

Coupled Simulation of Groundwater-Soil-Atmosphere Interaction for the TERENO Pre-Alpine Region (Topic 4)

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Reklim Workshop, Bad Honnef, 16.-18. April 2013



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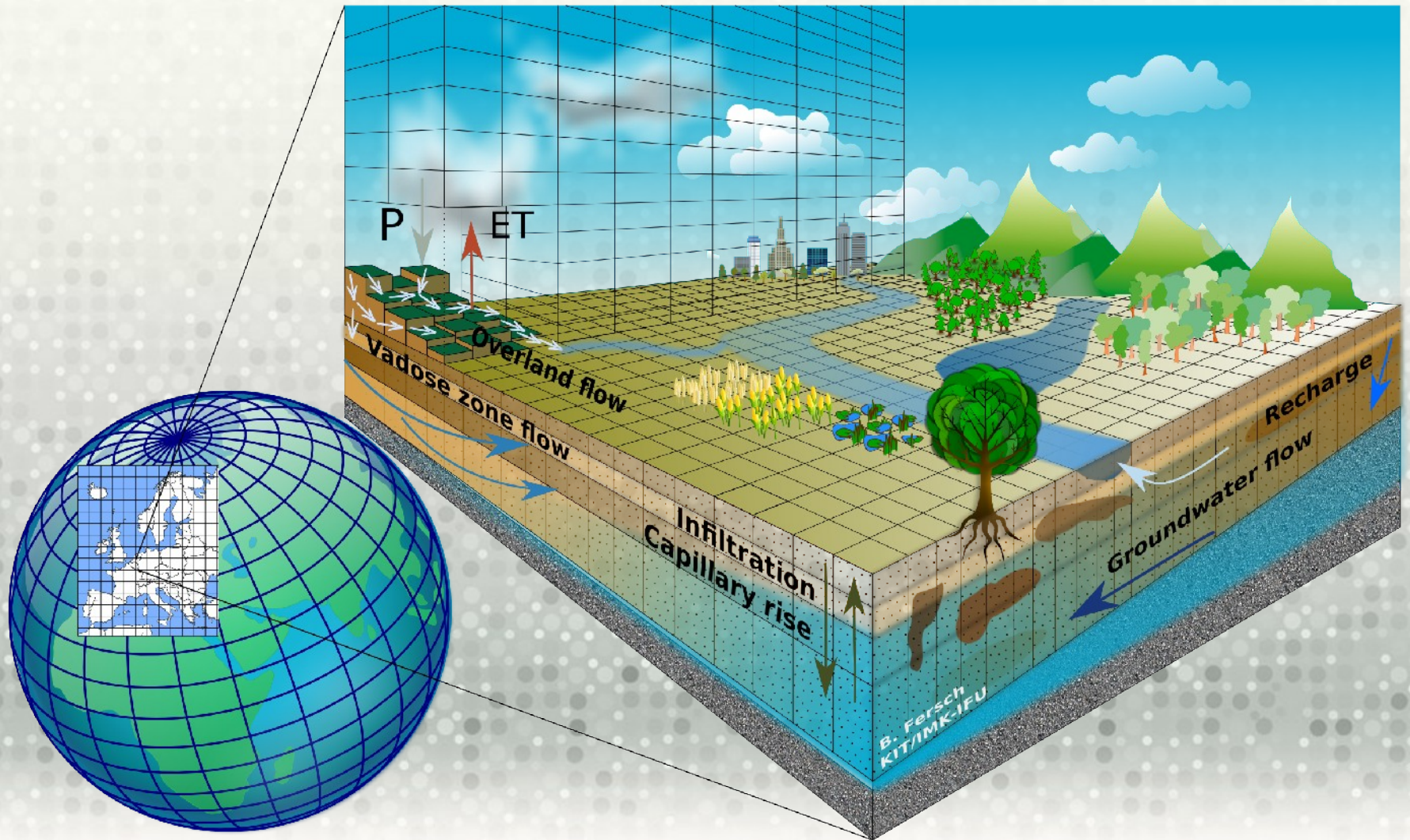


Motivation

- Closed expression of the regional water balance in dynamic model system
 - Regional atmospheric modeling studies focus usually on the prediction of temperature and precipitation
 - Coupling to hydrological models is typically realized in a one way mode and by using bias correction
 - Such approaches violate the closure of the water balance of the full system from the subsurface via the surface to the atmosphere

- Towards fully coupled cross-compartment hydrological-atmospheric model systems
 - Consider effects of lateral redistribution of surface water
 - Reinfiltration
 - Runoff routing
 - Account for subsurface processes of the vadose and the phreatic zone
 - Percolation to the groundwater
 - Capillary rise to root layers
 - Horizontal transport
 - Provide feedback capability from the subsurface via the surface to the atmospheric model and vice versa

Coupled Atmosphere-Hydrology Model System



Coupled Atmosphere-Hydrology Model Systems

Headwater Scale
<10,000km²

WRF-Hydro
(NDHMS)

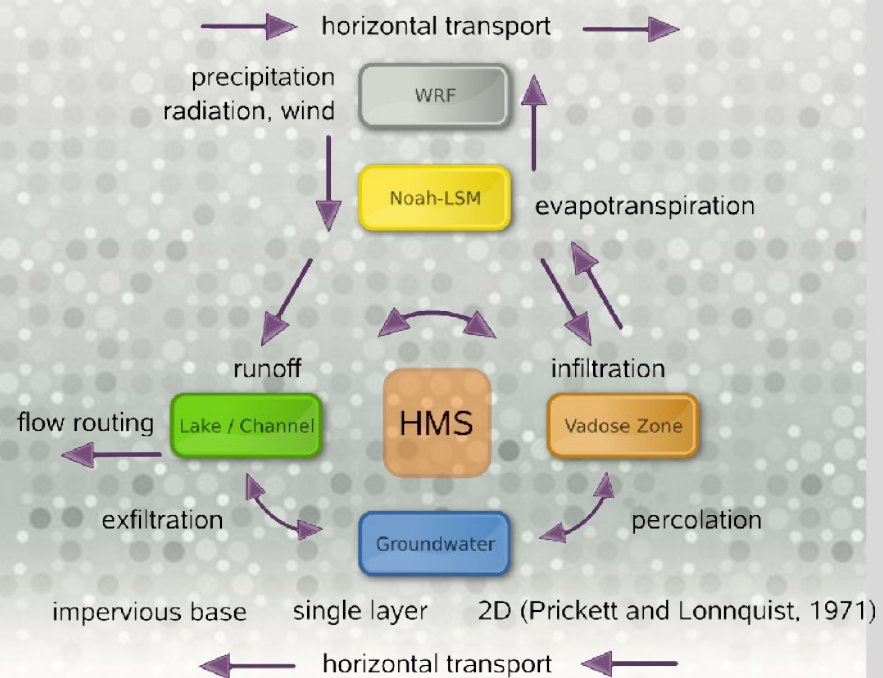
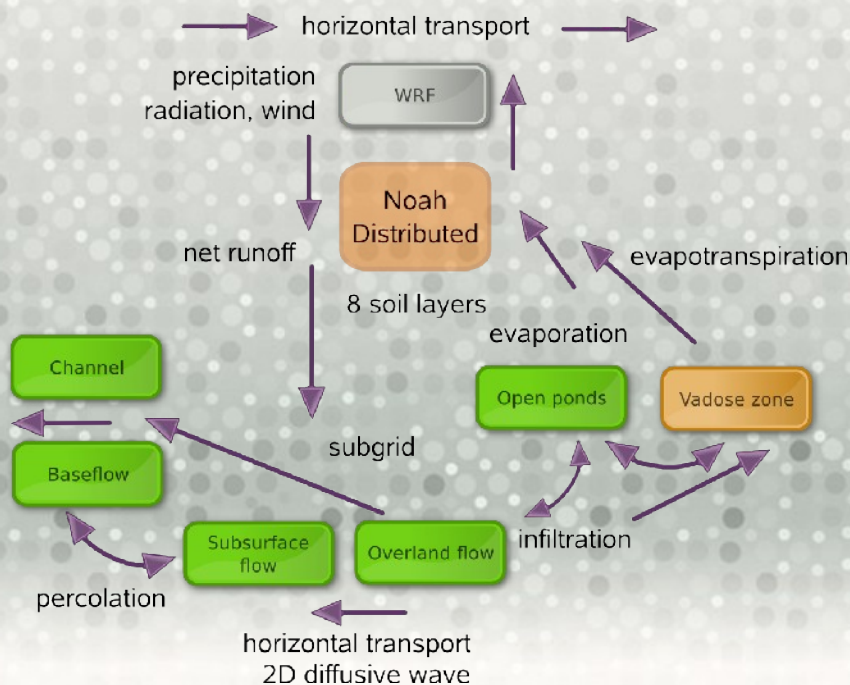
Mesh ~1km
Subgrid ~100m
Weeks to Months

River Scale
>10,000 km²

HMS/
Noah-LSM/
WRF

Mesh ~10km
No Subgrid
Years to Decades

vs.



Coupled Atmosphere-Hydrology Model Systems

Headwater Scale
<10,000km²

WRF-Hydro
(NDHMS)

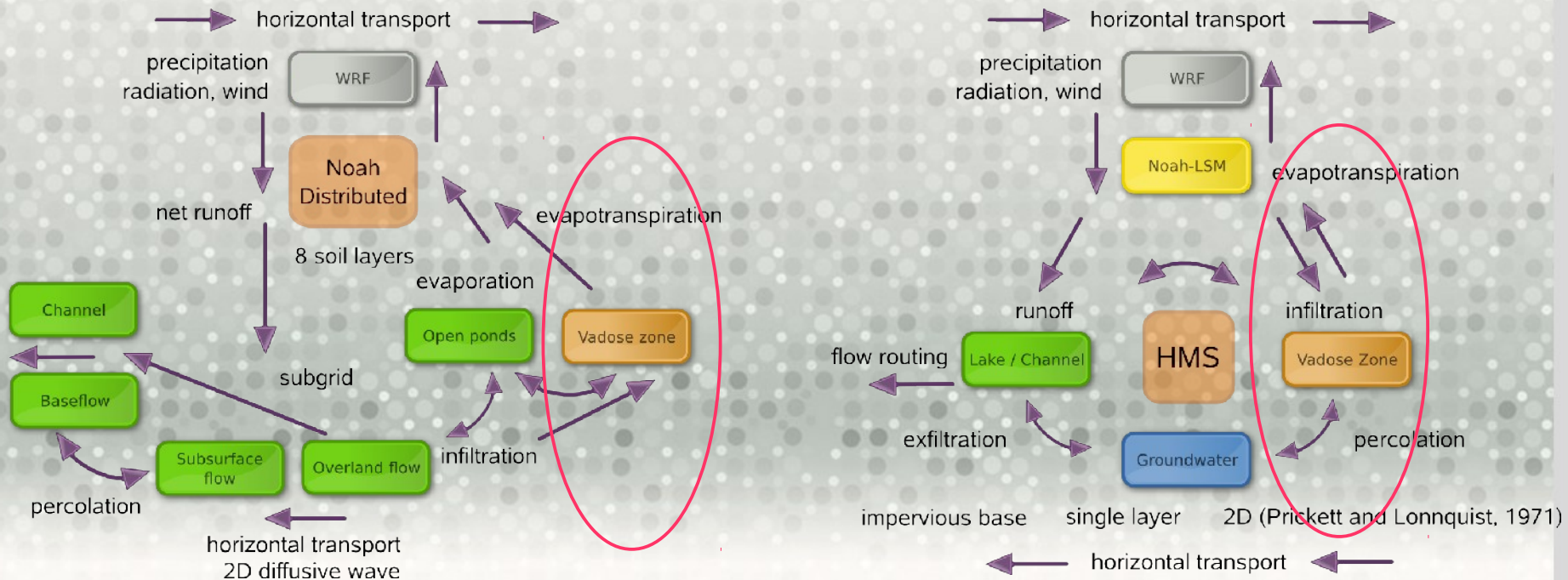
Mesh ~1km
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Weeks to Months

River Scale
>10,000 km²

HMS/
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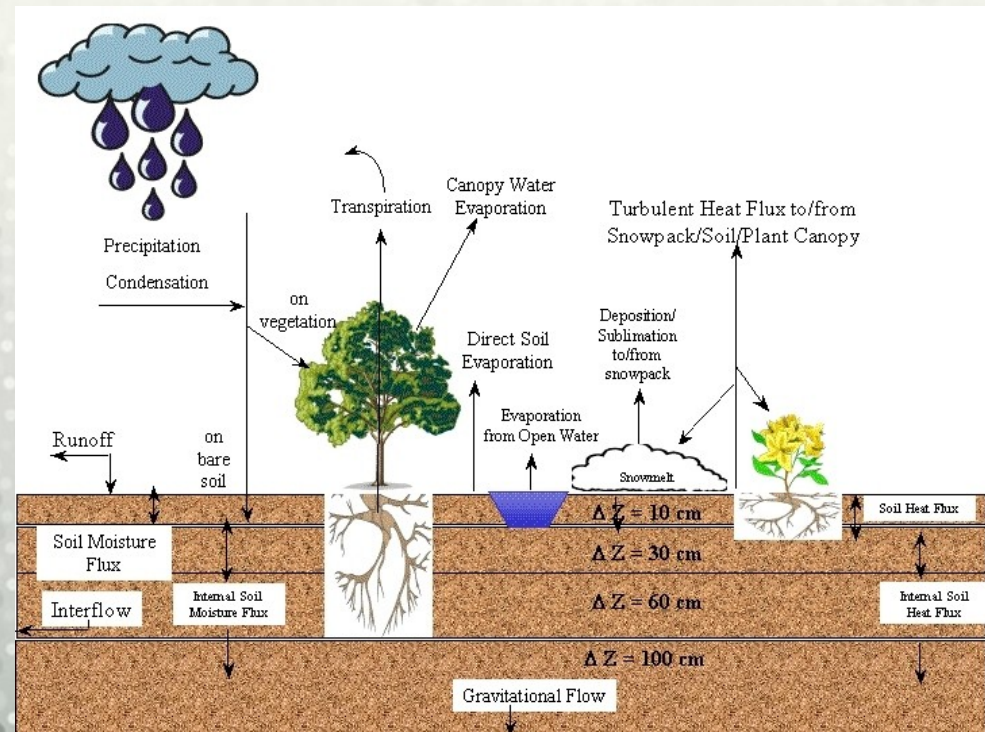
vs.



Common Land-Surface-Description in Regional Atmospheric Models

■ Standard column approach in LSMs

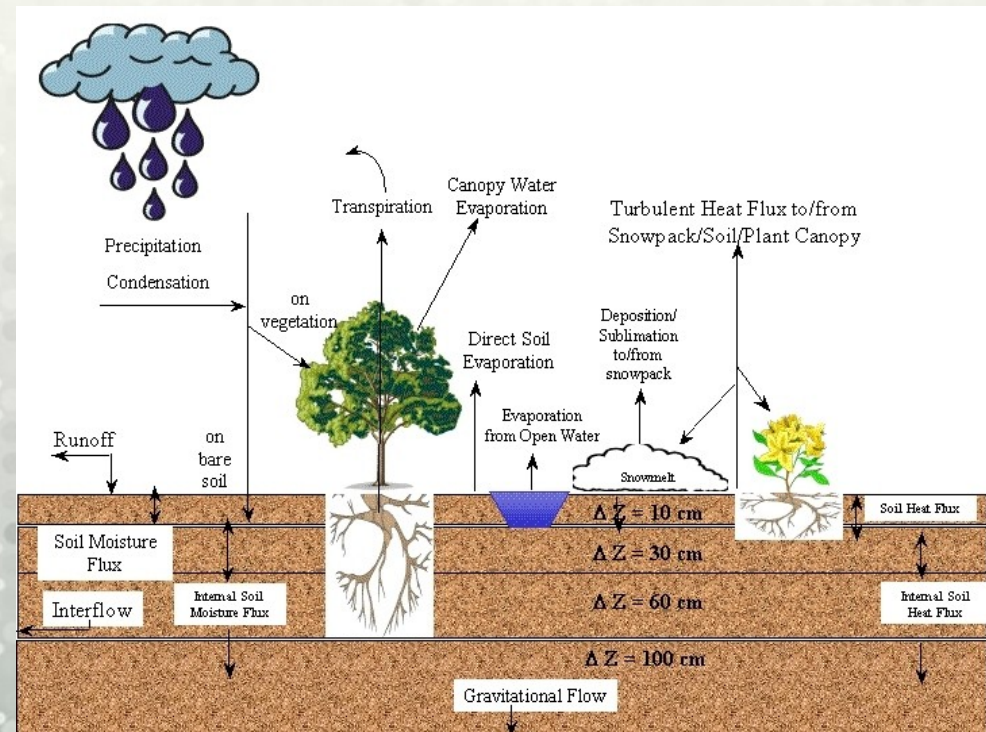
- Only vertical transport of moisture and energy is considered
- Infiltration excess and bottom drainage terms act as sink term with no possibility of returning
- Lower model boundary is defined by a downward gravitation driven flux



(Unified Noah / OSU Land Surface Model © NCAR)

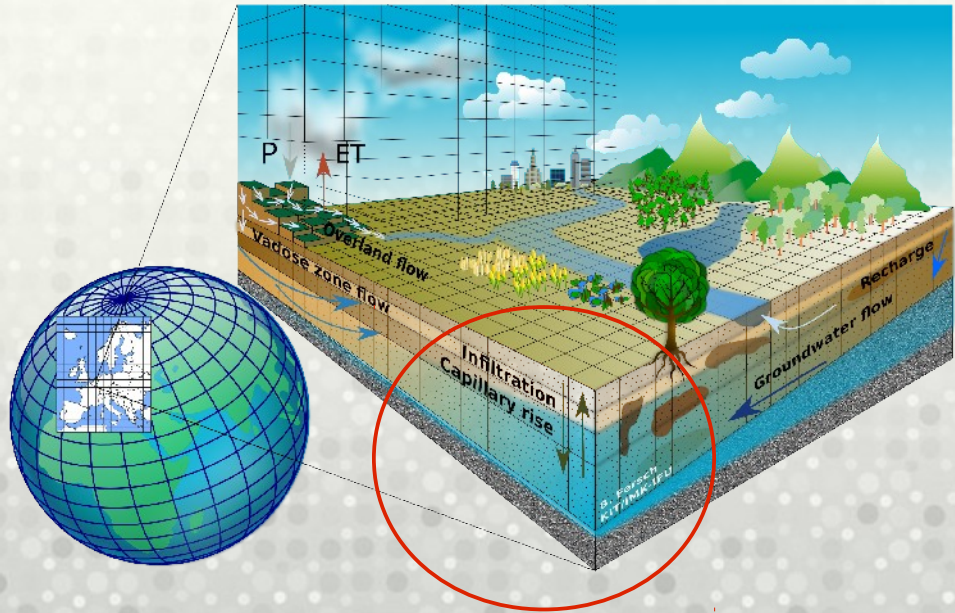
Common Land-Surface-Description in Regional Atmospheric Models

- Standard column approach in LSMs
 - No **lateral redistribution** of surface and subsurface water in the LSM
 - No **interaction** between **unsaturated** and **saturated** zone
 - **Mutual non-linear interaction of a closed water cycle cannot emerge**



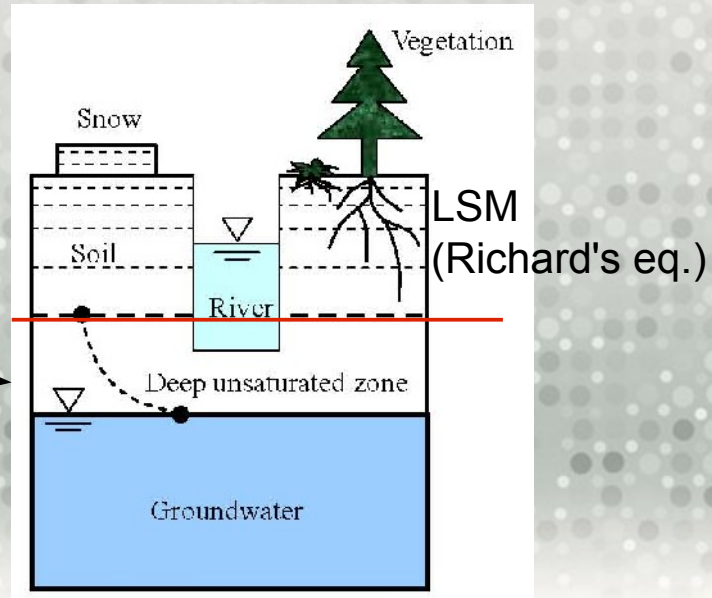
(Unified Noah / OSU Land Surface Model © NCAR)

Coupled atmosphere-hydrology model systems



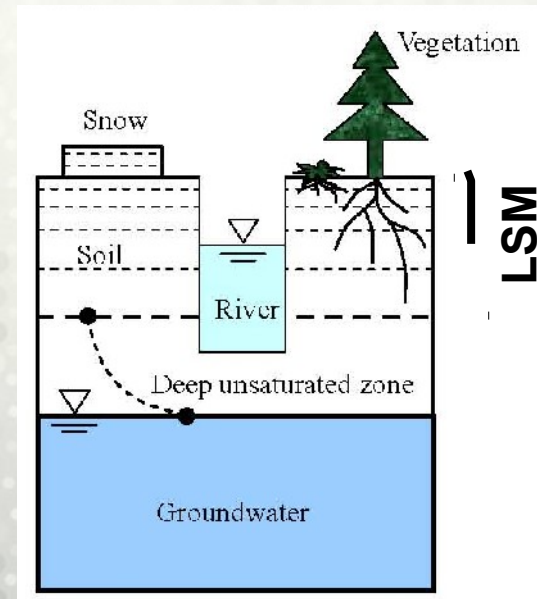
Description of moisture transport

GW-Head
Variable lower boundary



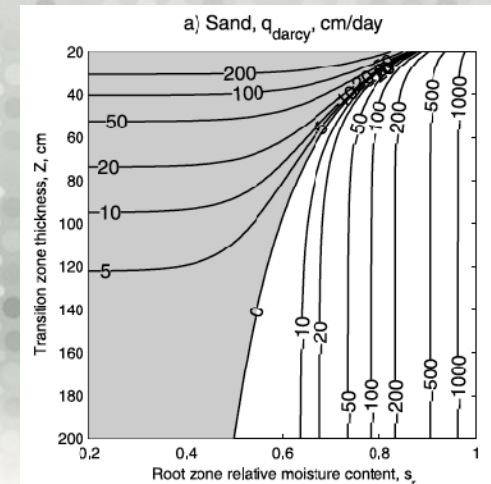
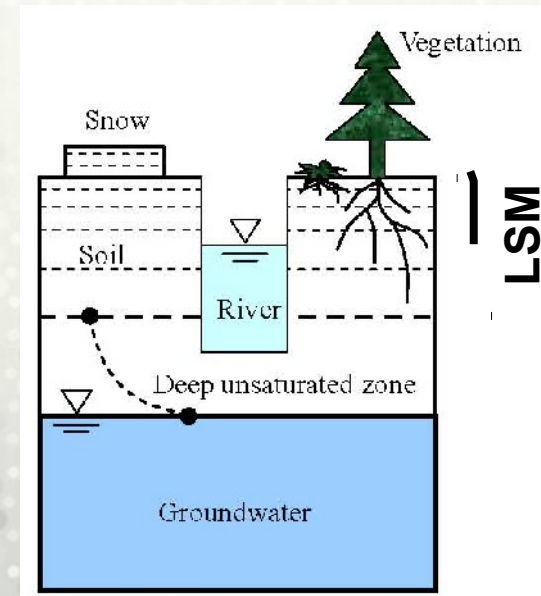
Groundwater Coupling Approach

- Approach 1: Richard's equation with fixed-head boundary condition (based on Zeng et al. (2009), De Rooij (2010))
 - Free drainage boundary condition of the LSM is replaced by bottom boundary condition which assumes an equilibrium soil moisture distribution of the deep unsaturated zone
 - The new bottom boundary condition is calculated by specifying the hydraulic head at the bottom of the LSM
 - New boundary condition realized with additional layer below the bottom of the Noah-LSM
 - Soil moisture content of additional layer depends on the distance to the hydraulic head
 - fully saturated, hydraulic head above or equal layer
 - partly saturated, hydraulic head within layer
 - unsaturated, hydraulic head below layer



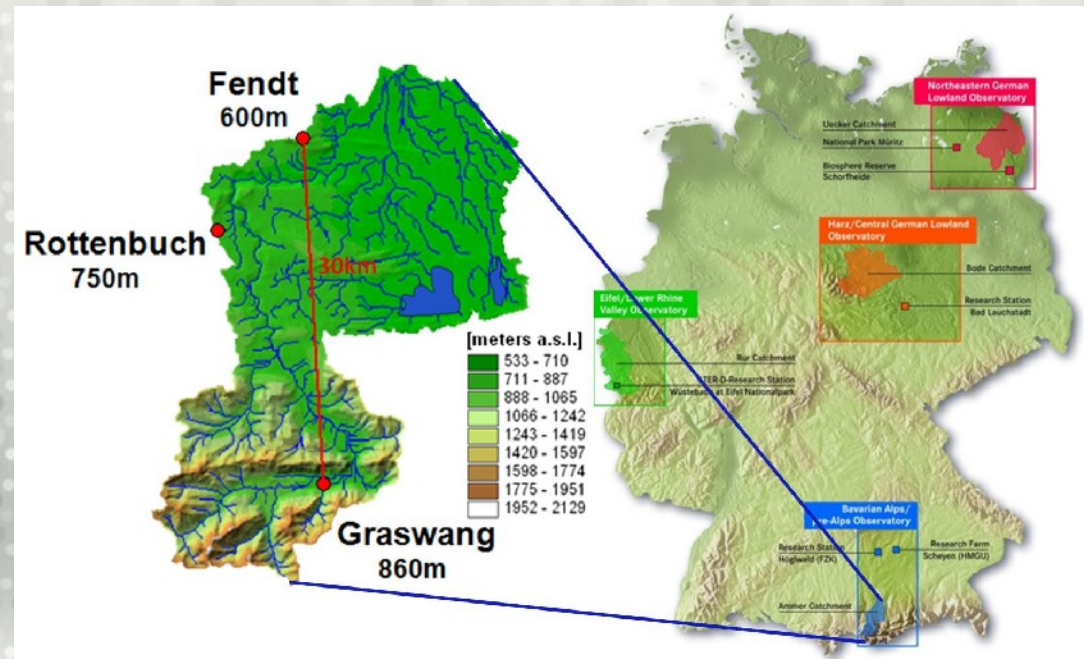
Groundwater Coupling Approach

- Approach 2: Darcy flux boundary condition (based on Bogaart et al. 2008)
 - Assumes a quasi steady-state moisture profile between groundwater head and the lowest soil layer of the LSM
 - Darcy equation is used to describe flow through this transition zone depending on the relative saturation at the bottom of the LSM and on the thickness of the transition zone
 - Parametrization that approximates net Darcy flux q_{darcy} for different thicknesses of the transition zone and different values of saturation for the lowest LSM soil layer
 - Approximation available for the 12 soil classes of Clapp & Hornberger (1978)

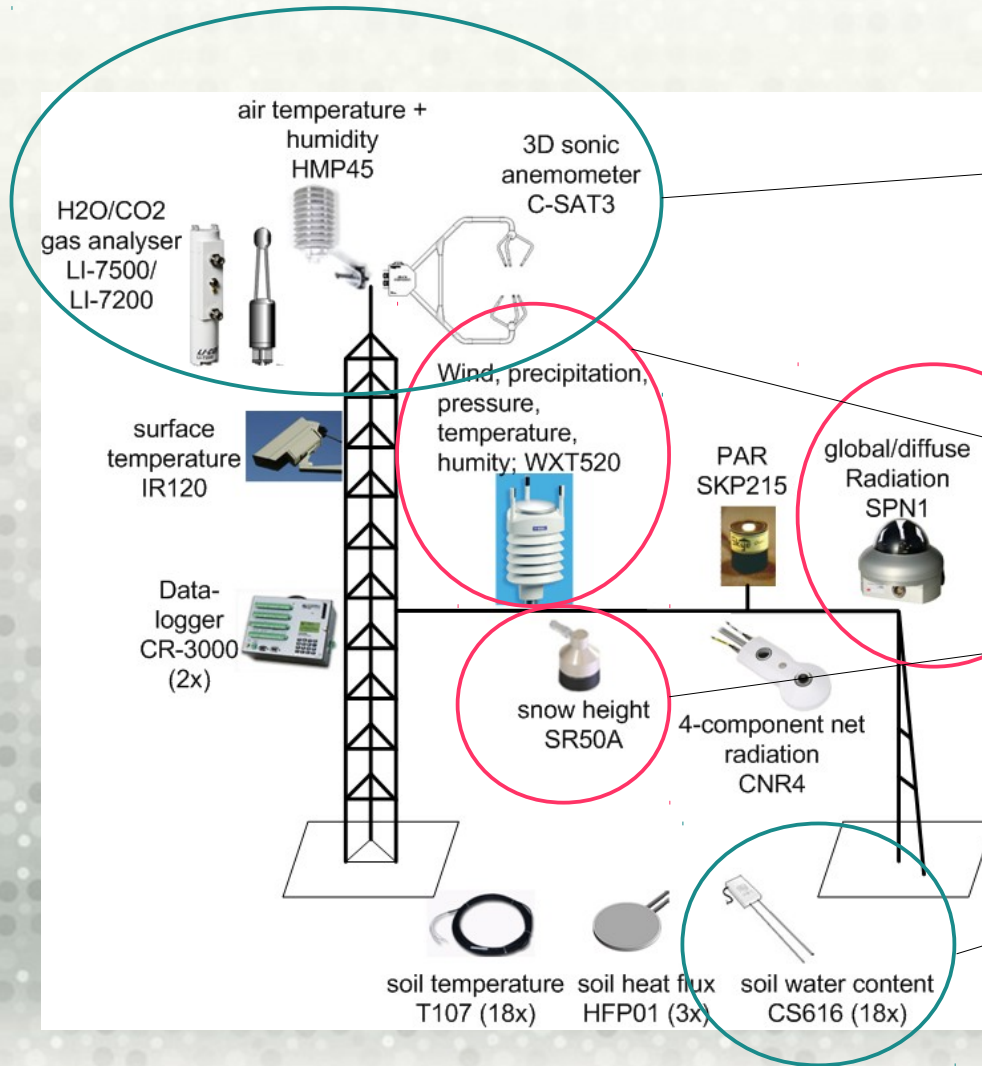


Proof of Concept – 1D Groundwater Coupling Study with the Noah-LSM and TERENO DATA

- The validity of the coupling approaches can hardly be analyzed in a distributed (2D) application
- 1D – stand-alone study is performed for a well measured site of the terrestrial environmental observatory (TERENO) pre-alpine
- Linder valley of the Ammer catchment
- Grassland, influenced by temporarily shallow groundwater table



Data at TERENO Climate Station Graswang



Time-series of eddy-flux evapotranspiration for evaluation

Noah-LSM stand-alone model driving

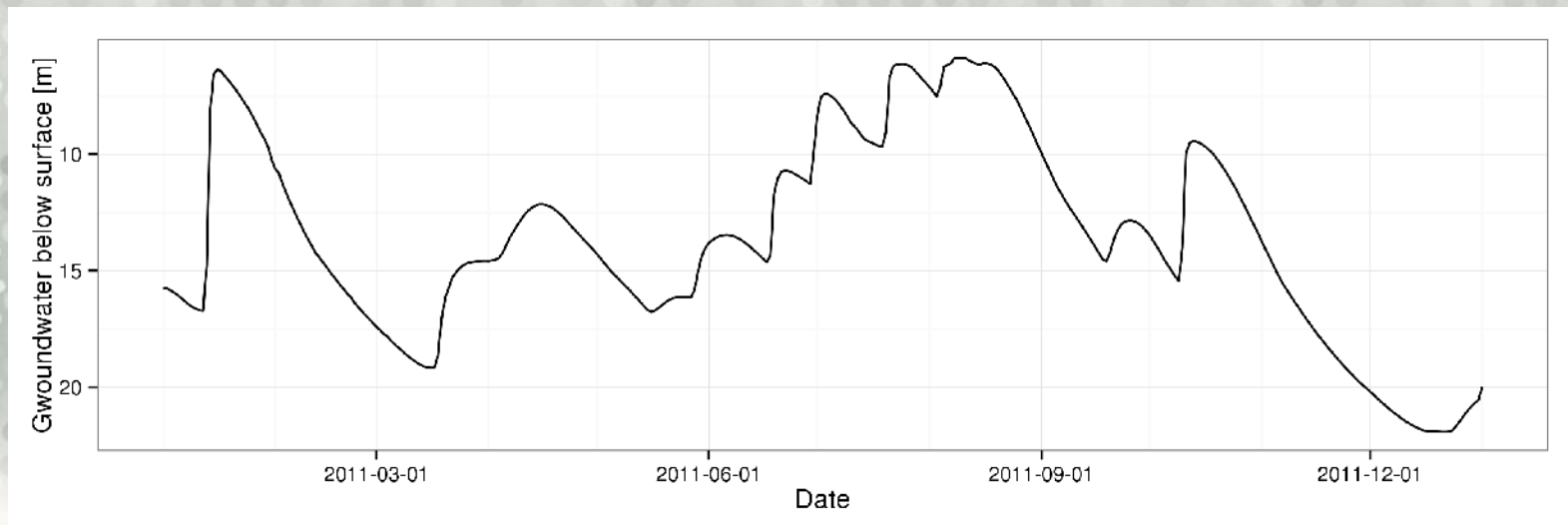
Time-series of soil-moisture for evaluation

Data Source: TERENO Site Graswang

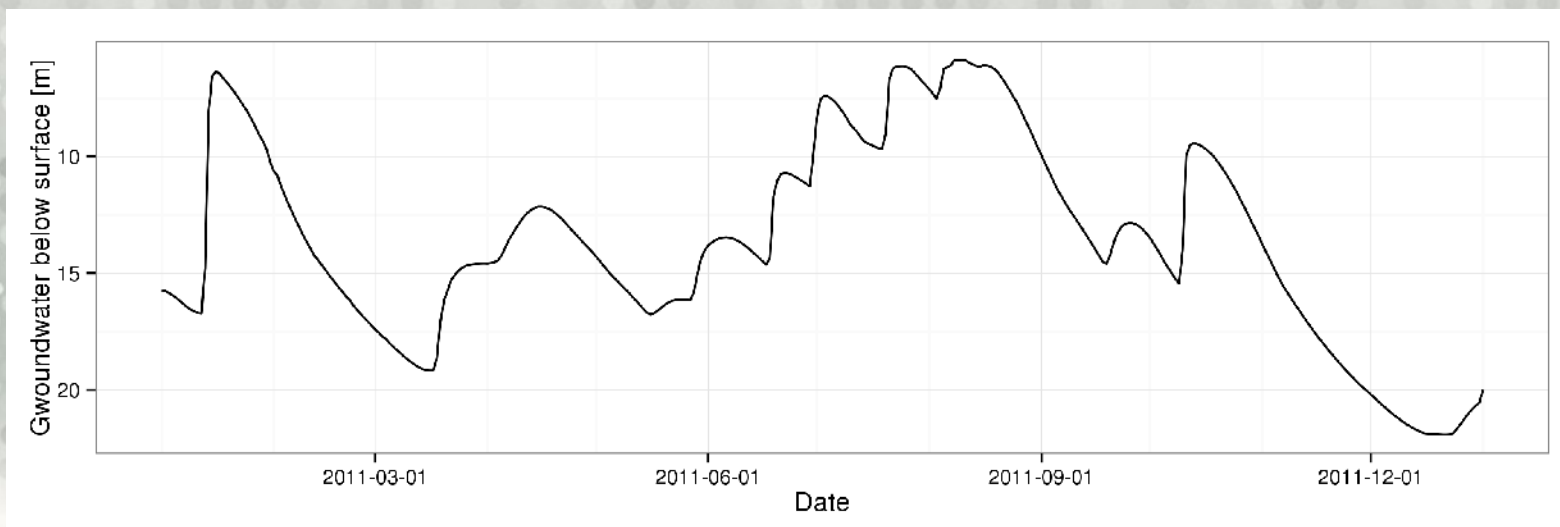
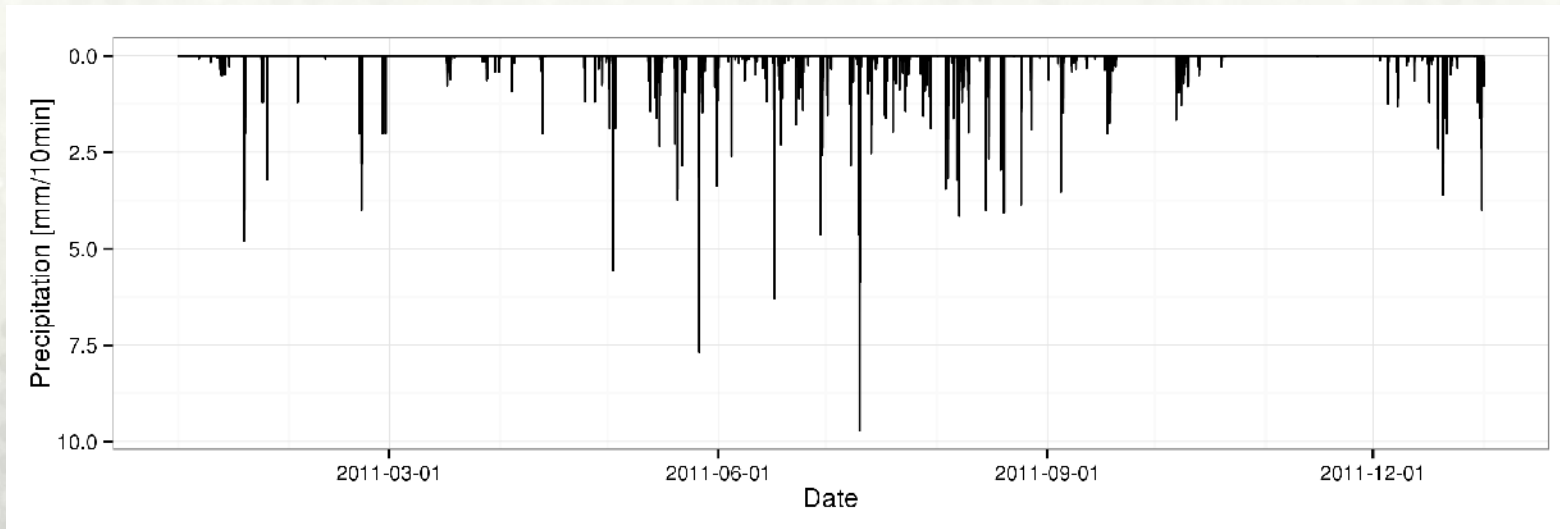


TERENO
TERRESTRIAL ENVIRONMENTAL OBSERVATORIES

Data Source: Groundwater Head Time-Series from the Local Water Authority (WWA Weilheim)



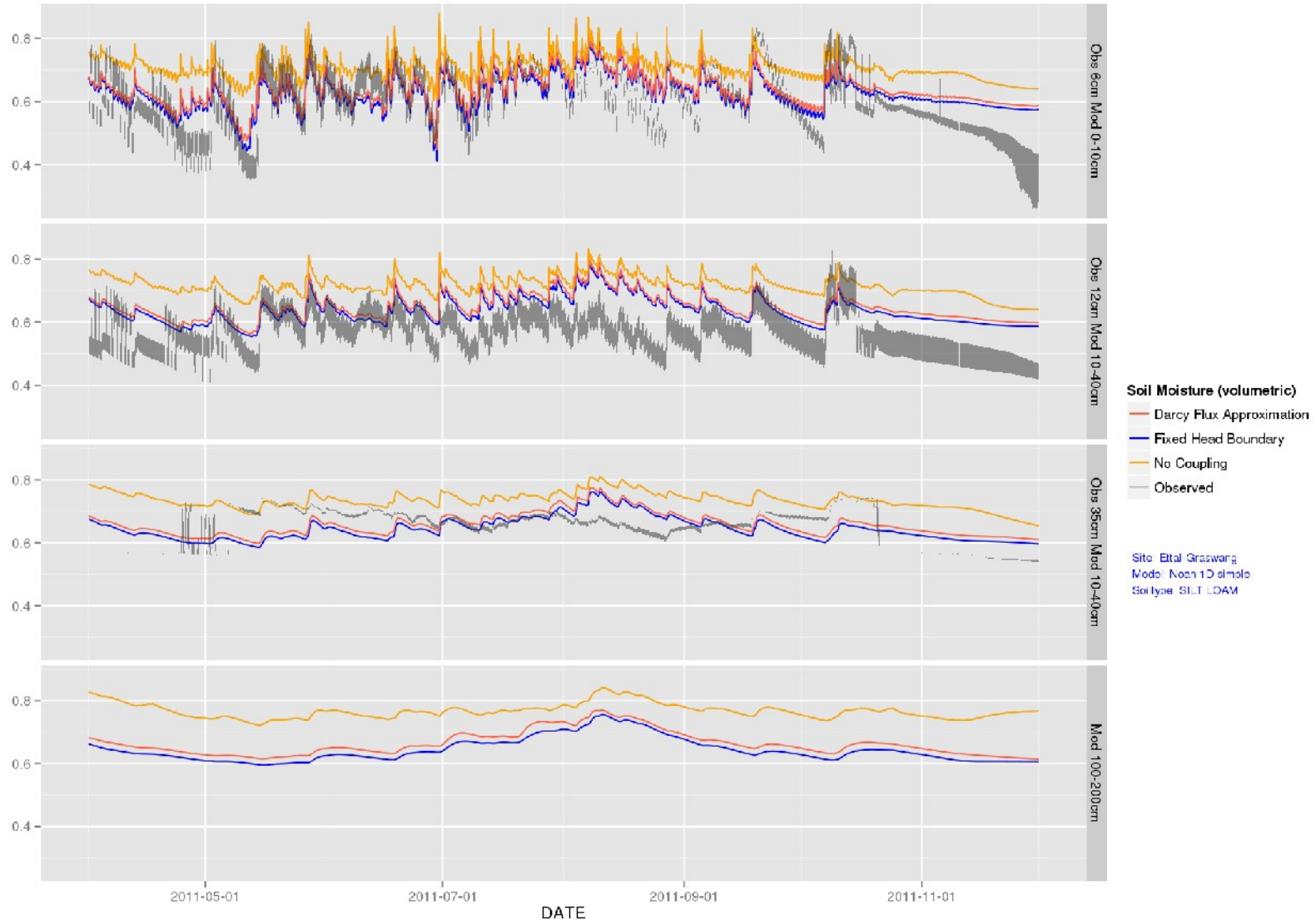
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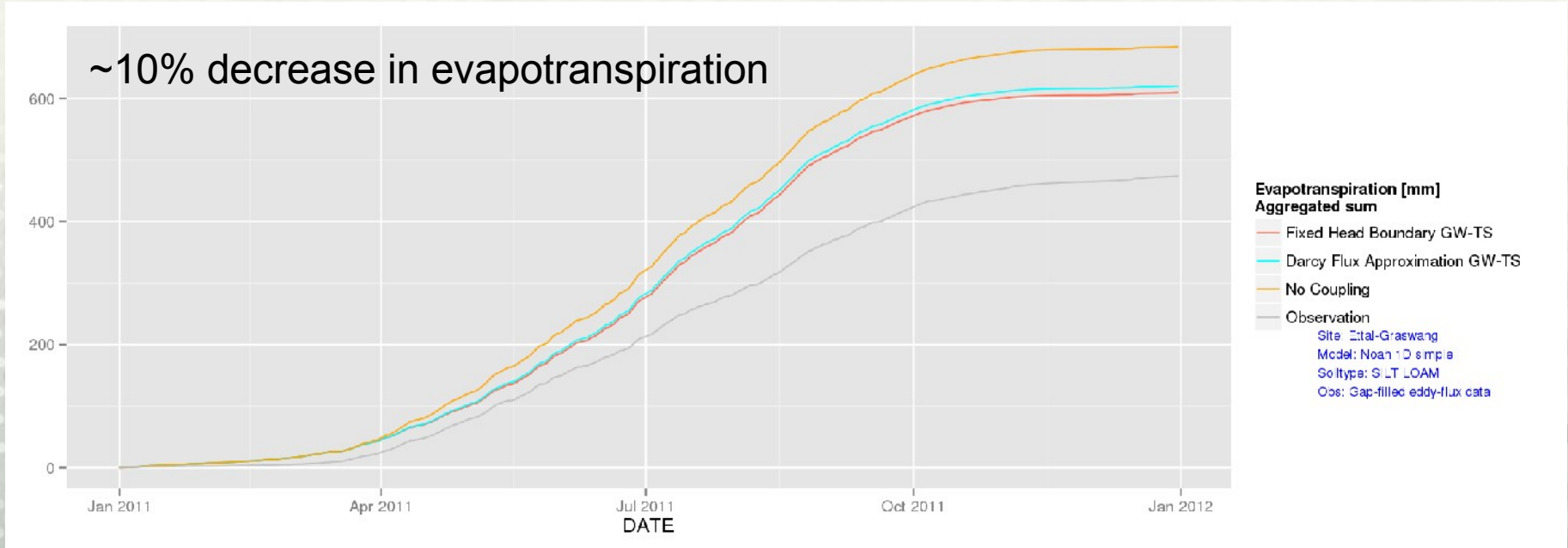
Configuration Noah-LSM for 1D Study

- Noah-LSM 1D, simple-driver, V.3.3
- Simulation period: 2010-07-01 to 2011-12-31 (half year of spin-up)
- Simulation time-step: 10 minutes
- Soil-type 4, silt loam
- Vegetation / land-use: grassland (USGS 7, modified)
- Monthly update of albedo, LAI, vegetation fraction, and roughness length
- 4 soil layers, 0-10, 10-40, 40-100, 100-200 cm (standard setup)
- Modes: Uncoupled, fixed-head-boundary and darcy-flux coupling with time-variant and constant (-7m) groundwater head

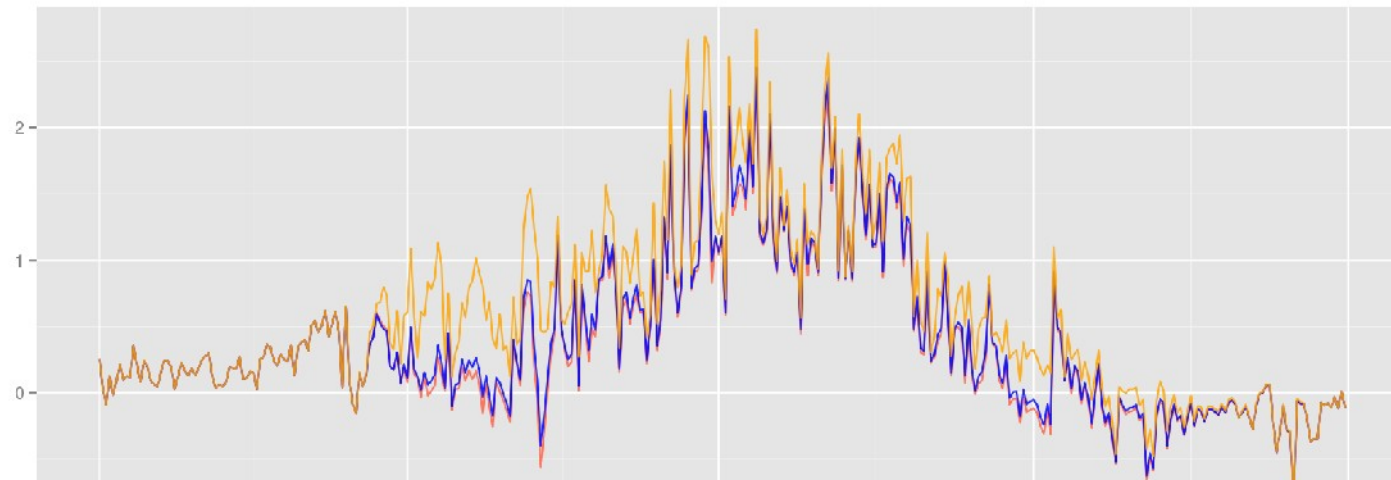
Impact of Coupling on Soil Moisture Dynamics



Impact of GW-Coupling on Evapotranspiration



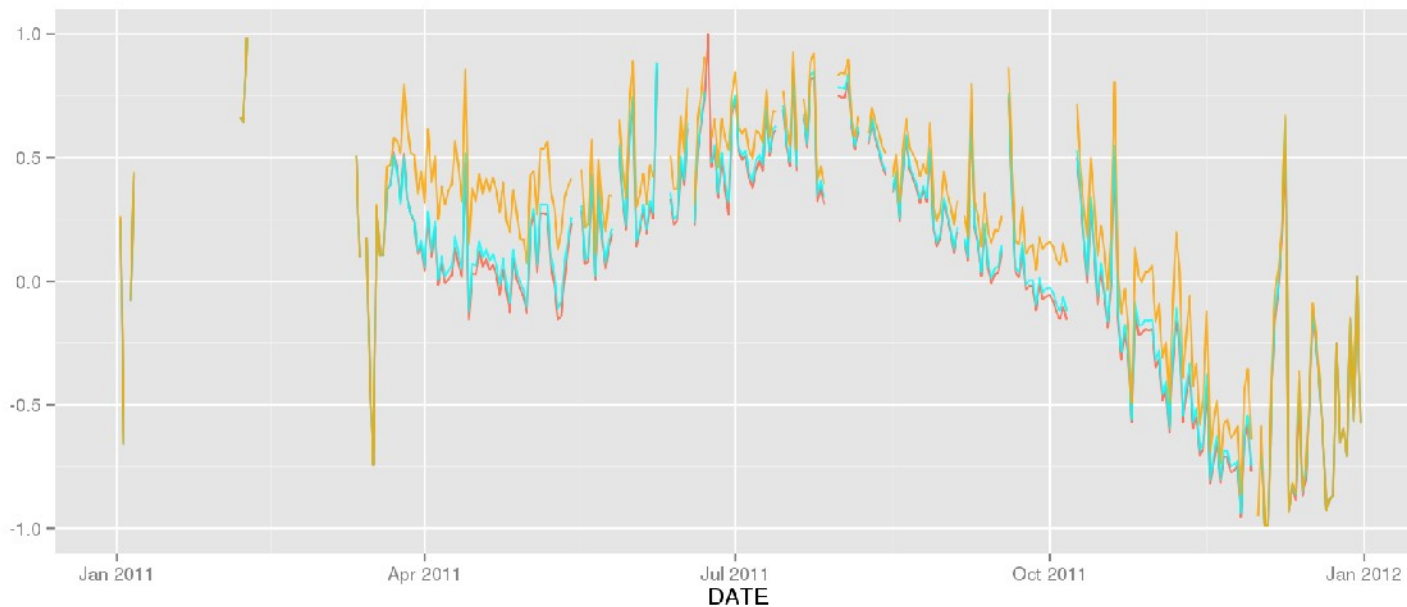
Daily Bias of Evapotranspiration



Evapotranspiration [mm/day]
Sim-Obs

- Fixed Head Boundary
- Darcy Flux Approximation
- No Coupling

Site: Etal-Grasweg
Model: Nash 1D simple
Soiltype: SILT LOAM
Obs: Gap-filled eddy-flux data



Evapotranspiration [mm/day]
(Sim-Obs)/Obs (rel. bias)

- Fixed Head Boundary GW-TS
- Darcy Flux Approximation GW-TS
- No Coupling

Site: Etal-Grasweg
Model: Noah 1D simple
Soiltype: SILT LOAM
Obs: Gap-filled eddy-flux data

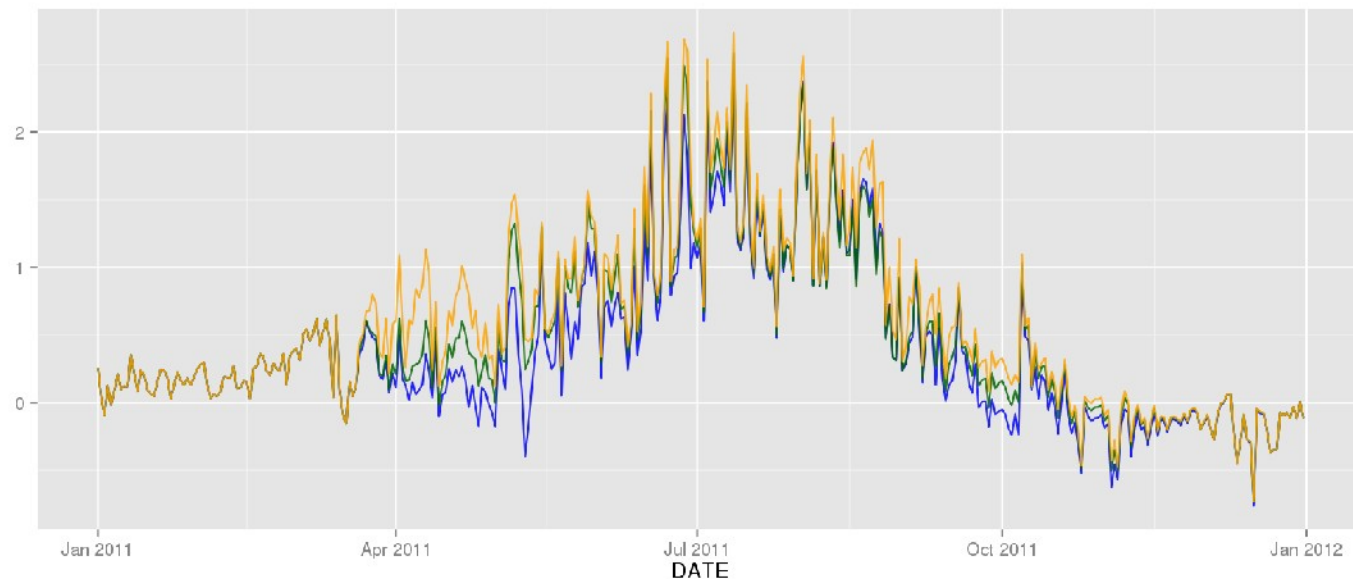
ETa, Comparison Static vs. Variable GW - Head



Evapotranspiration [mm/day] Sim-Obs

- Fixed Head Boundary GW-TS
- Fixed Head Boundary GW 7m
- No Coupling

Site: Etal-Graswang
Model: Noah 1D simple
Soiltype: SILT_LOAM
Obs: Gap filled eddy flux data

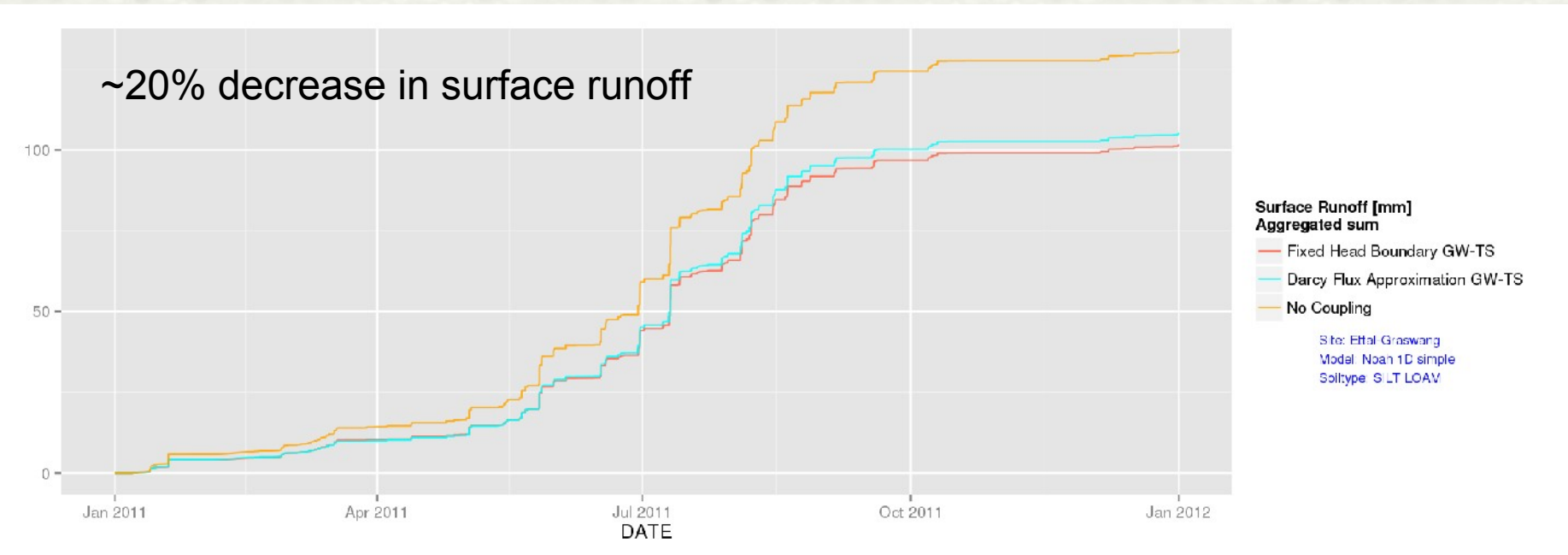


Evapotranspiration [mm/day] Sim-Obs

- Darcy Flux Approximation GW-TS
- Darcy Flux Approximation GW 7m
- No Coupling

Site: Etal-Graswang
Model: Noah 1D simple
Soiltype: SILT_LOAM
Obs: Gap filled eddy flux data

Impact of GW-Coupling on Infiltration Excess



Distributed Model Systems WRF-Hydro & WRF-HMS

Headwater Scale
<10,000km²

WRF-Hydro
(NDHMS)

Mesh ~1km
Subgrid ~100m
Weeks to Months

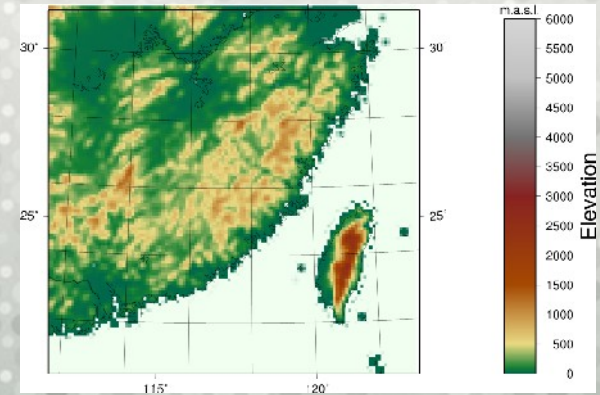


Process Study and Forecast Application

River Scale
>10,000 km²

HMS/
Noah-LSM/
WRF

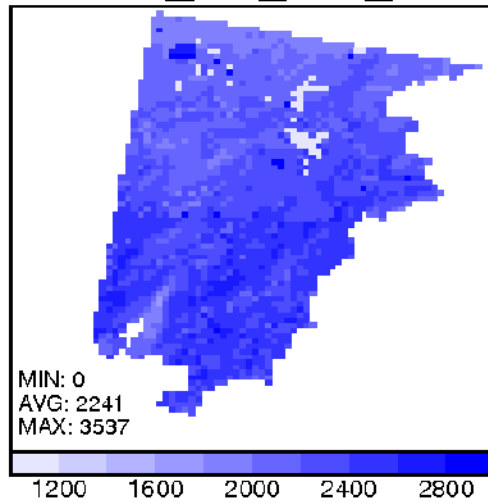
Mesh ~10km
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Years to Decades



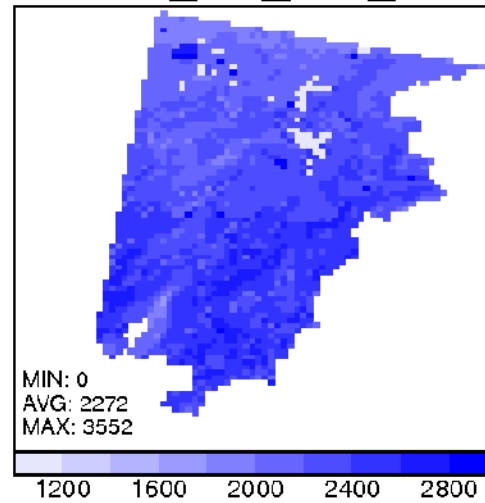
Climate and Long Term Application

Mean yearly sums [mm/year], EVAP

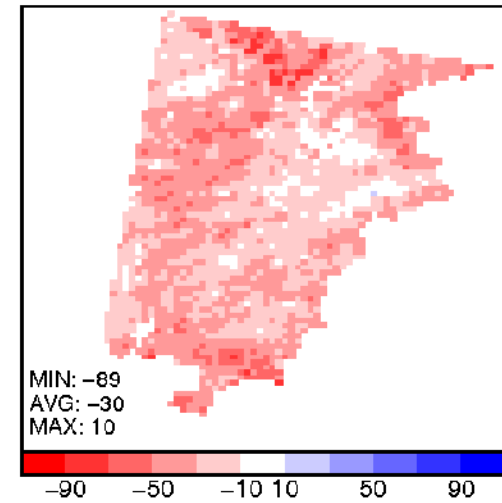
EVAP_DS_mm_M0



EVAP_DS_mm_M4



DIFF

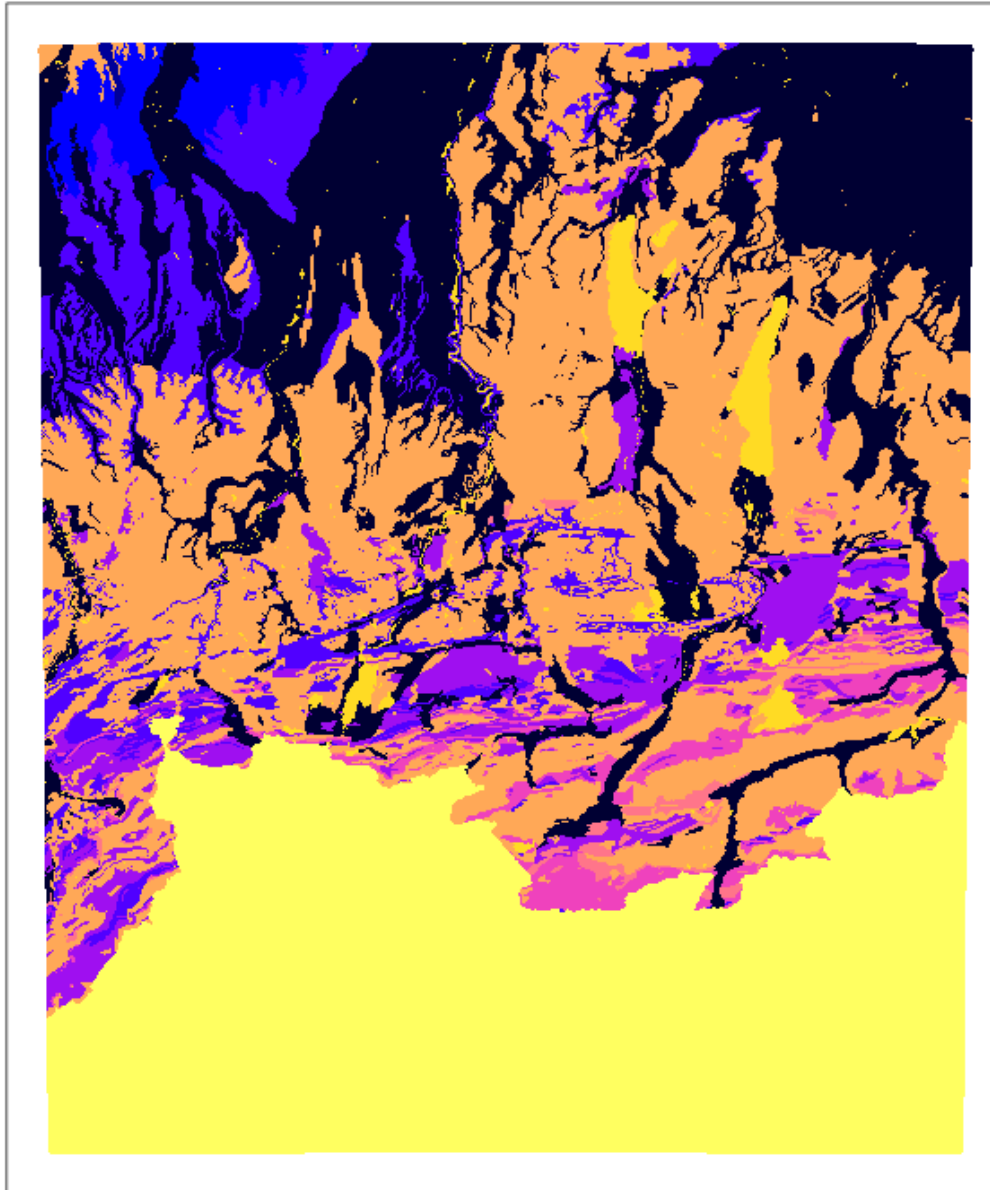


poyang_EVAP_DS_mm_M0_EVAP_DS_mm_M4_EVAP_yearsum_ymonmean_mrf.eps

Difference in potential evaporation for groundwater coupled HMS-WRF simulation.

- Lower boundary condition for LSM Richard's equation important for simulation of near-surface soil moisture field
- For 1D column study with TERENO input data, improvement of soil moisture states of the upper LSM soil levels
- Groundwater coupling impacts also evapotranspiration and surface runoff with considerable amounts

- Additional comparison of 1D Study for other TERENO sites throughout Germany (but groundwater is not monitored for most of the observatories)
- Application of a groundwater enabled version of the model system WRF-Hydro for the Ammer catchment (incl. 2D groundwater model)
- Distributed application requires distributed hydrogeological information about saturated conductivity, porosity, and thickness for aquifers → approximation strategy
- Validation of coupled simulations with respect to soil moisture patterns (sensor networks, ACROSS, remote sensing products)



Classes of hydraulic conductivity

- 2
- 3
- 4
- 5
- 6
- 9
- 10
- 12
- 99
- 100