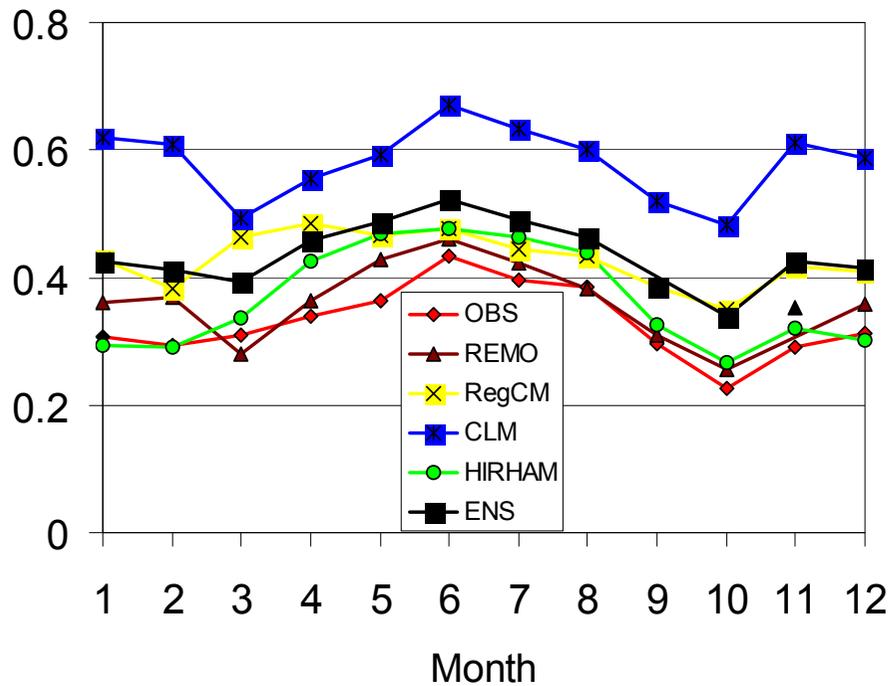
1961-90

FRE: Frequency (fraction) of days with

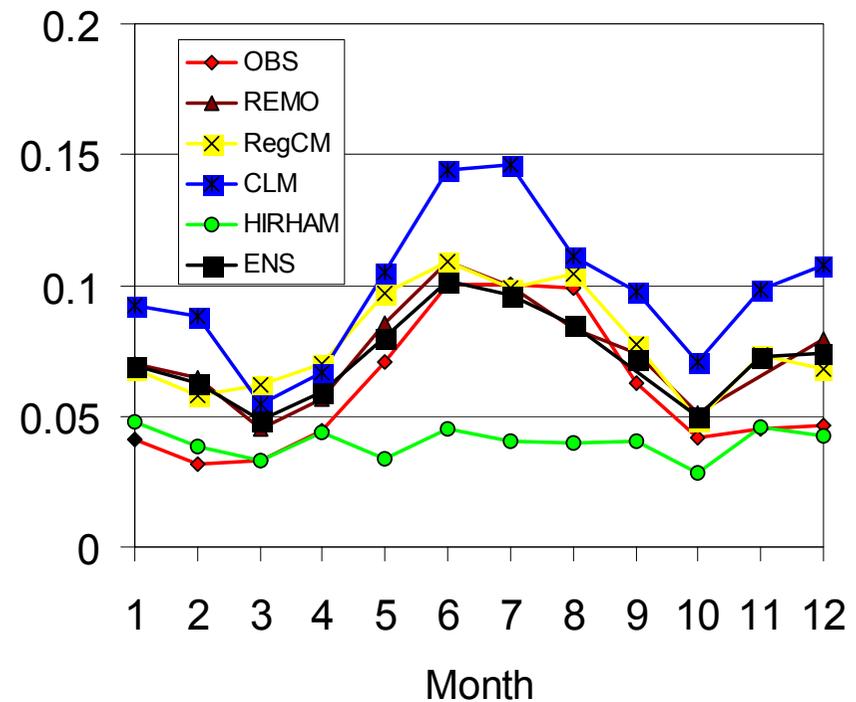
UG1



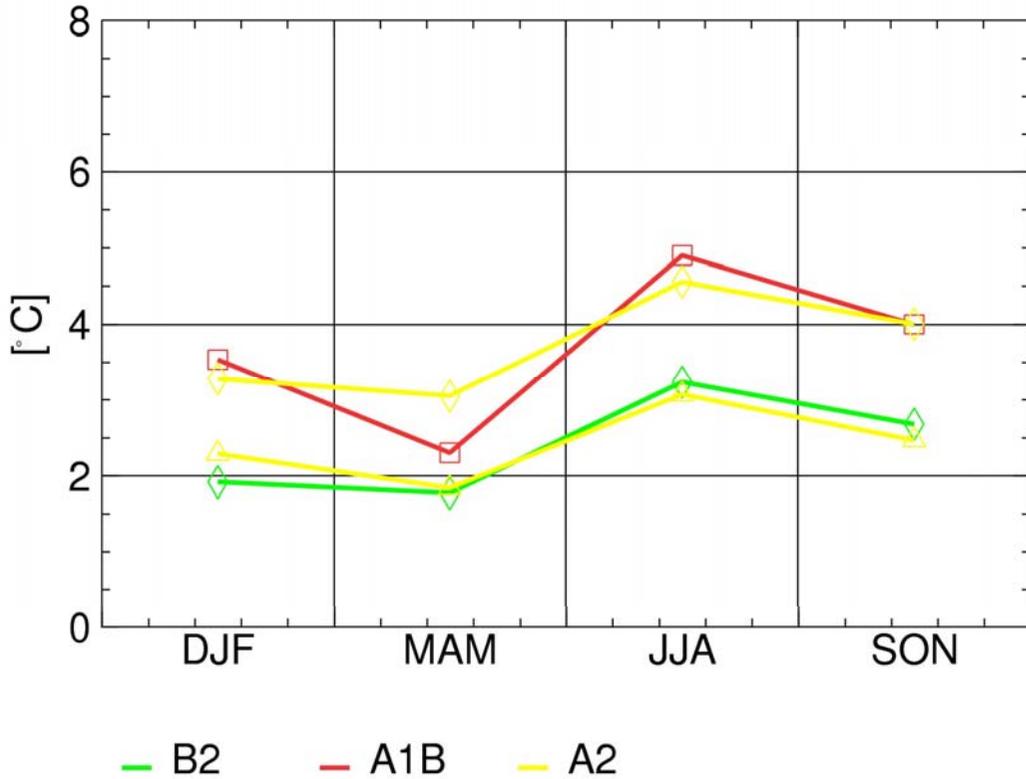
FRE 1: $P > 1\text{mm}$



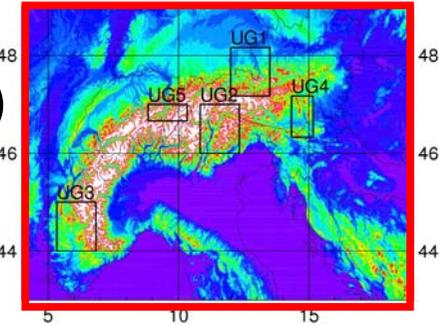
FRE 15: $P > 15\text{mm}$



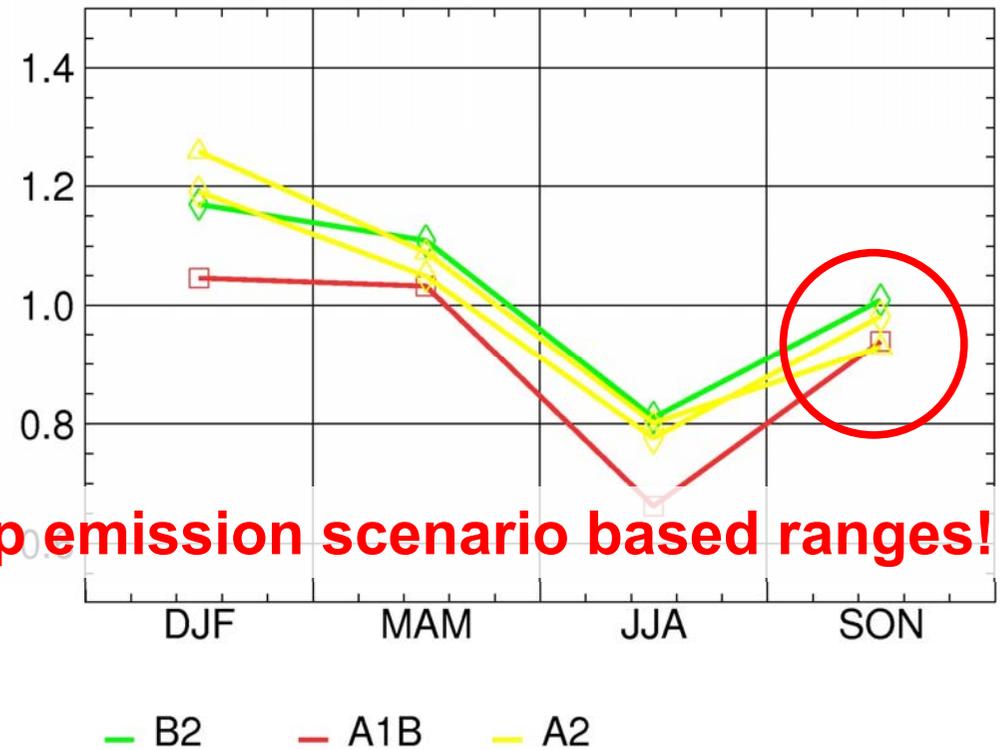
Temperature GAR Ts-Tc



Δ: HIRHAM (HadCM3)
 □: CLM (ECHAM5)
 ◇: RegCM (HadCM3)



Precipitation GAR Ps/Pc



RCM ranges overlap emission scenario based ranges!

c - control run (1961 – 1990)
 s - scenario run (2071 – 2100)

From RCM from hydrological impact analysis...

An example from the Alps

Regional Climate Modeling & Hydrological Impact Analysis

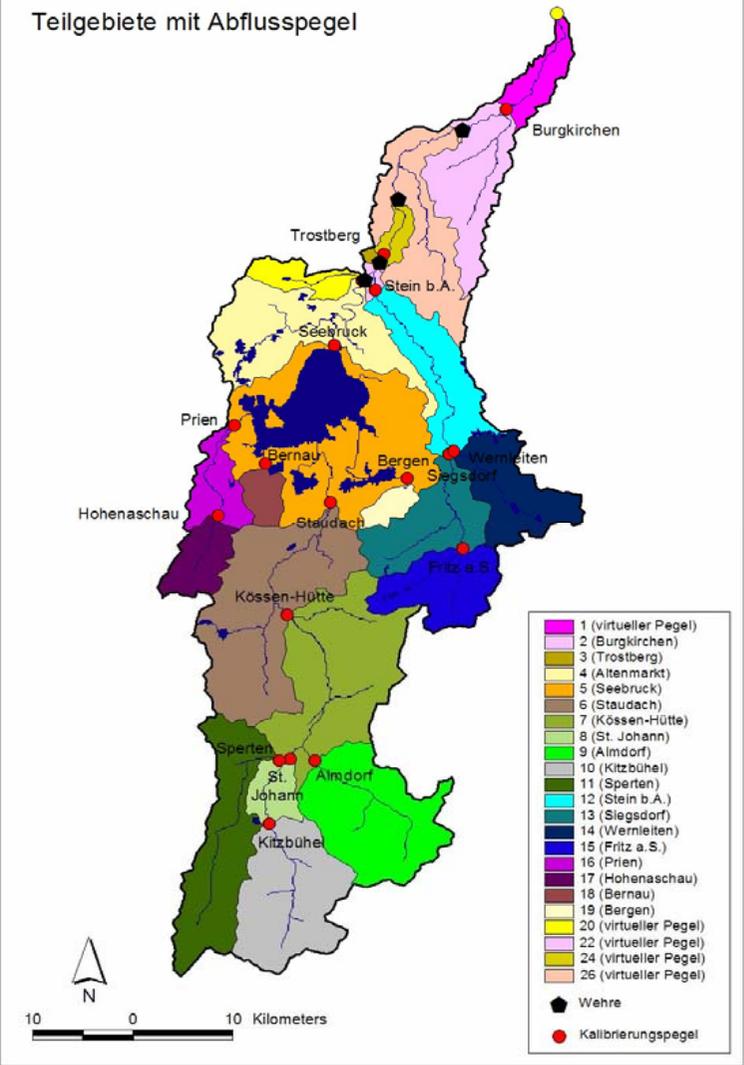


Alz Catchment

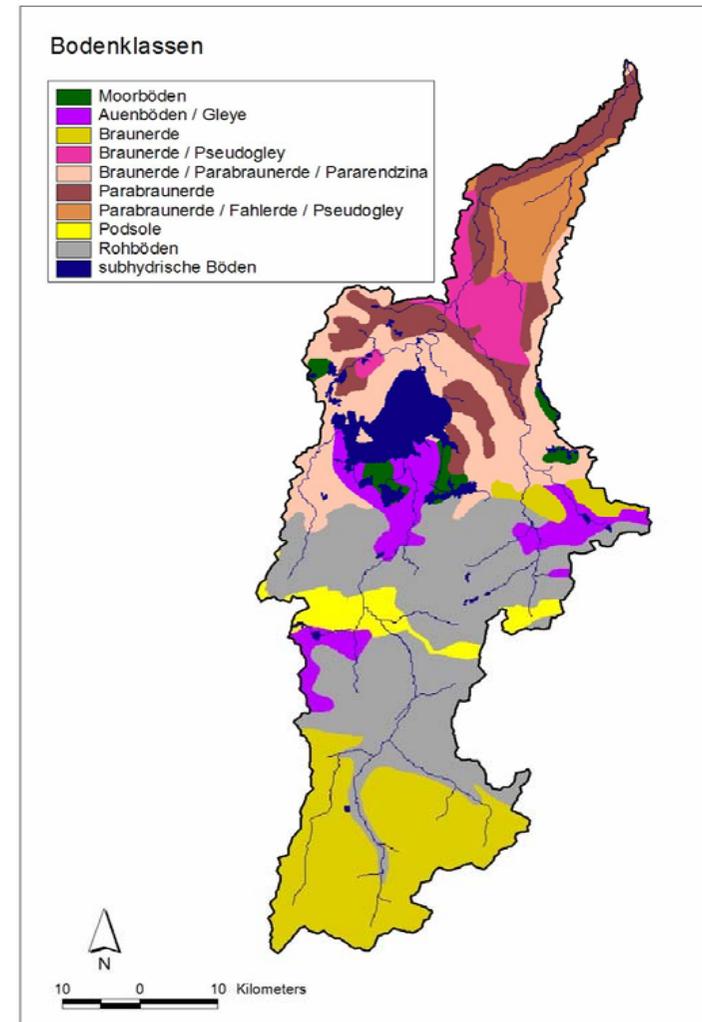
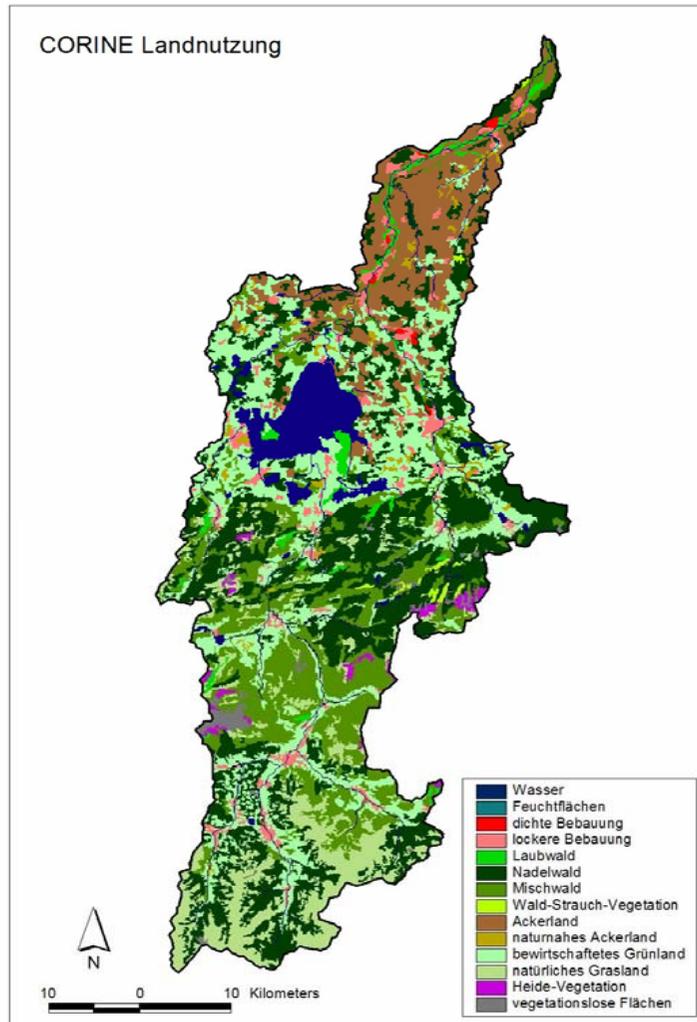
- Area: 2256 km²
- Complex Topography
- Elevation: 359-2328 N.N. **Chiemsee**
- Mean precipitation: 1450 mm/a (1900mm/a mountains, 850 mm/a low lands)
- Days with snow cover: $\approx 104/a$ (Kitzbühl)
- Temperature gradient: $\approx 0.6 \text{ }^\circ\text{C}/100\text{m}$
- Distributed hydrological modeling $\Delta x \Delta y = 500\text{m}$

Lake Chiemsee & Chiemgau Alps

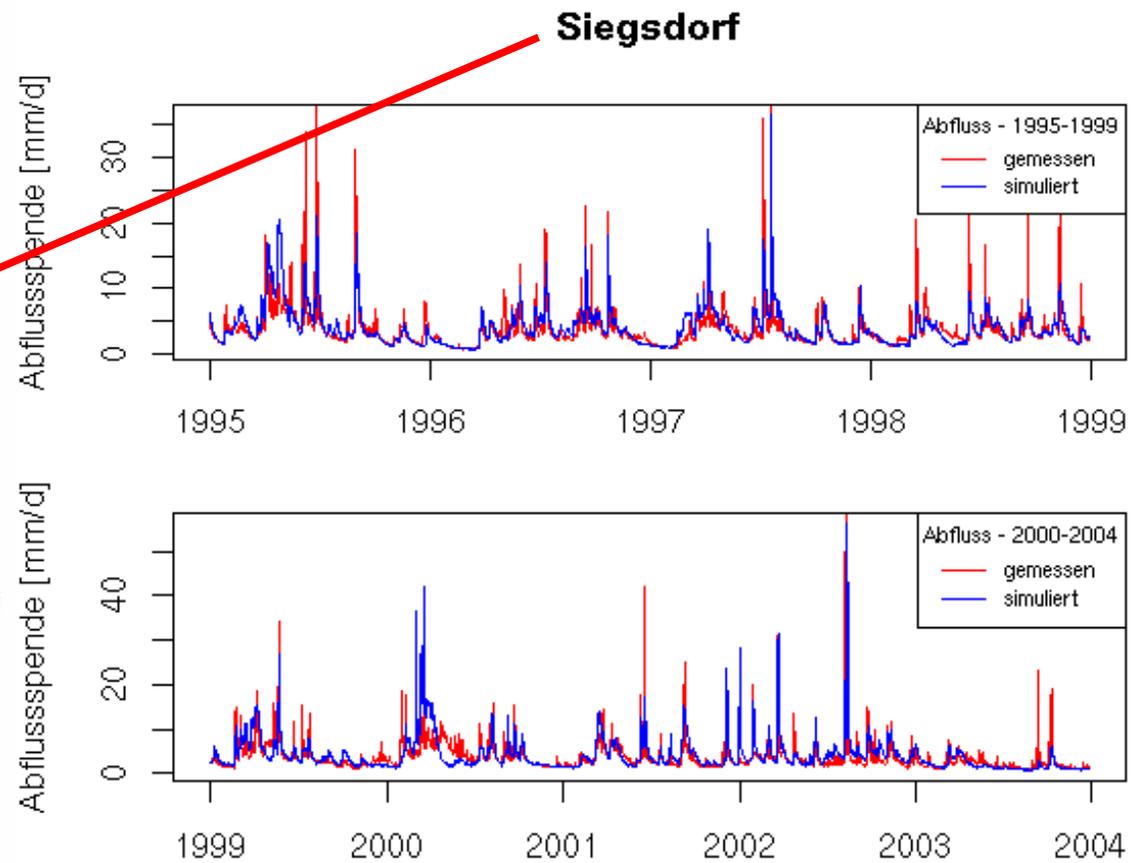
Teilgebiete mit Abflusspegel



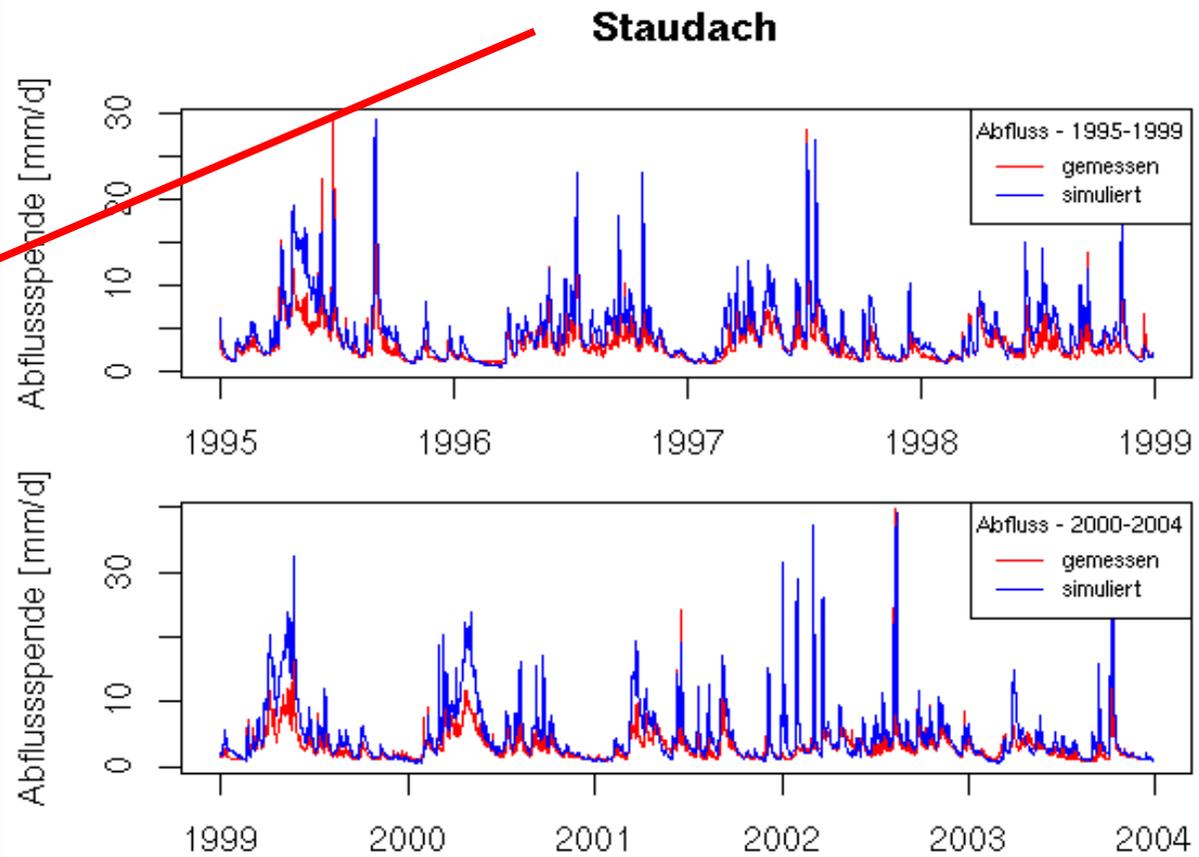
Time Invariant Data for Hydrological Model



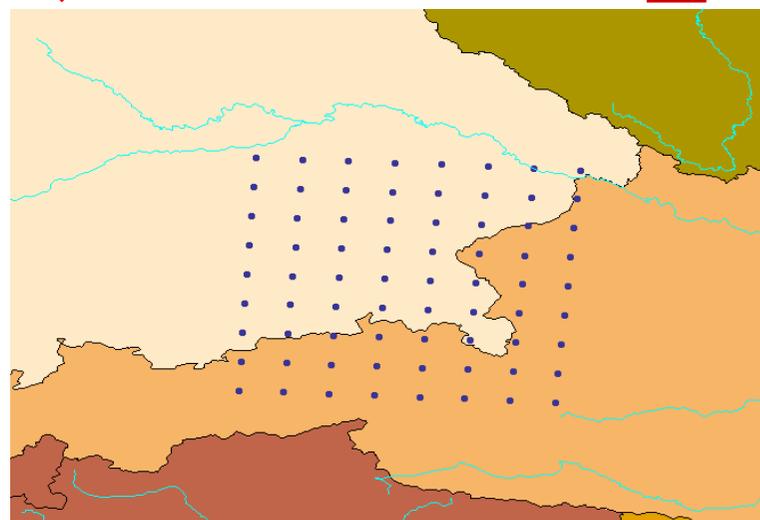
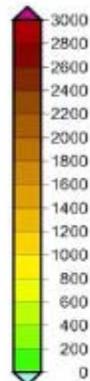
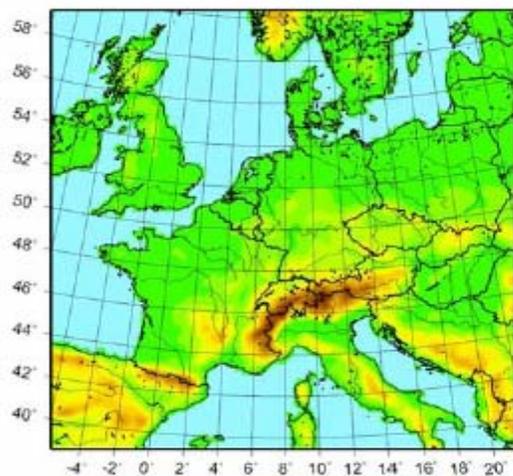
Water balance simulation with WaSiM



Water balance simulation with WaSiM

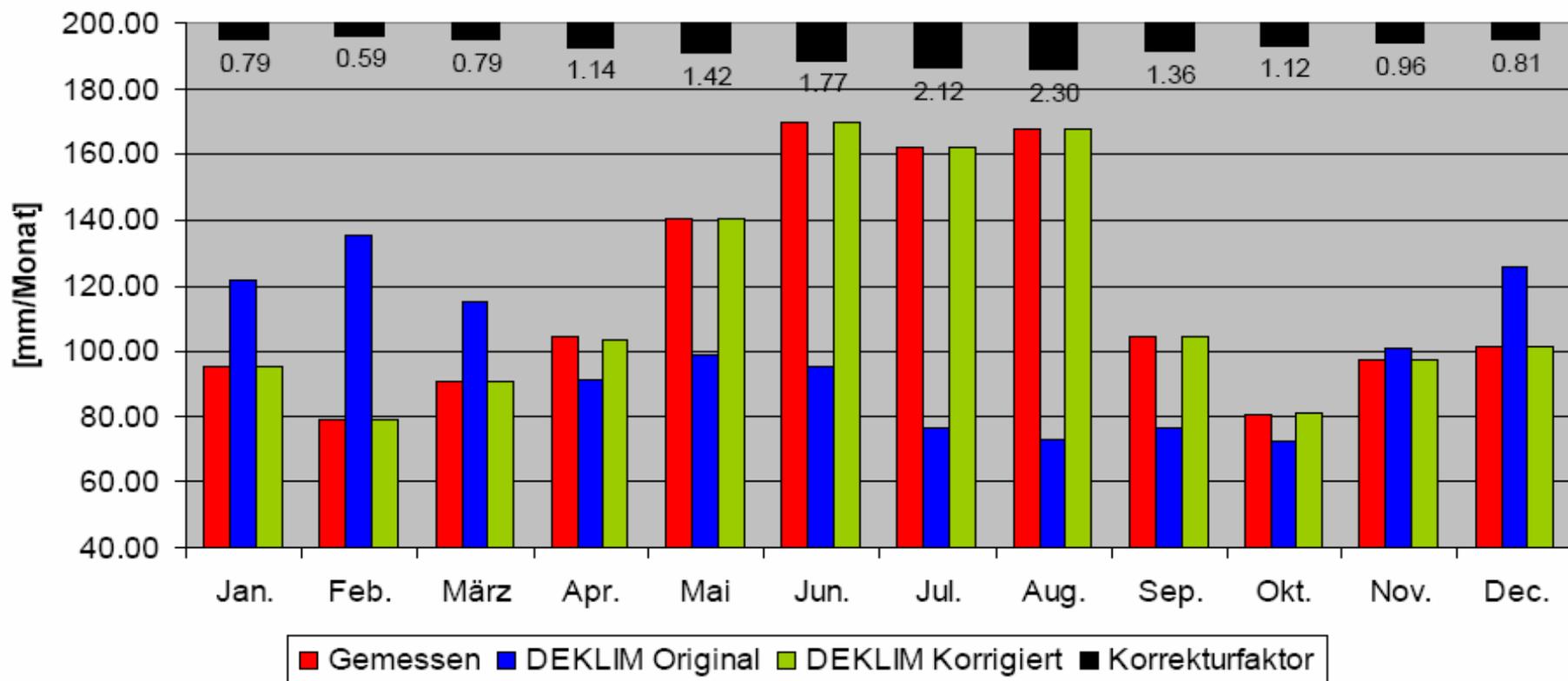


Model Chain ECHAM4(B2) – MM5@18km – WaSiM@500m



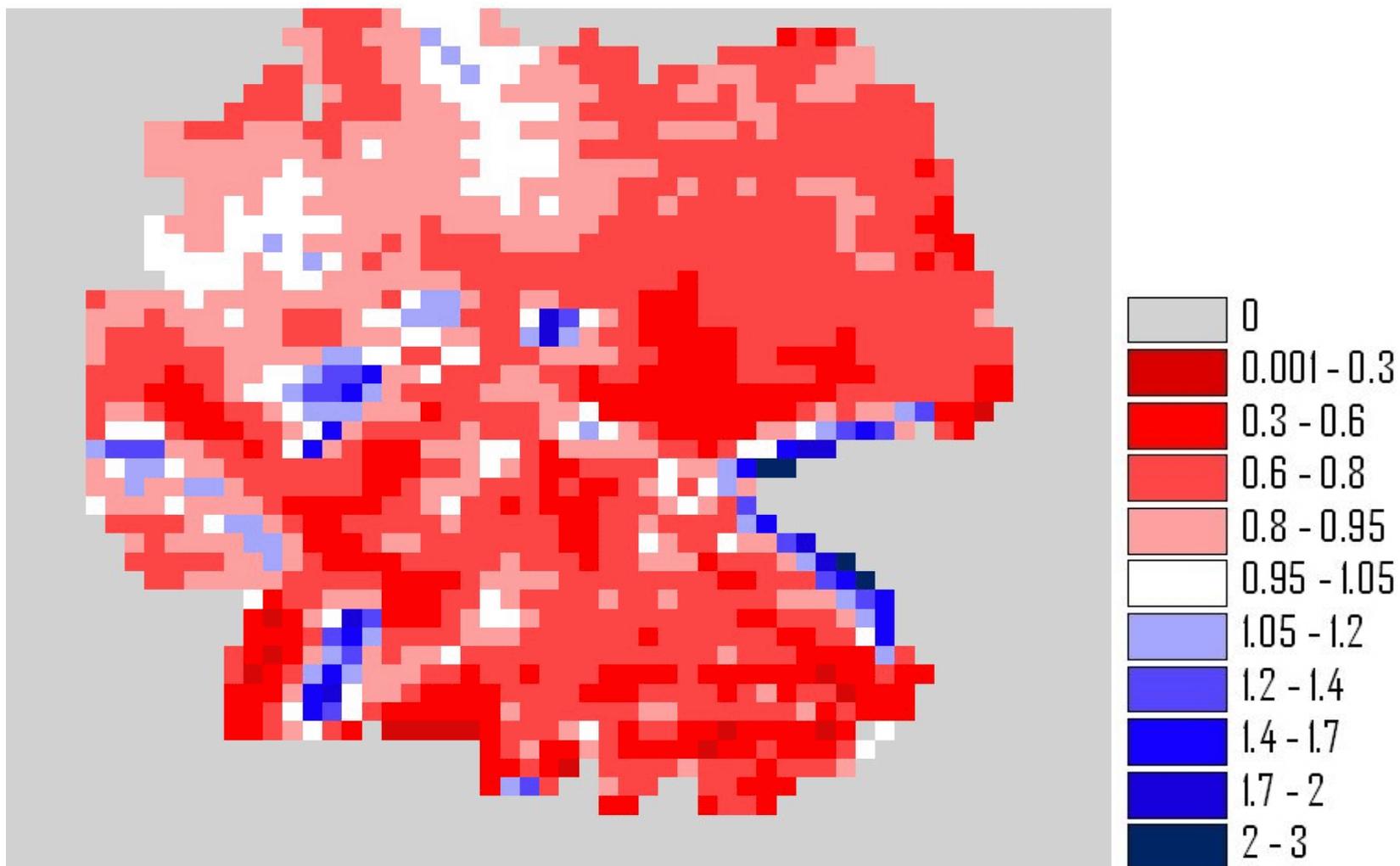
Position of Grid Points

Method & Bias-Correction after
Kunstmann et al., 2004 (HESS, Ammer EZG),
Kunstmann et al., 2007 (PCE, Volta EZG),
Kunstmann et al., 2007 (IAHS, oberes Jordan EZG)



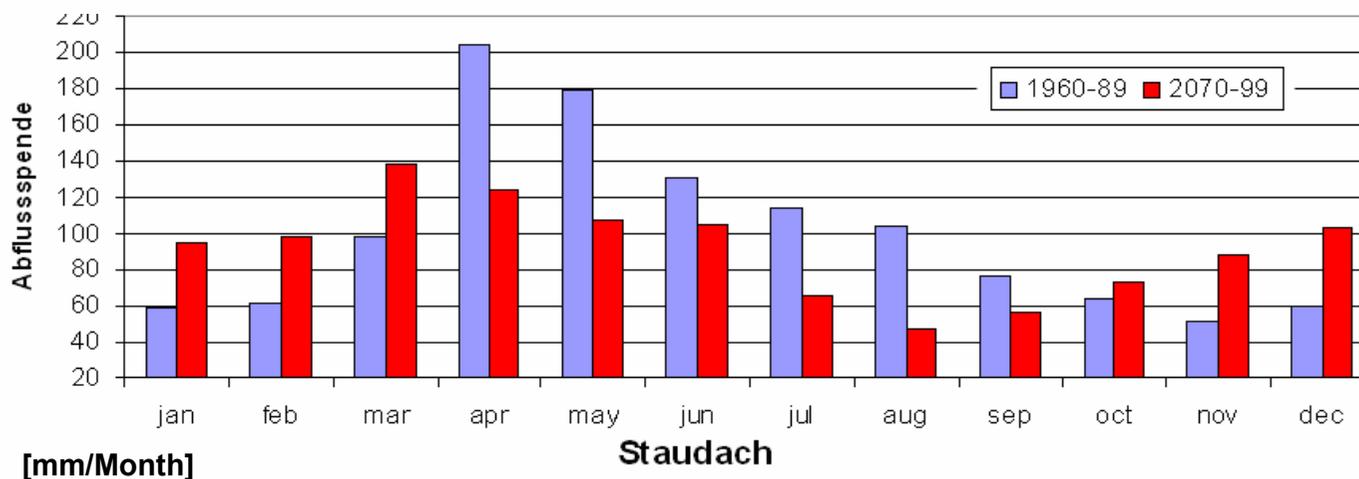
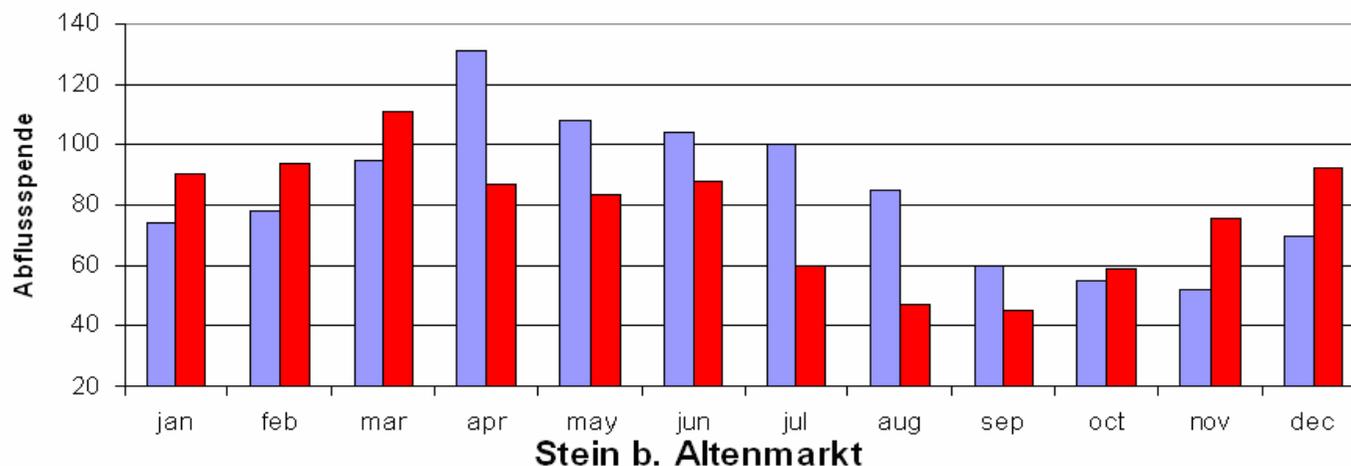
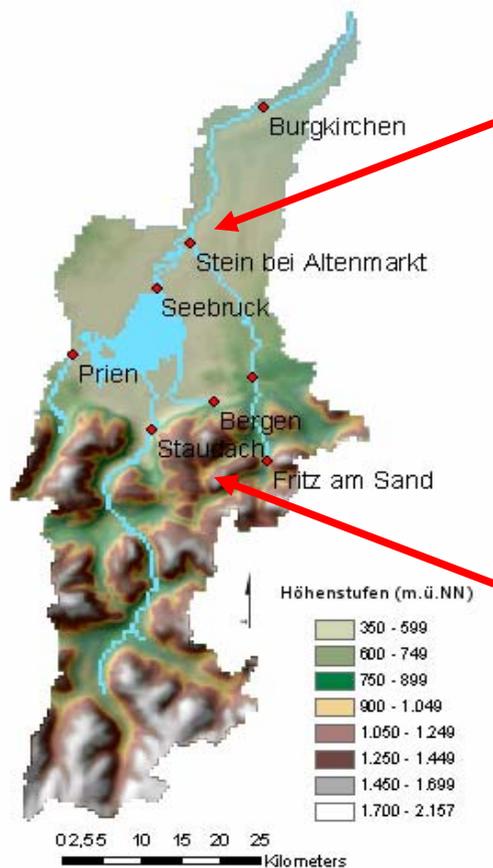
Alz catchment:
 Monthly bias-correction factors for ECHAM4-MM5@19km: CTRL vs. climatology

Area:
Germany+
Austrian part of
Alz catchment



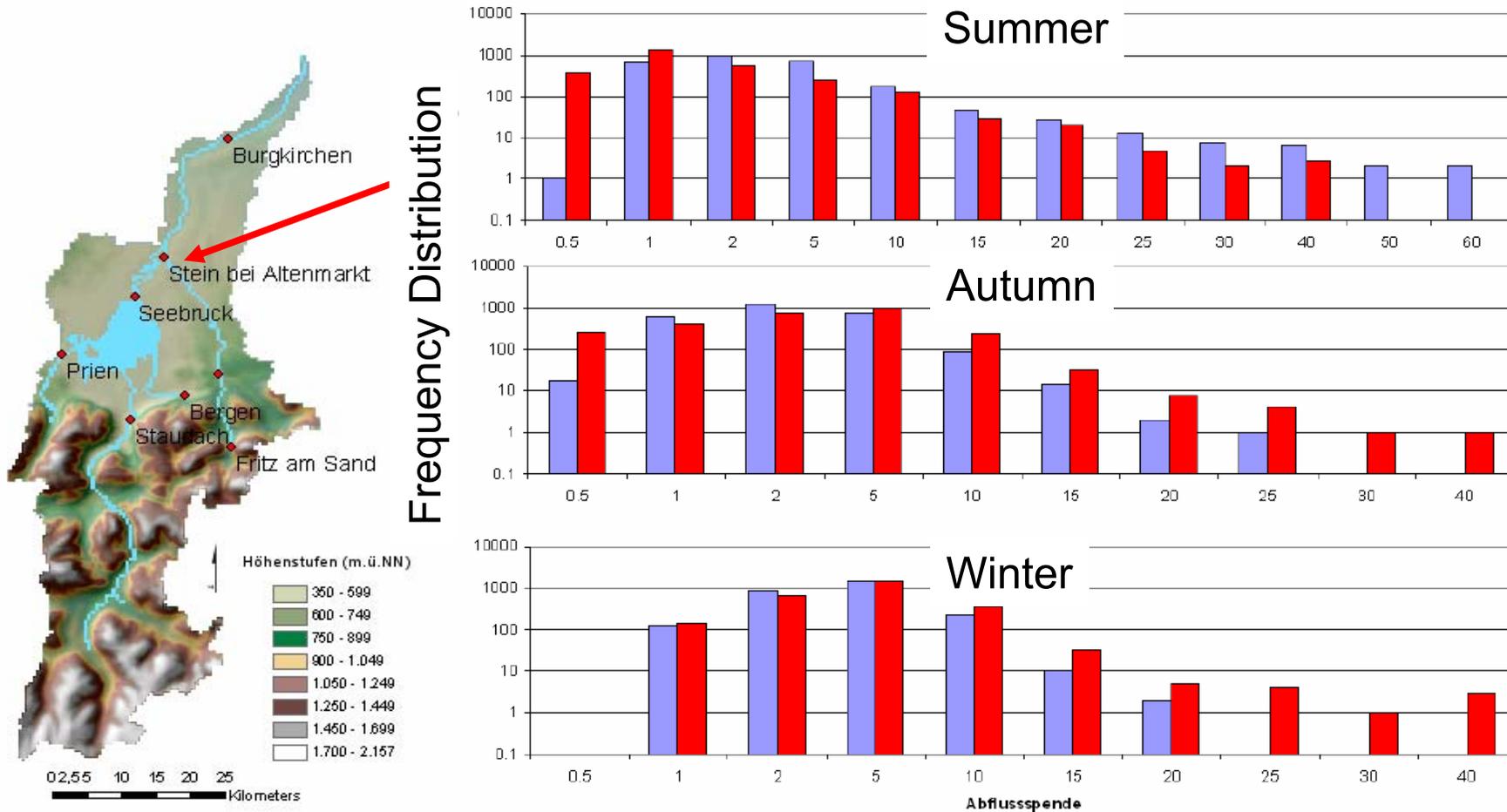
Monthly bias-correction factors for ECHAM5-CLM@18km: CTRL vs. climatology

Joint Model System ECHAM4(B2) – MM5@18km – WaSiM@500m



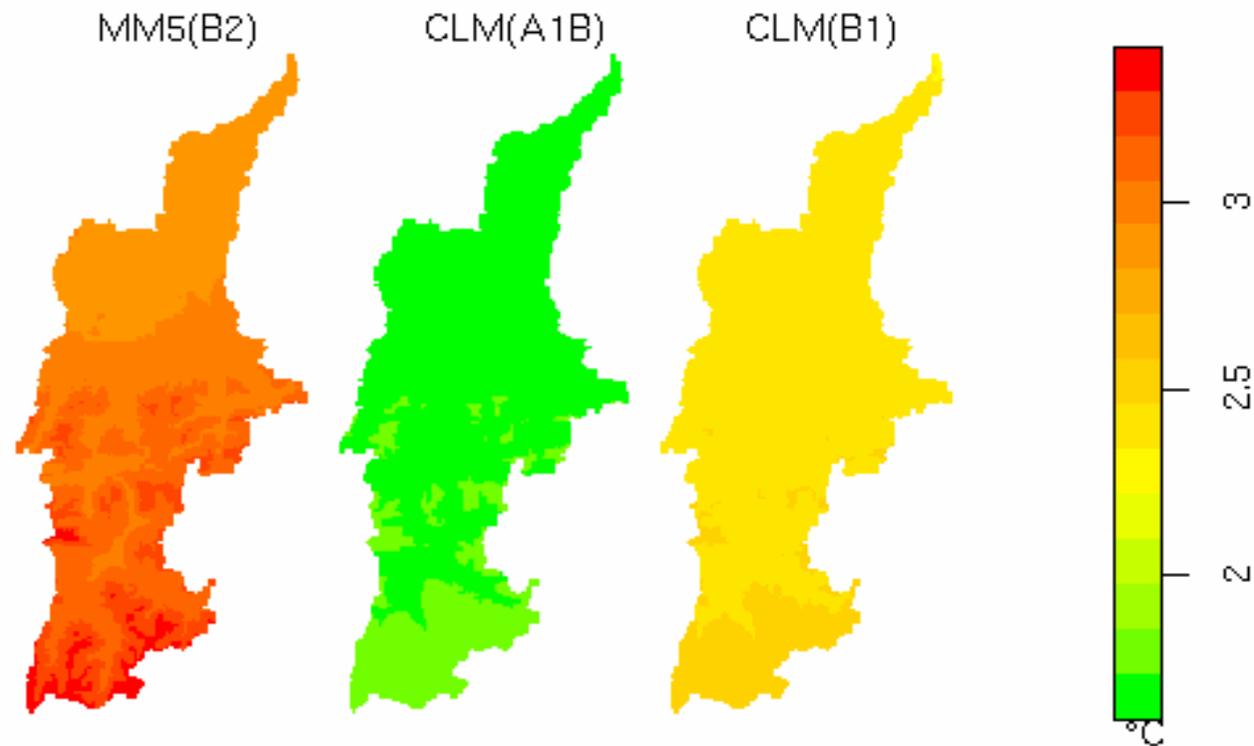
Decrease of summer, increase of winter runoff

Joint Model System ECHAM4(B2) – MM5@18km – WaSiM@500m



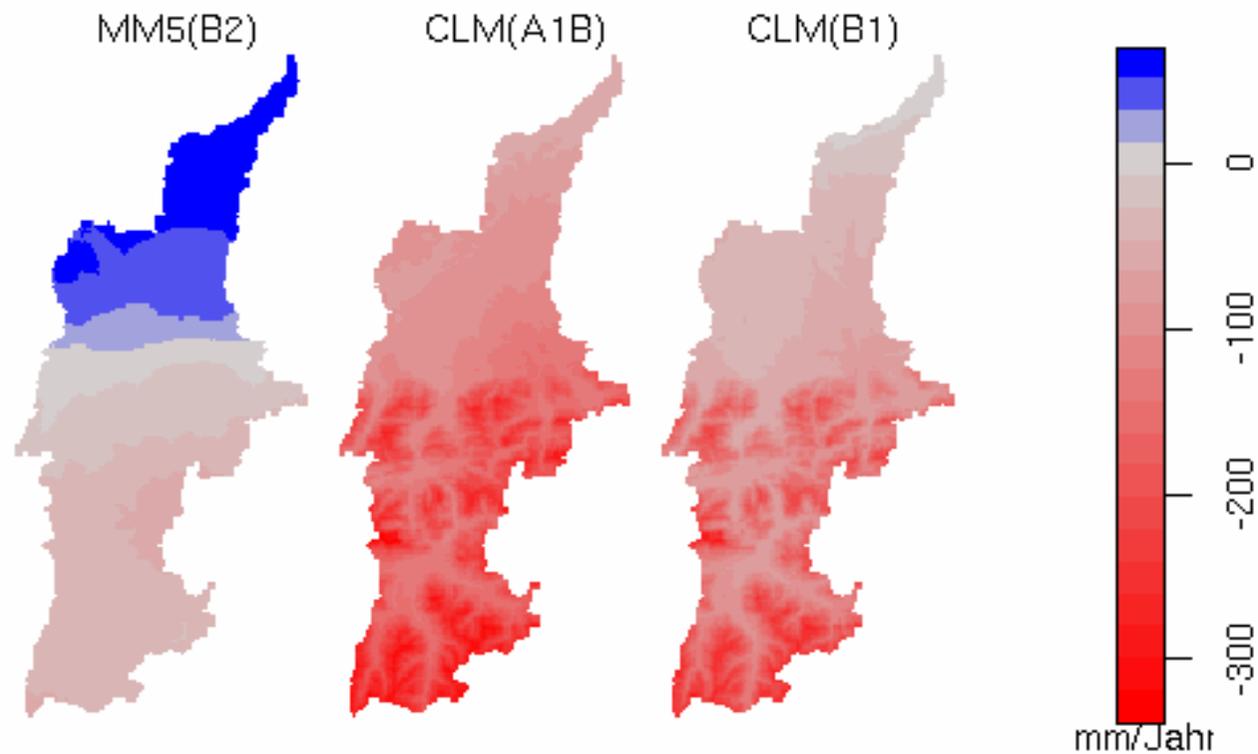
Changes in frequencies: Increase of both low flow & high flow situations!

ECHAM4-MM5@18km vs. ECHAM5-CLM@19km – WaSiM@500m



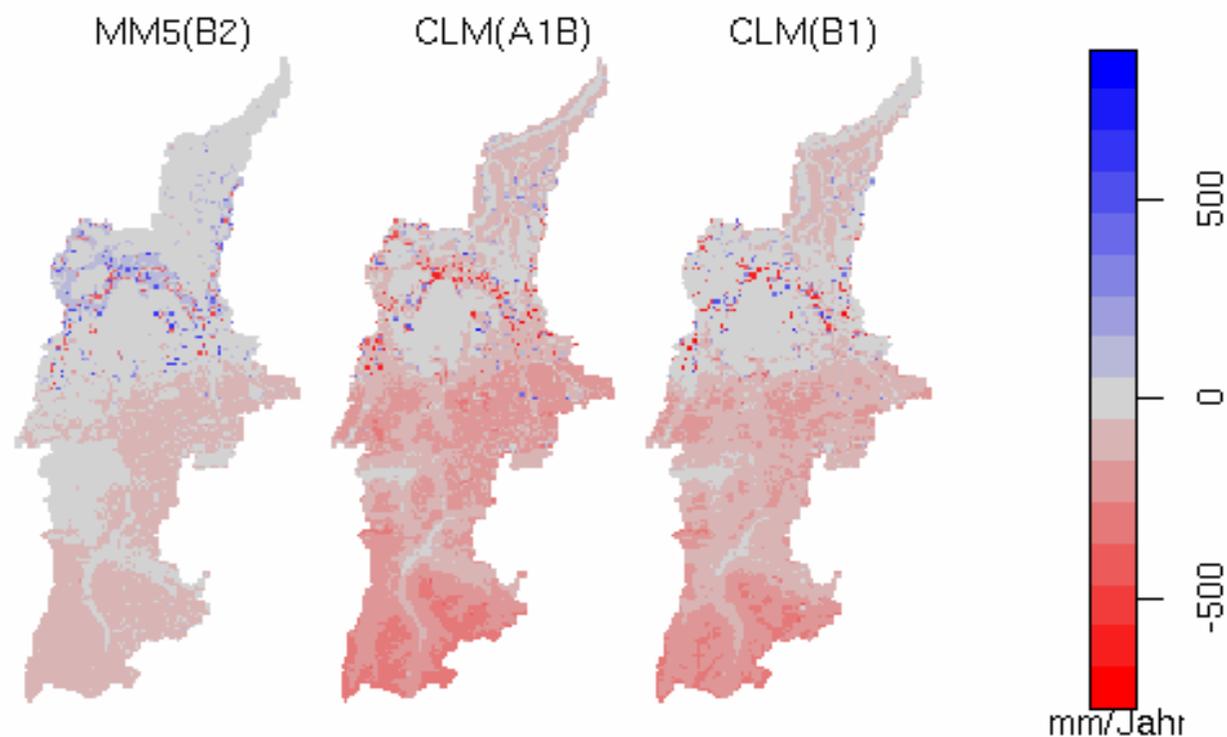
Temperature Change 2070-99 minus 1960-89

ECHAM4-MM5@18km vs. ECHAM5-CLM@19km – WaSiM@500m



Precipitation Change 2070-99 minus 1960-89

ECHAM4-MM5@18km vs. ECHAM5-CLM@19km – WaSiM@500m



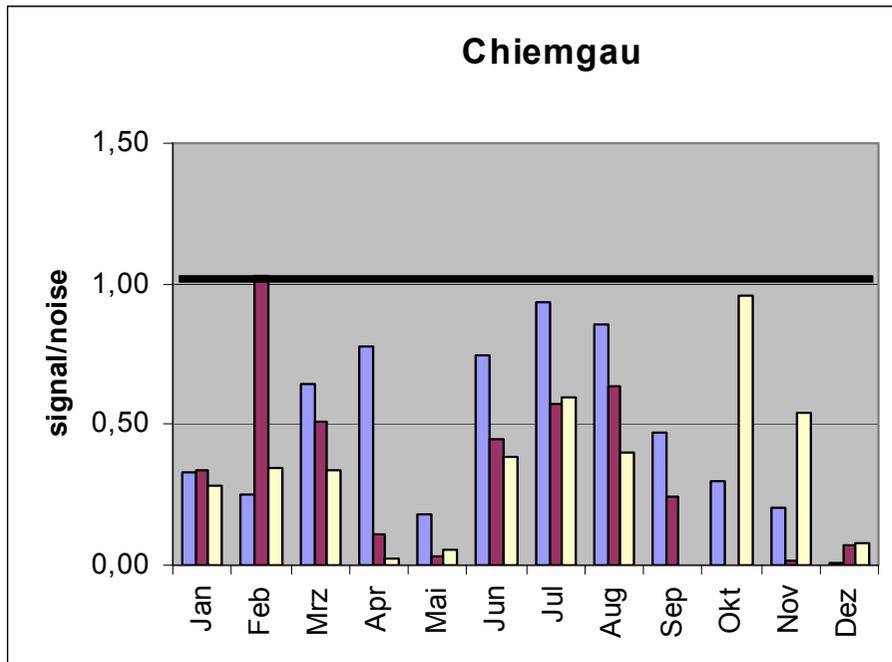
Total Runoff Change 2070-99 minus 1960-89

Expected Climate Change: Trend vs. Natural Variability

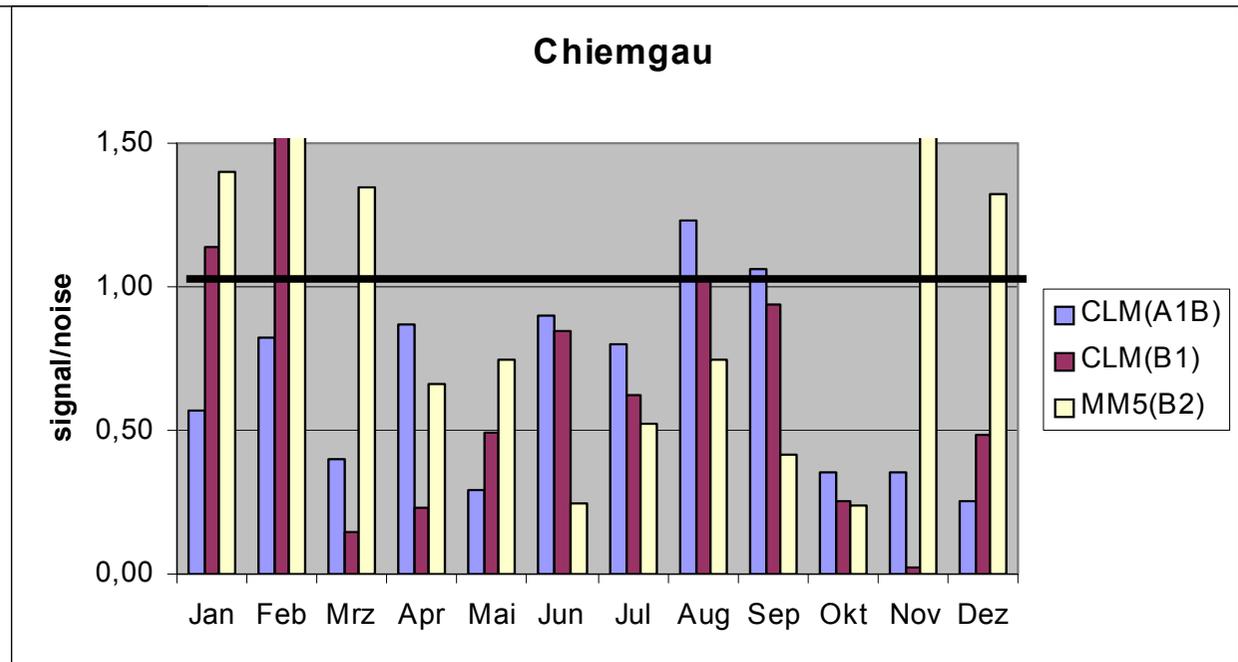
Model Chain ECHAM4/5 - MM5@18km/CLM@19km - WaSiM@500m

Signal to Noise ratio:

$$SN = \frac{|\bar{X}_{fut} - \bar{X}_{pres}|}{\sigma} > 1?$$



SN for Precipitation



SN for Discharge

Signal climate change mostly within inter-annual variability!

Summary: Current Capabilities & Limitations

- Regional climate models: spatial resolutions down to 5-10km for climate runs
- Non-hydrostatic models for $\Delta x < 10$ km
- **No single model can be identified as best:**
 - performance depends on selected variable and area
 - ⇒ Hydrological impact studies by ensemble data set
- There is further a clear need for high resolution RCM data
- More detailed climatology needed (daily station-statistics)
- **Significant biases in precipitation detected!**
 - ⇒ Biases in precipitation usually require statistical correction techniques
- **RCM ranges overlap emission scenario based ranges!**

... How about the Philippines?



Thank you for your attention